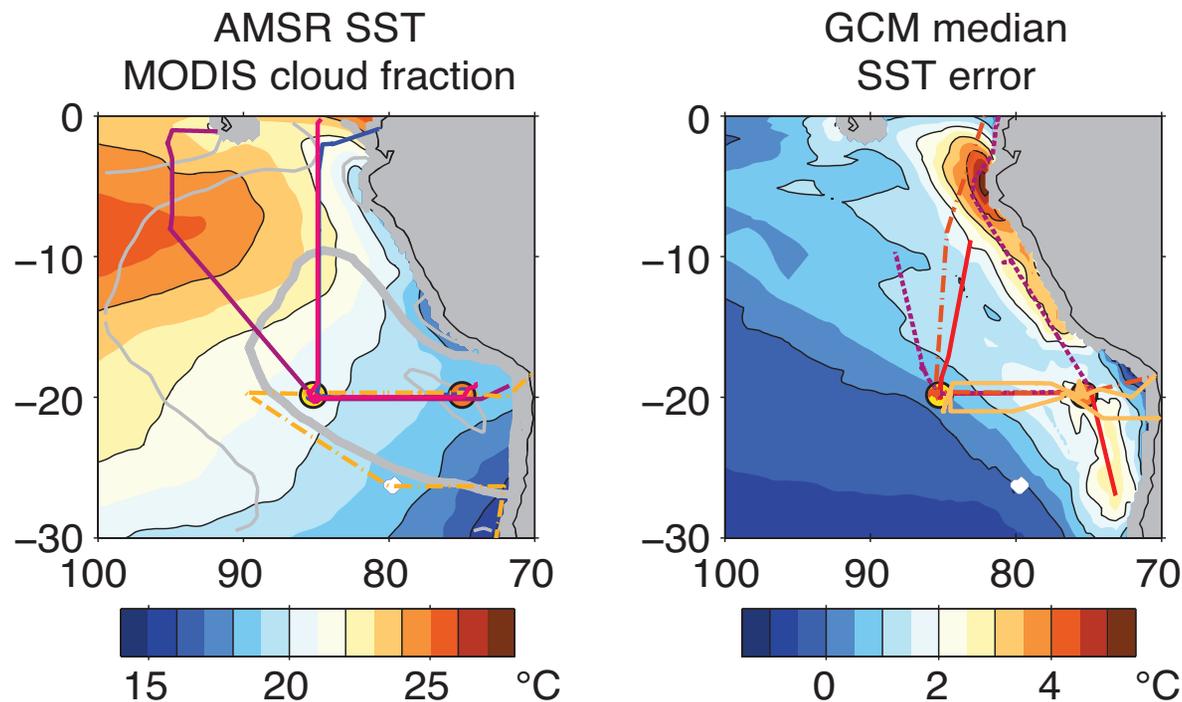


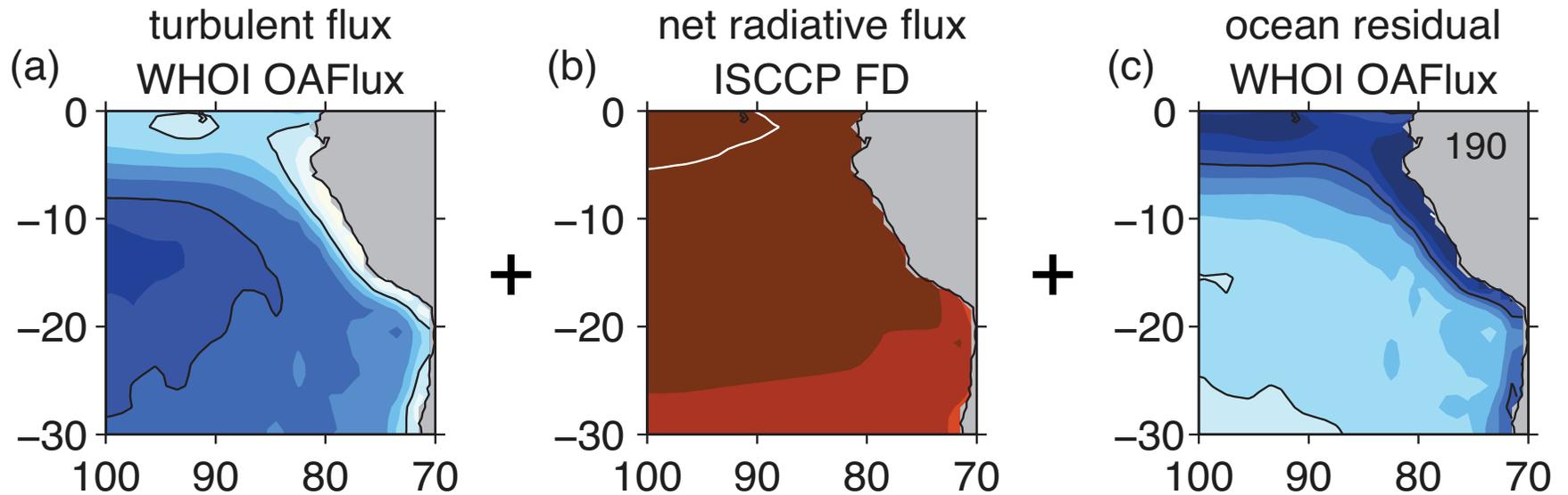
Eastern Pacific surface fluxes observed and simulated by coupled GCMs



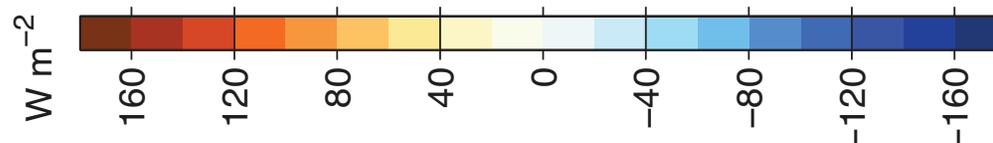
Simon de Szoeke & Chris Fairall

de Szoeke et al. 2010: Surface Flux Observations on the Southeastern Tropical Pacific Ocean and Attribution of SST Errors in Coupled Ocean–Atmosphere Models, *J. Climate*.

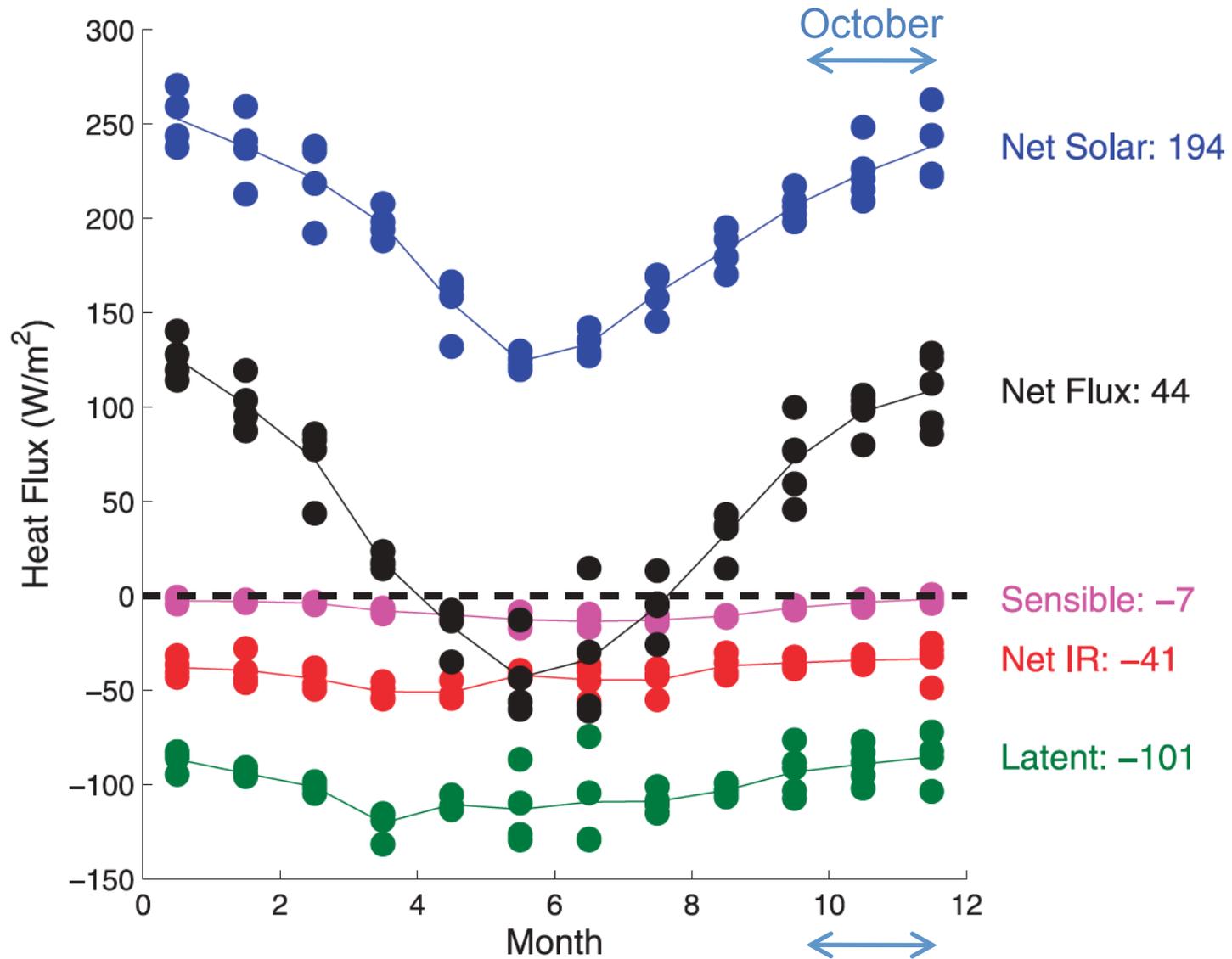
annual average heat budget



= 0



surface ocean heat budget

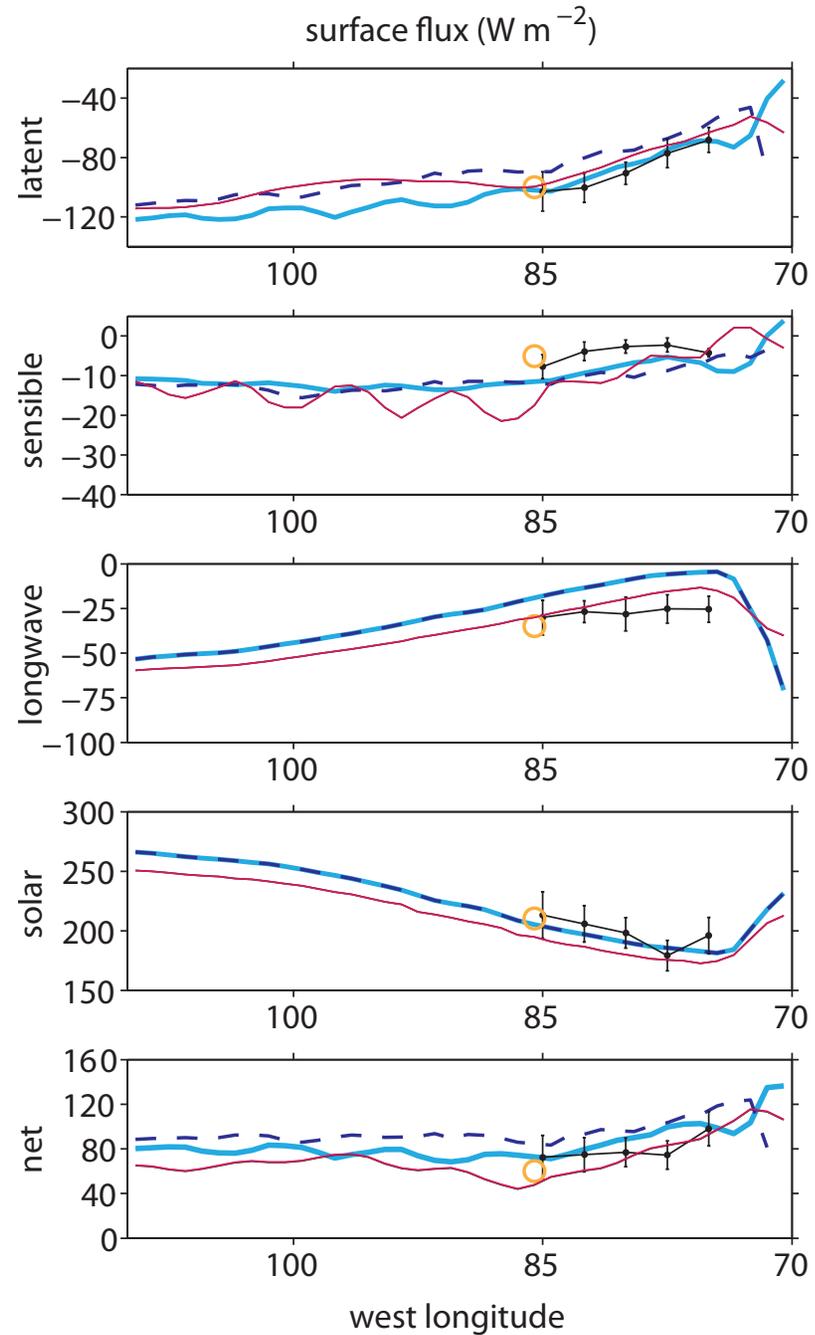


Colbo and Weller 2007

ship-observed heat fluxes

20°S, October

- WHOI OAFlux
- - UW Hybrid
- NCAR CORE
- NOAA PSD ship
- WHOI buoy



surface heat budget

0 =

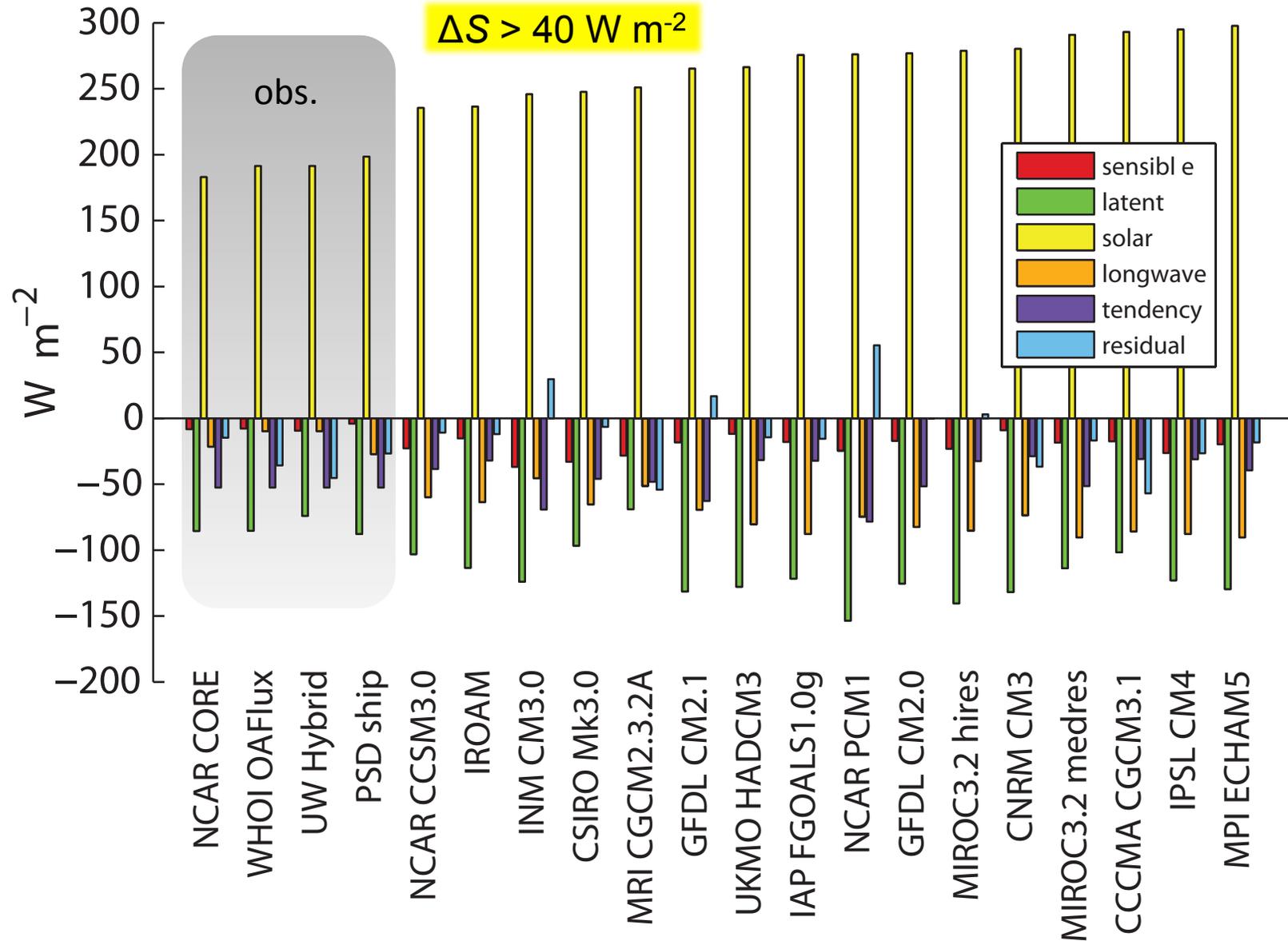
$-\partial/\partial t$ SST

turbulent evaporation and sensible flux

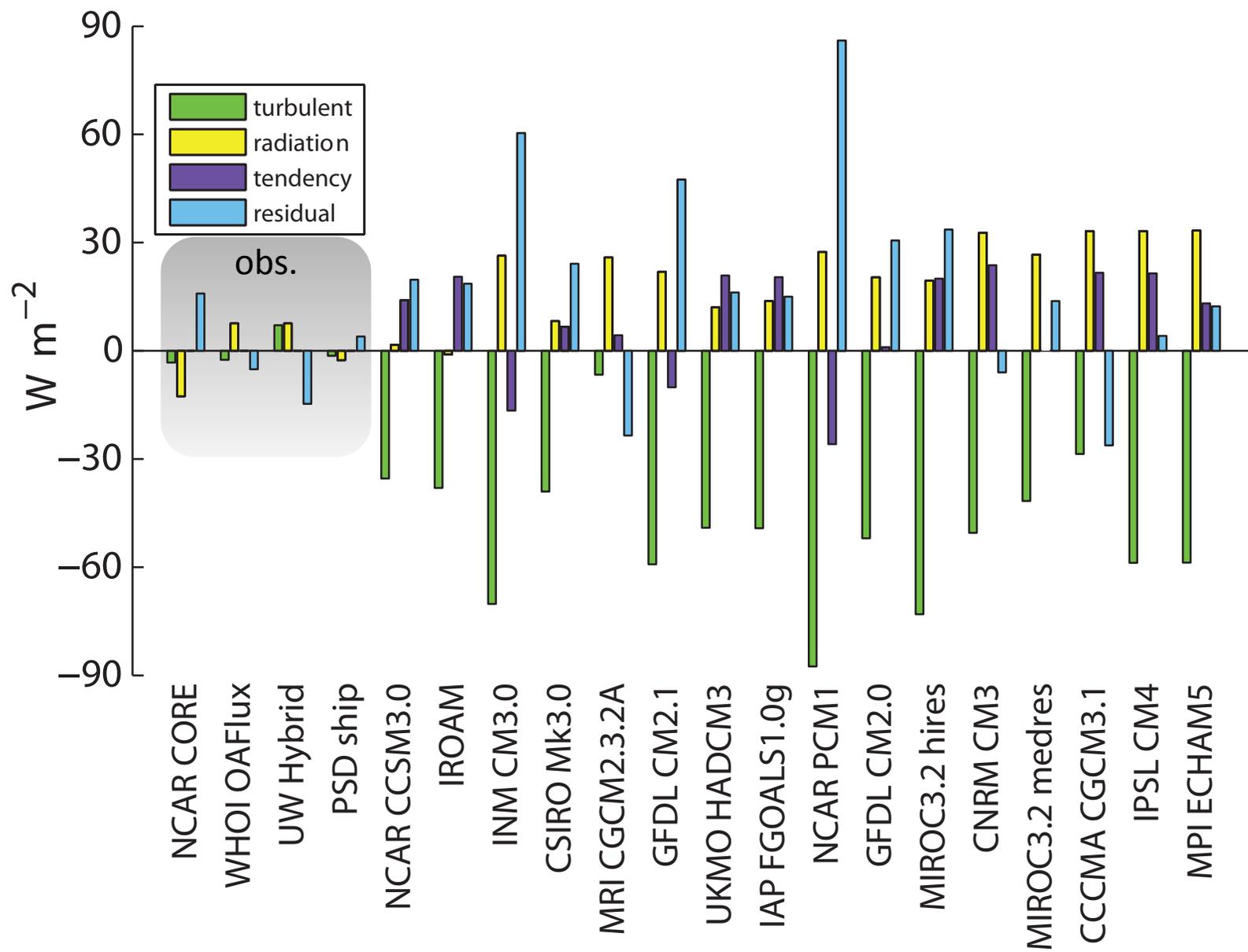
net radiation

ocean residual

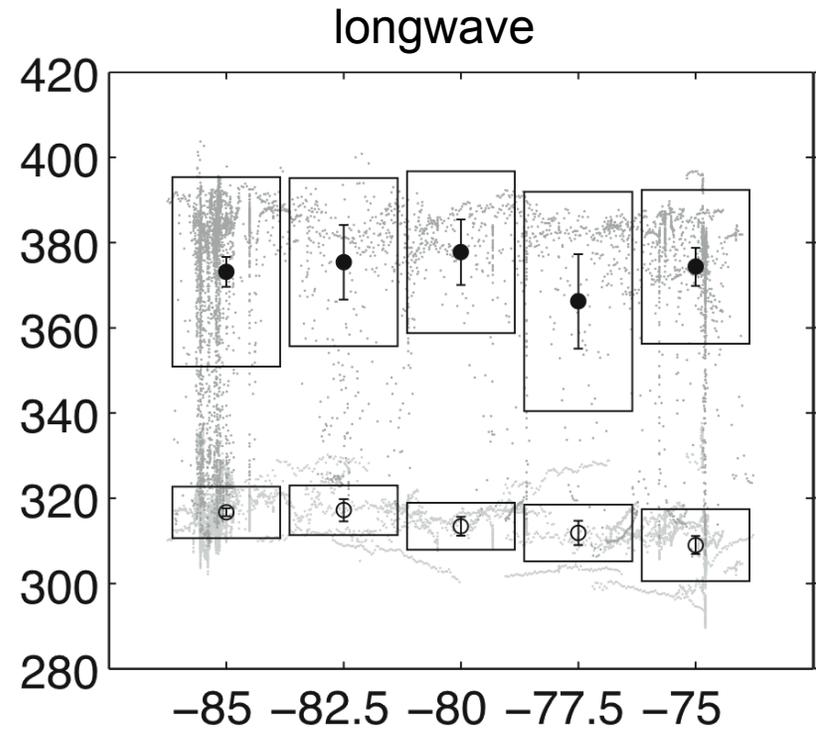
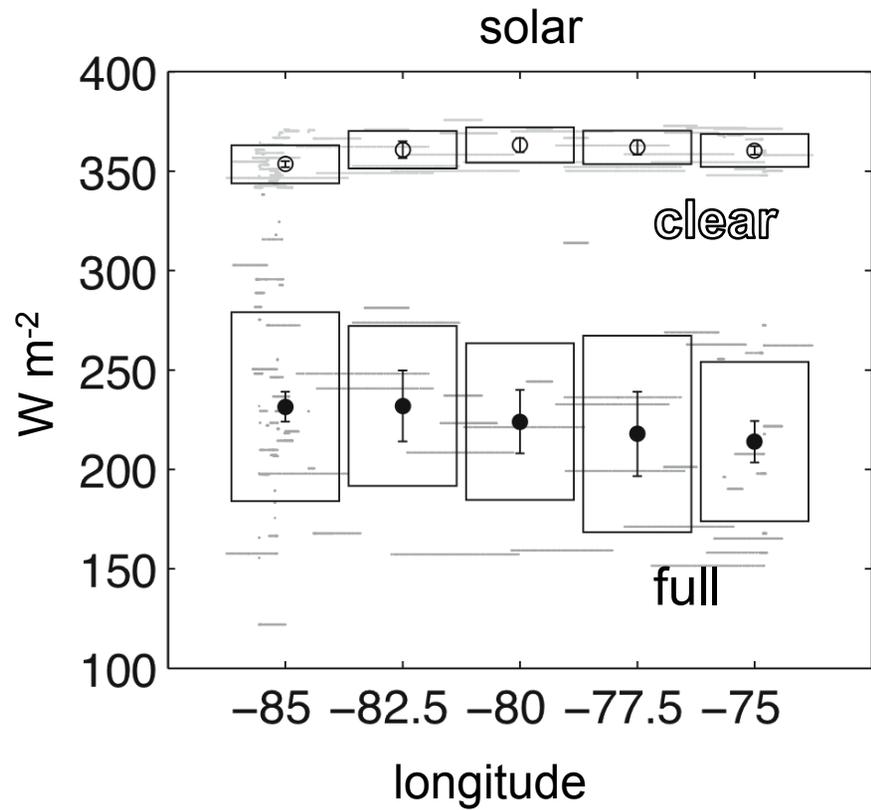
surface heat budget



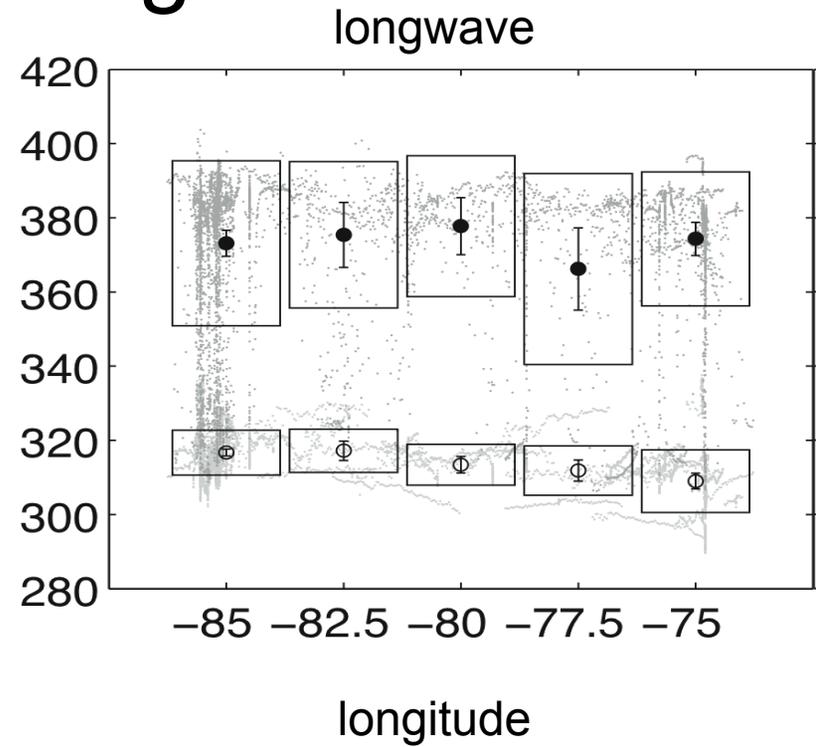
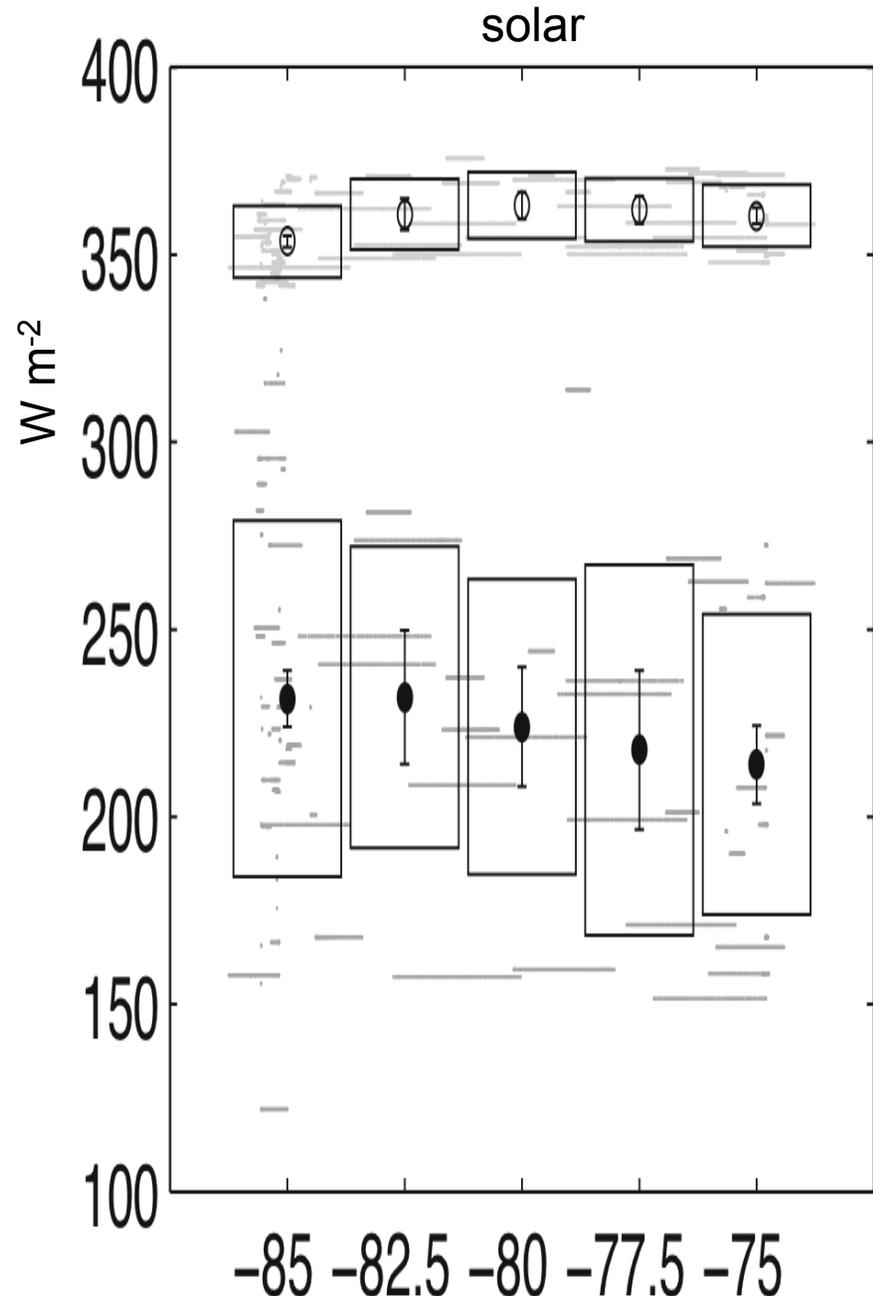
surface heat budget *errors*



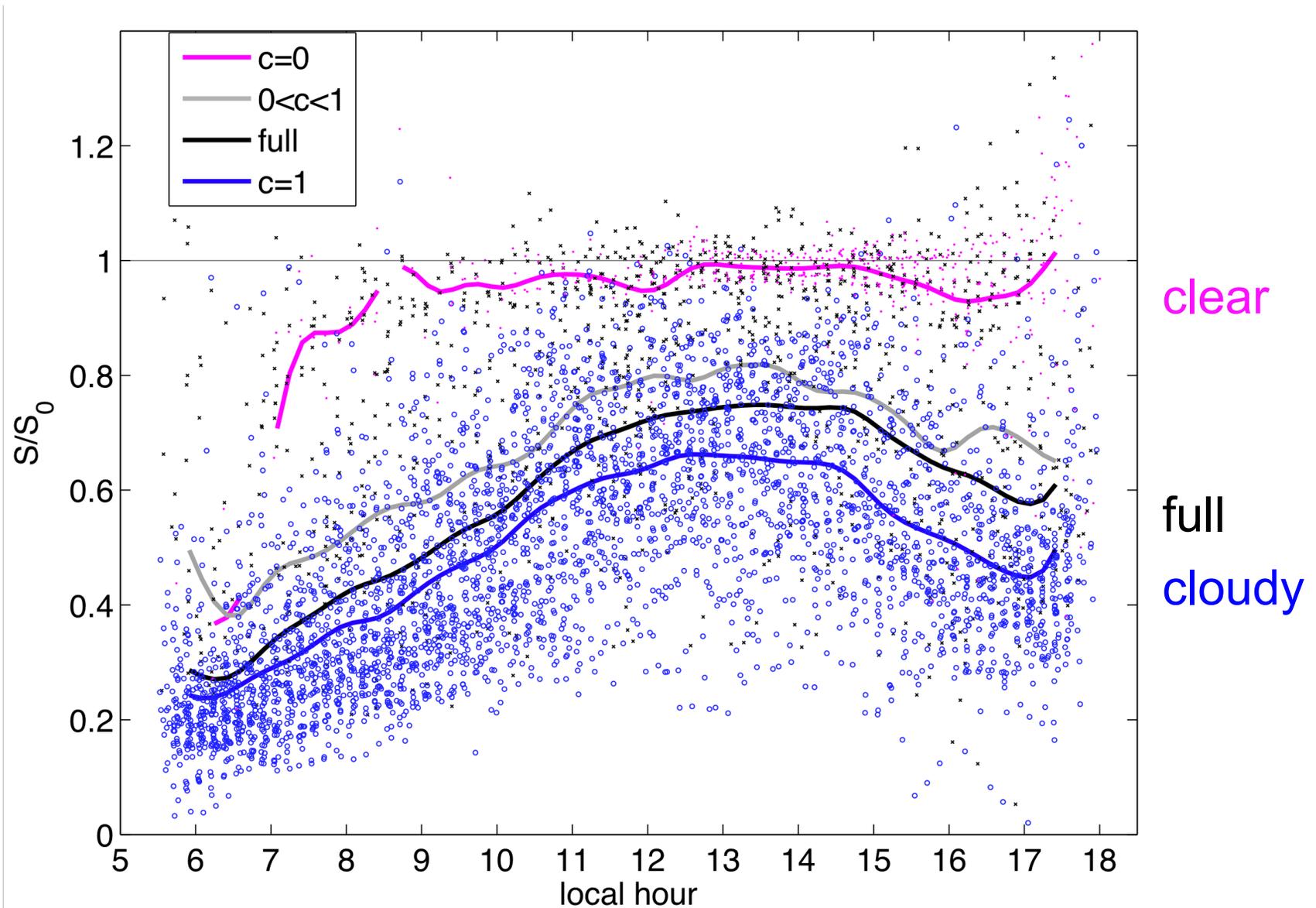
surface downwelling radiation



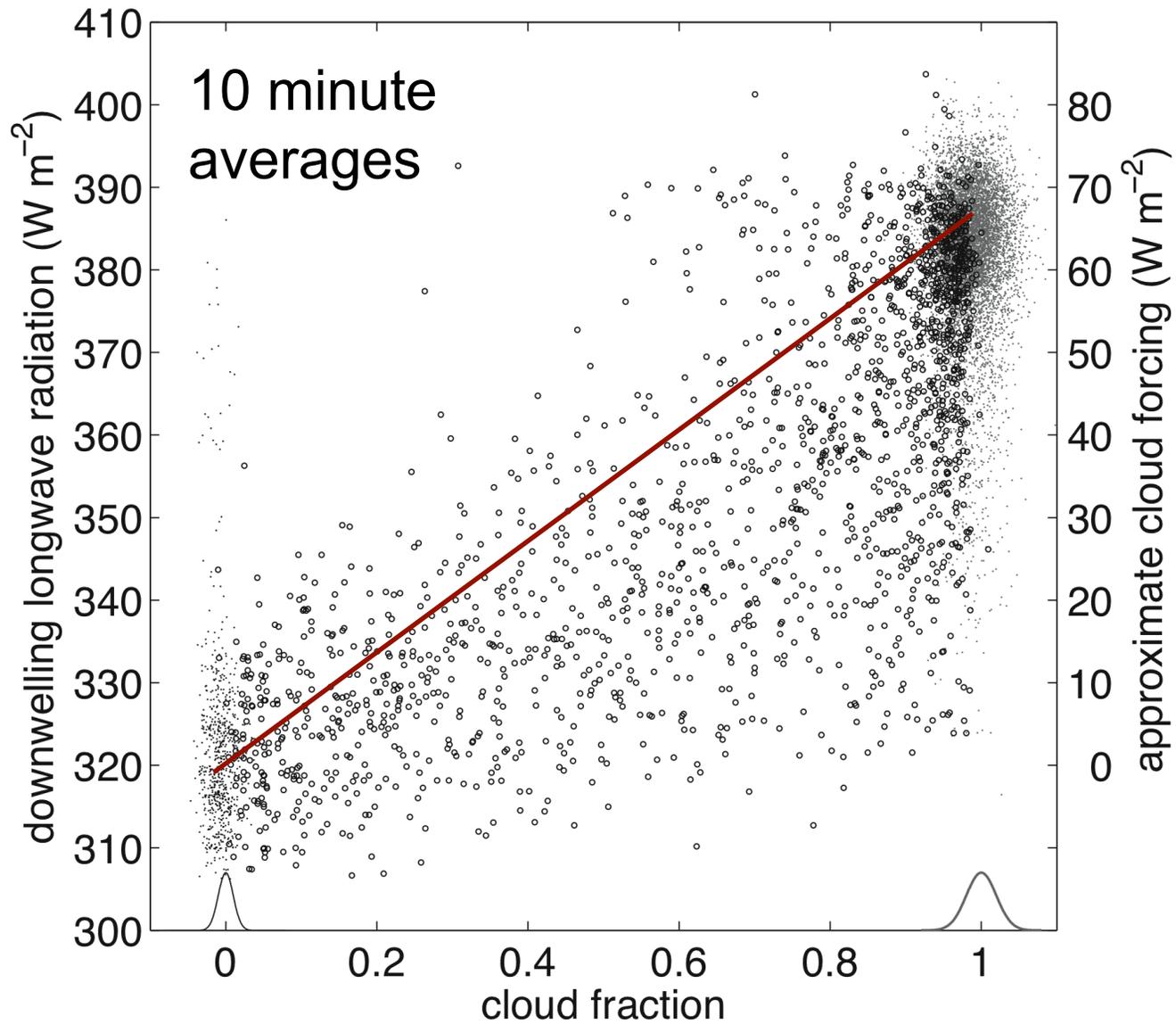
surface downwelling radiation



solar transmission



longwave cloud effect

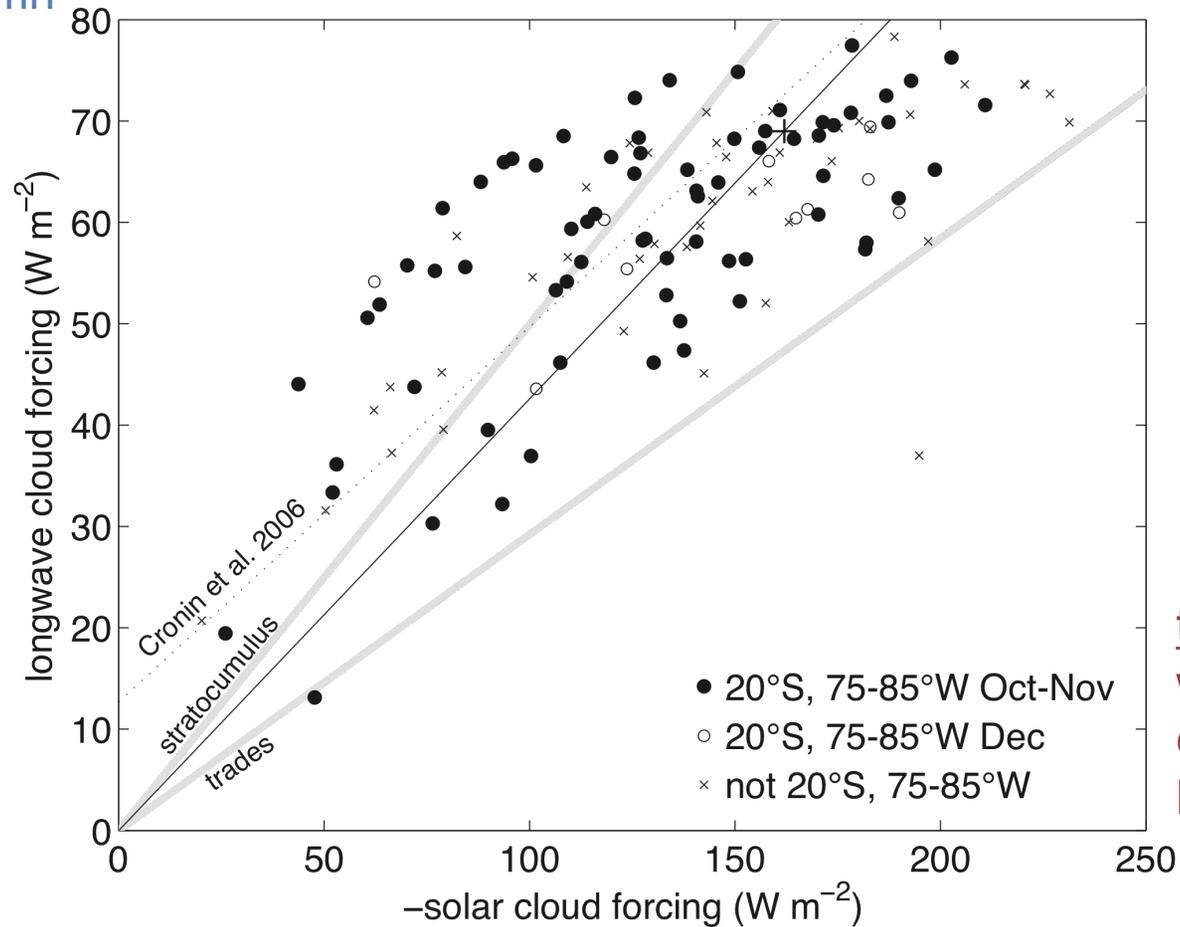


Cloud forcing phase diagram

polar

cold dry column

low sun ←

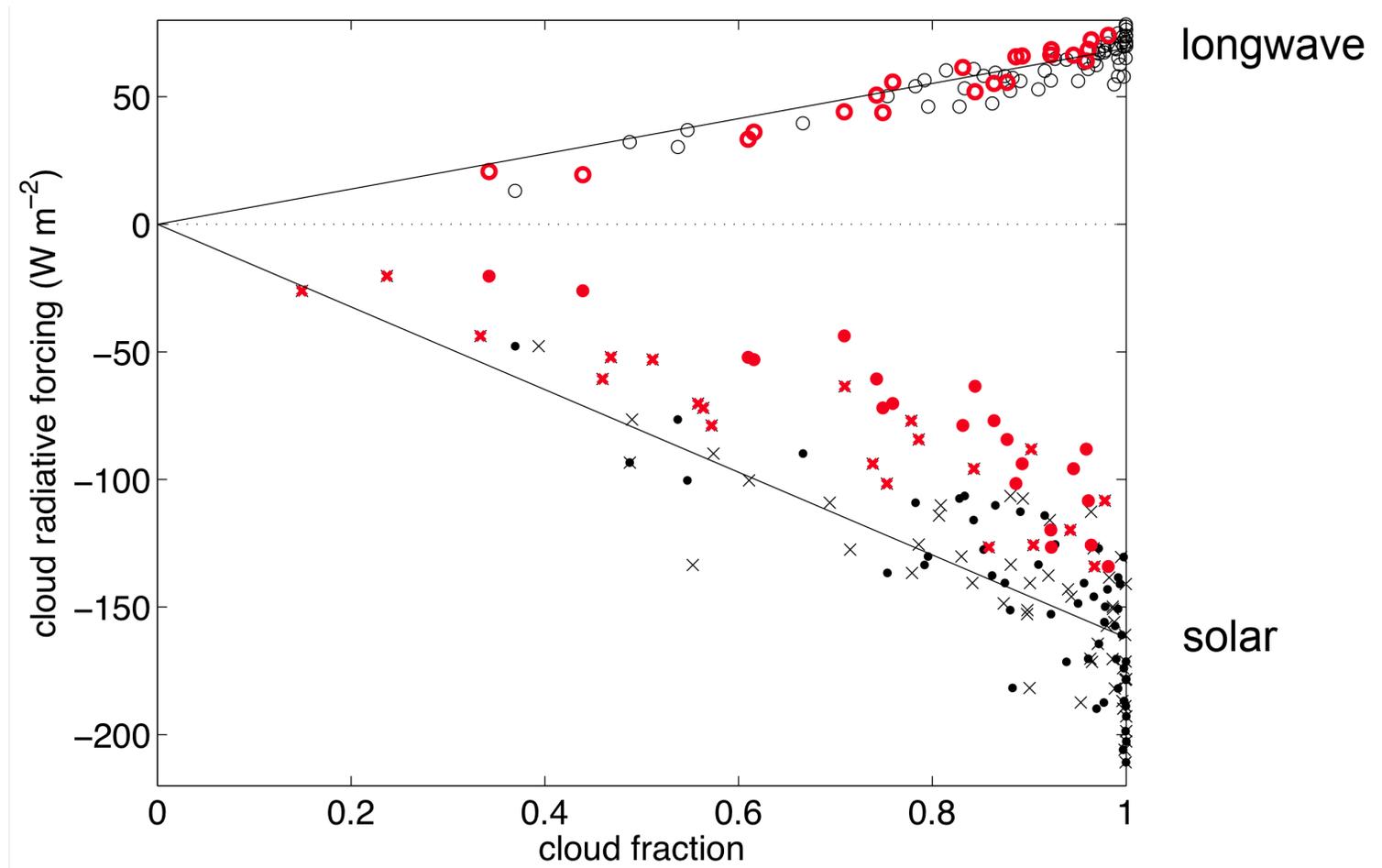


tropical

warm moist
column,

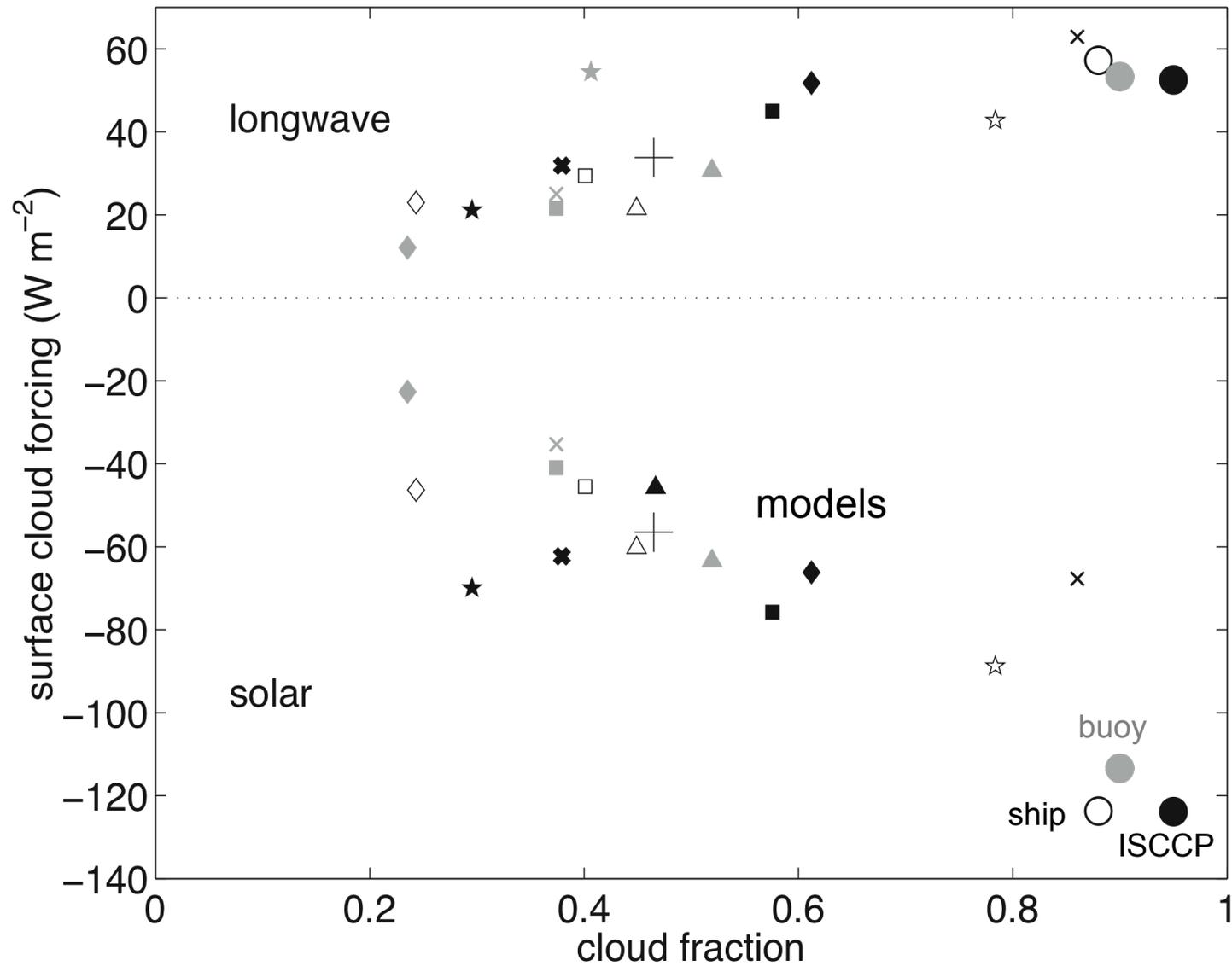
high sun →

Daily average cloud radiative forcing

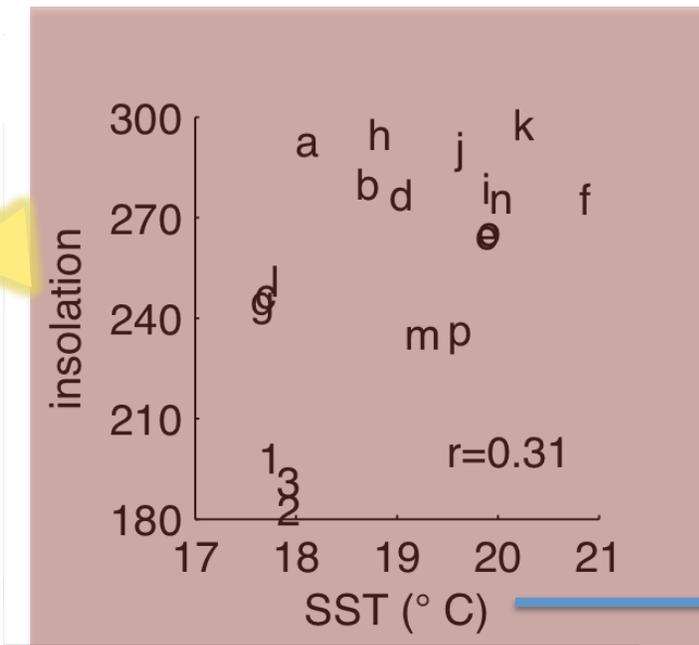
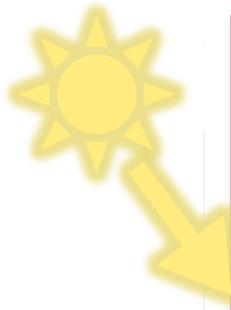


xx: solar-weighted cloud fraction
red: $2(R-R_0) > |S-S_0| + 10 \text{ W m}^{-2}$

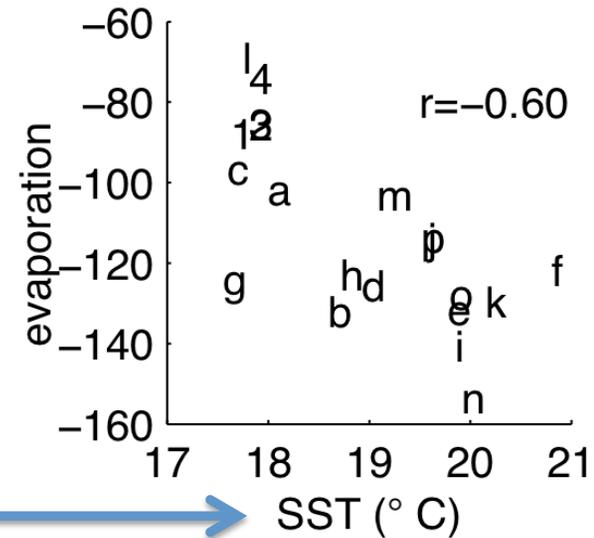
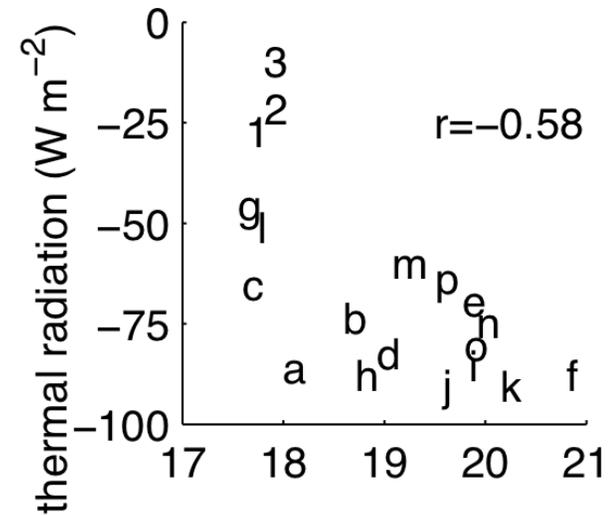
model surface cloud forcing



Does insolation explain SST?



no



surface heat budget

0 =

$-\partial/\partial t$ SST

turbulent evaporation and sensible flux

net radiation

ocean residual =

Ekman transport

geostrophic transport

upwelling

eddy flux divergence

} damp SST

the end

