ROSENSTIEL
SCHOOL of MARINE &
ATMOSPHERIC SCIENCE

An aircraft-based investigation of the first aerosol indirect effect during VOCALS-REX

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 1B "The small effective radii measured from space over SEP are primarily controlled by anthropogenic rather than natural, aerosol production, and entrainment of polluted air from the lower freetroposphere is an important source of cloud condensation nuclei"

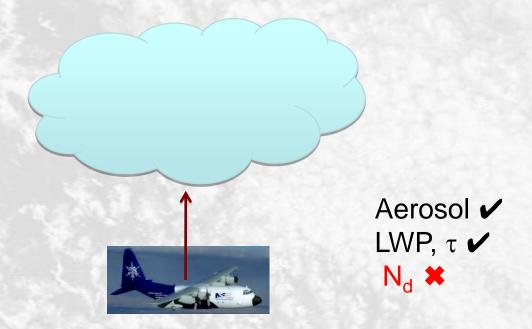
- ↑Aerosols → ↑# of droplets → ↑albedo
 - One of the largest source of uncertainty
 - Quantification of the aerosol indirect effect

Cloud-aerosol interactions

- Requirements:
 - Measurements of microphysics (r_e and/or N_d)
 - Measurements of aerosols (CCN, accumulation mode)
 - Isolate the aerosol effect from effects associated with:
 - Precipitation
 - Cloud dynamics (LWP)

Studying cloud-aerosol interactions

- In-situ measurements: C-130 aircraft
 - Subcloud observations: larger dataset



N_d can be parameterized as a function f(LWP,τ)

Retrievals

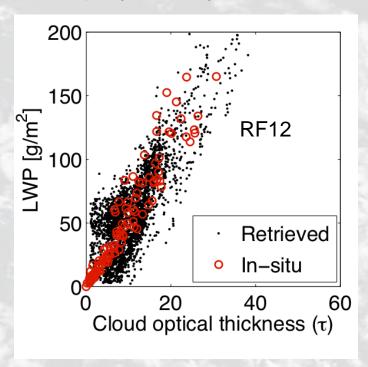
 Iterative algorithm based on radiative transfer simulations.

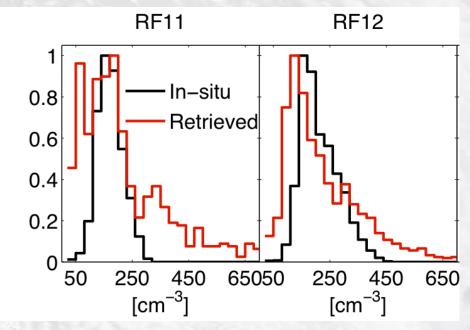
LWP: GVR microwave radiometer

τ: pyranometer

τ vs LWP and N_d

 Good correlation between τ and LWP: retrievals are physically consistent





$$N_d = A \times \frac{G}{k} \frac{t^3}{LWP^{5/2}}$$

Retrieved N_d tends to present wider distributions.

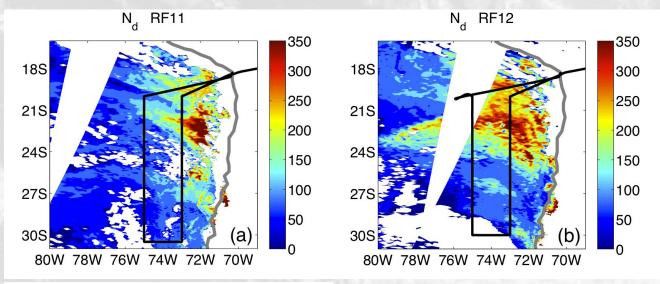
Quantifying cloud aerosol interactions

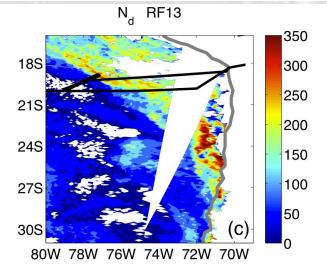
• 1st aerosol indirect effect: Increase in aerosol concentration (N_a) leads to an increase in cloud droplet number concentration (N_d) $N_a \ \ \, \square \ \, N_d \ \ \, \square \ \, t^3$

$$ACI_{N} = \frac{\text{flog}(N_{d})}{\text{flog}(N_{a})} \gg 3ACI_{t} = 3 \times \frac{\text{flog}(t)}{\text{flog}(N_{a})}$$

- Theory imposes constraint in ACI_N:
 - 0 < ACI_N < 1
 - It is necessary to control for the cloud dynamics (LWP) because:
 - Changes in N_d are controlled by both: LWP and N_a
 - We are interested in isolating the aerosol effect

Regional perspective: MODIS N_d



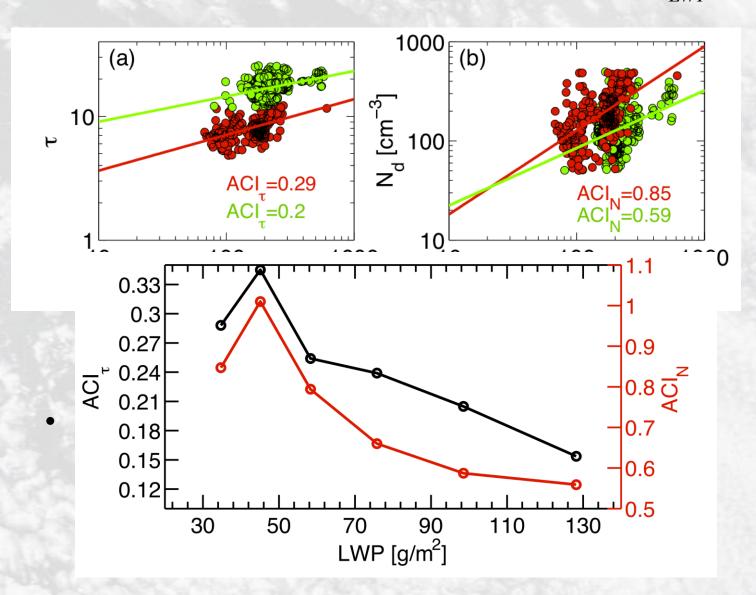


N_a-N_d correlation

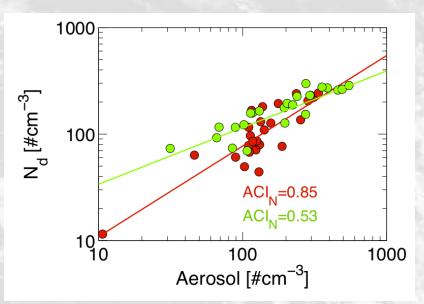
- Strategy:
 - Stratify observation based on LWP:
 - LWP_{i+1}=1.3.LWP_i
- Statistically significant correlations

LWP bins [g/m ²]	N° of samples	r
30-39.7	207	0.55
39.5-50.7	315	0.58
50.7-65.9	355	0.5
65.9-85.7	362	0.52
85.7-111.4	312	0.4
111.4-145	133	0.27

$$\frac{\mathsf{ACI}_{\mathsf{N}}}{\mathsf{Nog}\,N_a}\bigg|_{\mathit{LWP}}$$



ACI_N from cloud probes

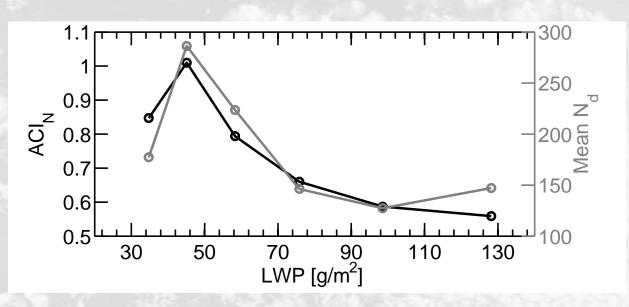


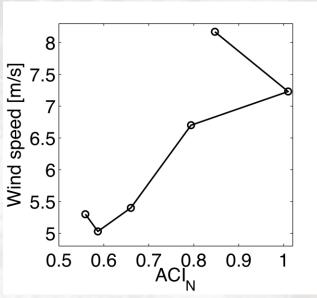
- ACI_N estimates for aricraft ascent/descent only:
 - N_a right before/after ascents/descents
 - Mean N_d of the cloudy column.

LWP<40 LWP>50

- A dependence of ACI on LWP is confirmed
- High values consistent with some other aircraft studies (e.g Twohy et al., 2005) and satellite retrievals (Painemal and Zuidema, 2010).
- Higher values than ARM site-based investigations (ACI< 0.5).
 - Temporal resolution?
 - Chile-Peru Sc higher sensitivity?

Controlling factors





- ACI decreases with N_d
- ACI increases with wind speed
- Implications for the overall cloud radiative response?

Radiative implications

• $Q = 800 \text{ W/m}^2$

LWP	% Na	% Nd	% A	$\Delta Q [W/m^2]$
30-39.7	10 %	8.5	2.2	-17.6
39.5-50.7	10 %	10	2.45	-19.6
50.7-65.9	10 %	7.9	1.66	-13.28
65.9-85.7	10 %	6.6	1.14	- 11.2
85.7-11.4	10 %	5.8	0.93	- 7.44
11.4-145	10 %	5.6	0.68	-5.44

• e.g. IPCC estimates < |-3 W/m²|

Conclusions

- Combination of remote sensing allow estimate the cloud aerosol first indirect effect.
- Quantitative assessment of the first indirect effect
 - High temporal resolution(1s)
 - Relatively high N_a-N_d correlations
- Results consistent with cloud probes-based estimates.
 - Chile-Peru Sc present high sensitivities to aerosol variability.
- High radiative forcing, magnitudes 5-10 times larger than GCM and satellite estimates

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