

Observations of the Diurnal Cycle of Marine Stratocumulus Clouds and Precipitation

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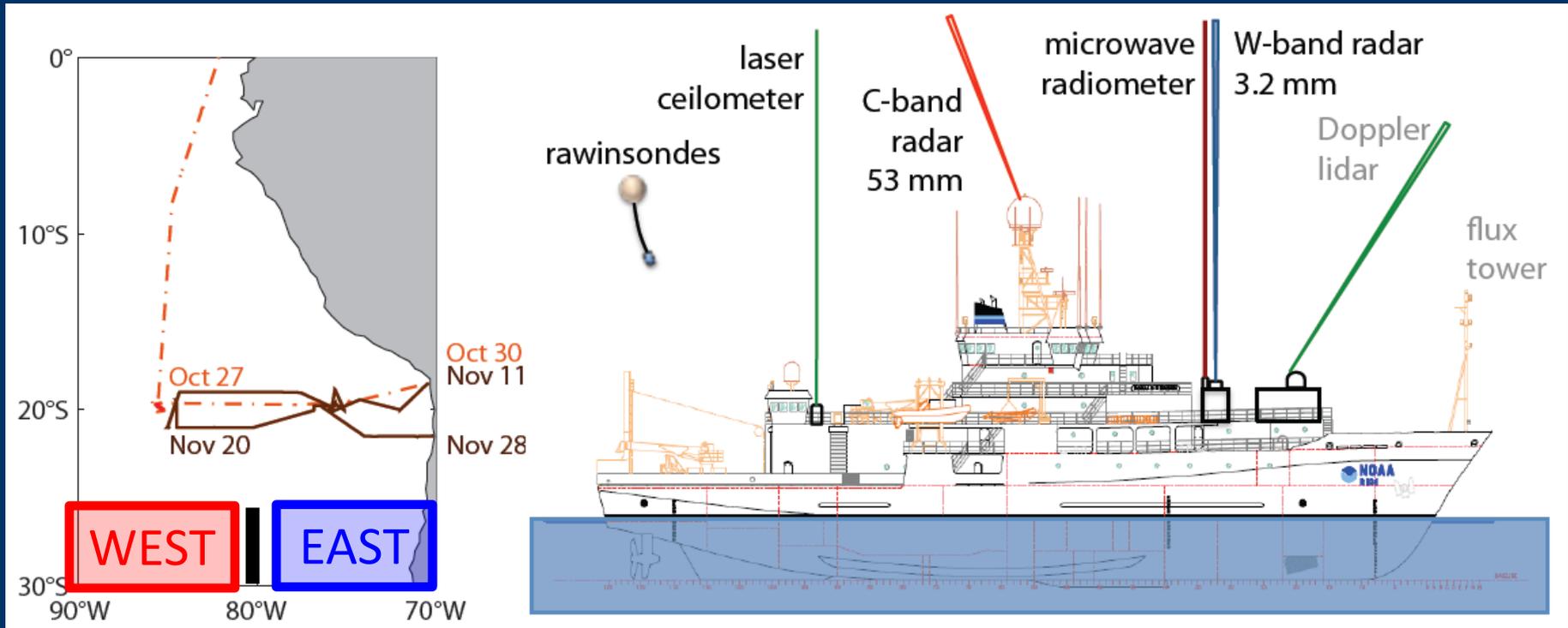
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Goals

- 1) Document the diurnal variation of important features of the cloud deck to aid in validation of model output.
- 2) Examine to what degree drizzle varies in time and space.
- 3) Begin to understand the controlling factors that cause drizzle to start and stop.

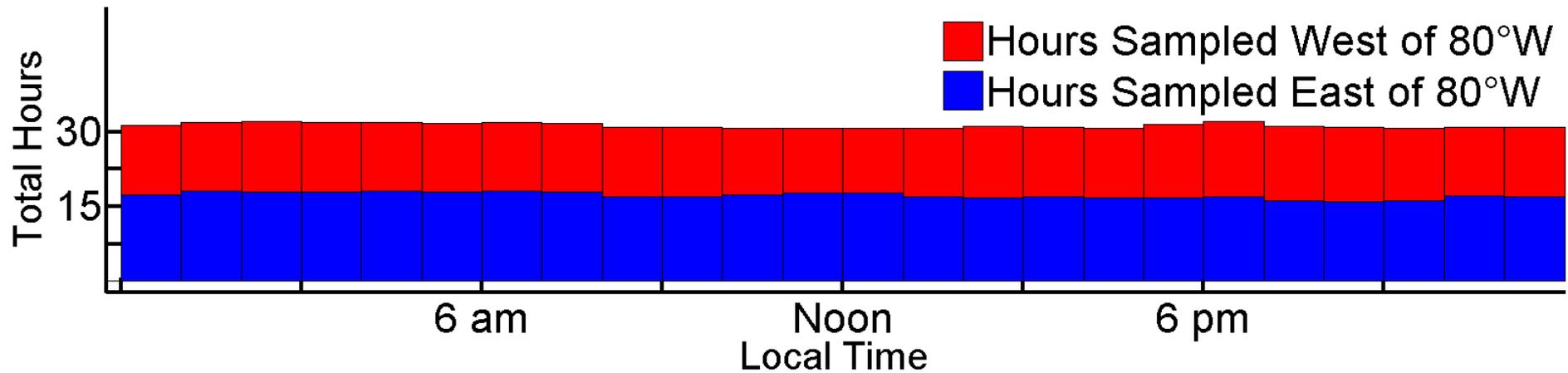
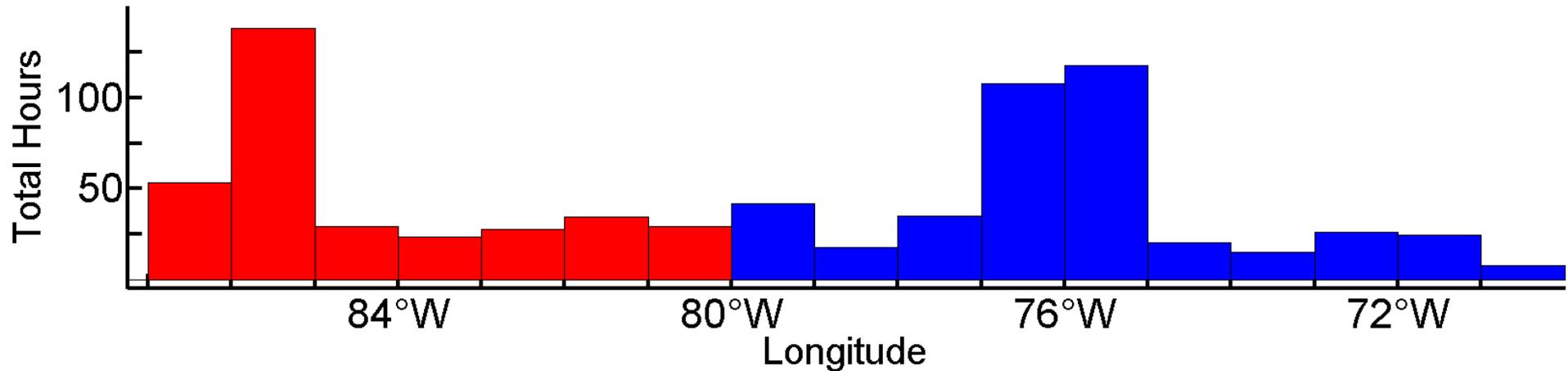
Data from the R.H. Brown



Data were divided into two regions, east and west of 80°W, to account for longitudinal changes to boundary layer characteristics.

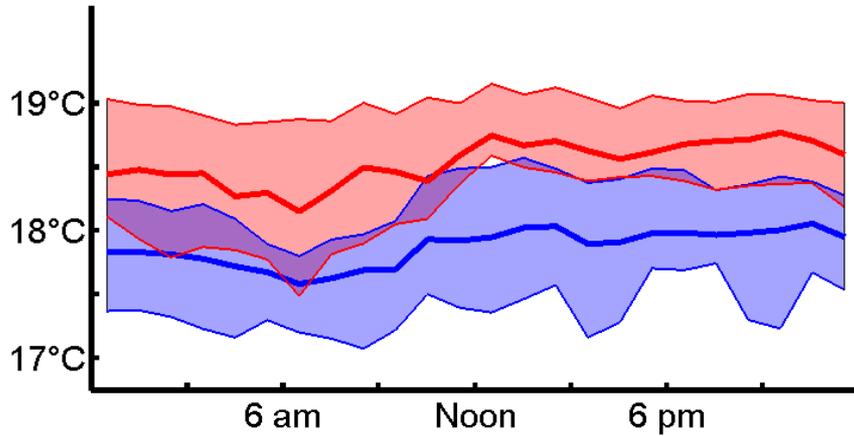
C-Band Radar Sampling

Hours of Radar Data in Space and Time

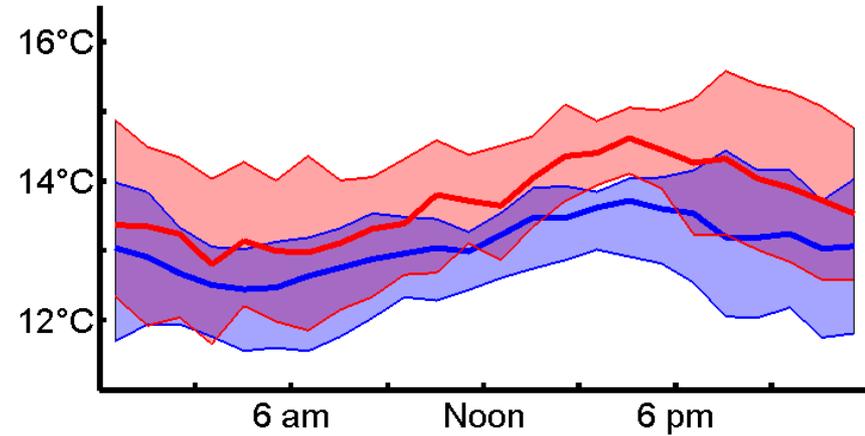


Surface Observations

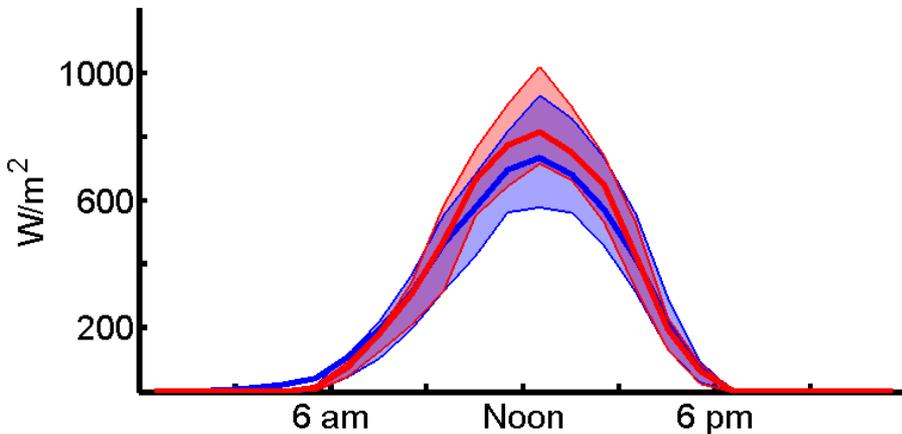
Temperature



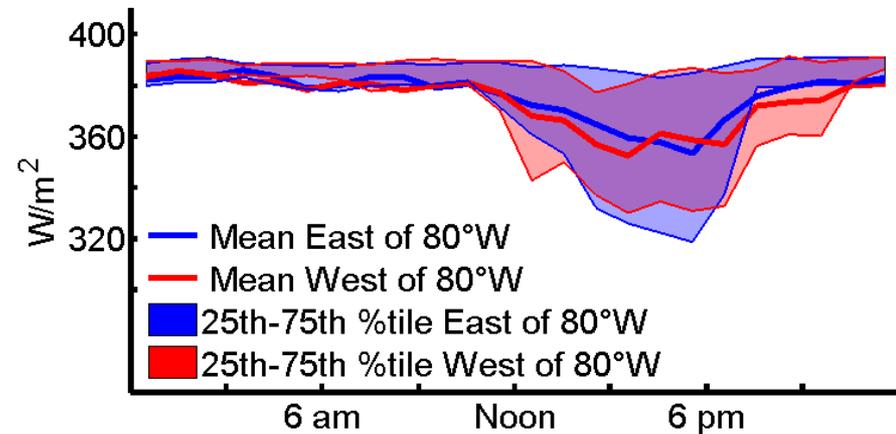
Dew Point Temperature



Shortwave Radiation

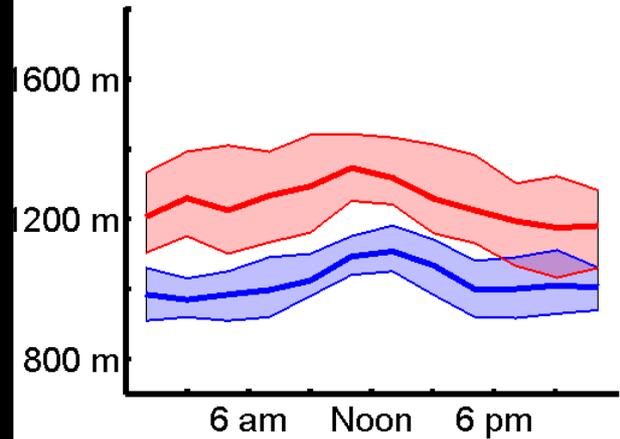


Longwave Radiation

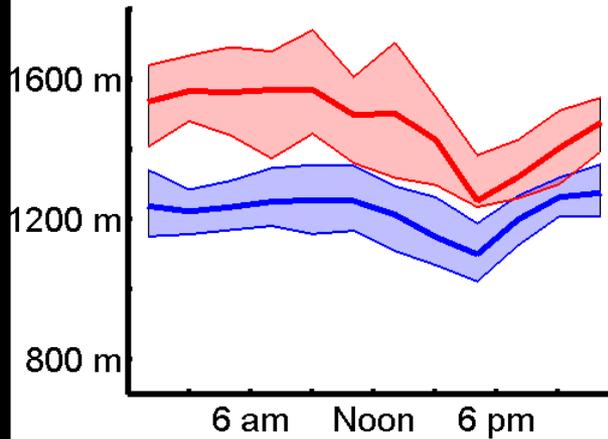


Cloud Observations

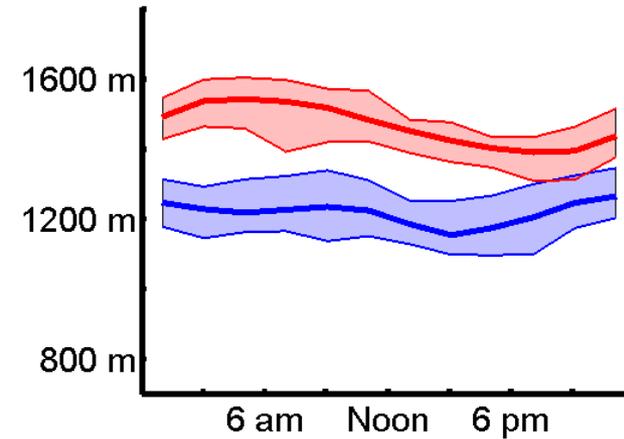
Ceilometer Cloud Base



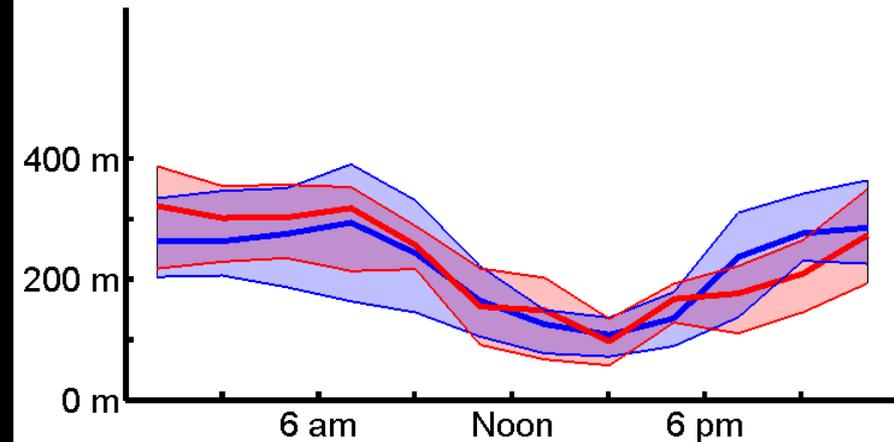
Radar Cloud Top



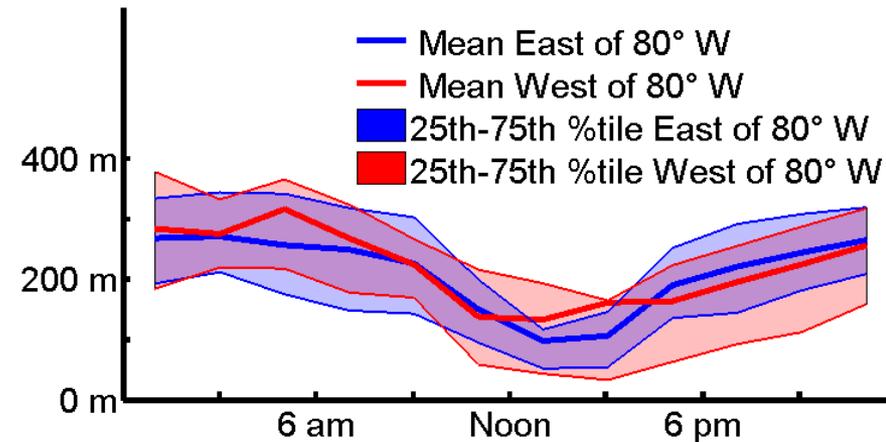
Sounding Cloud Top



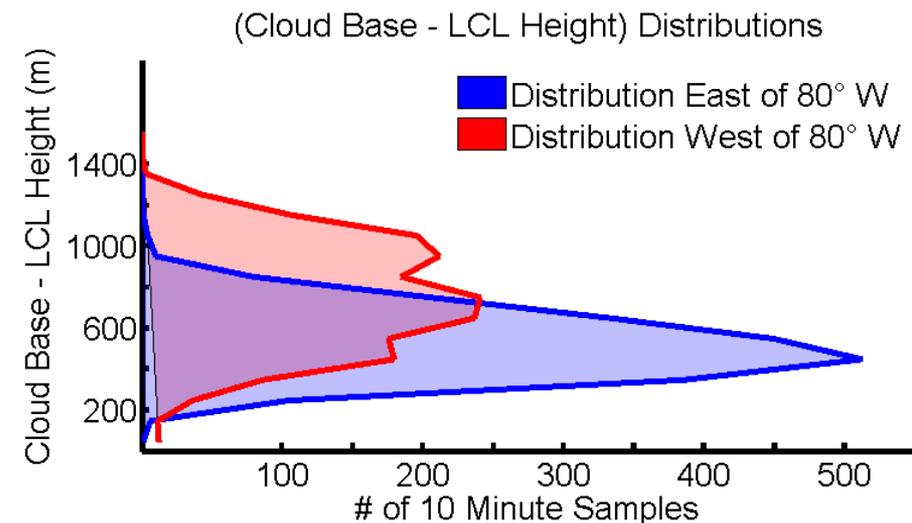
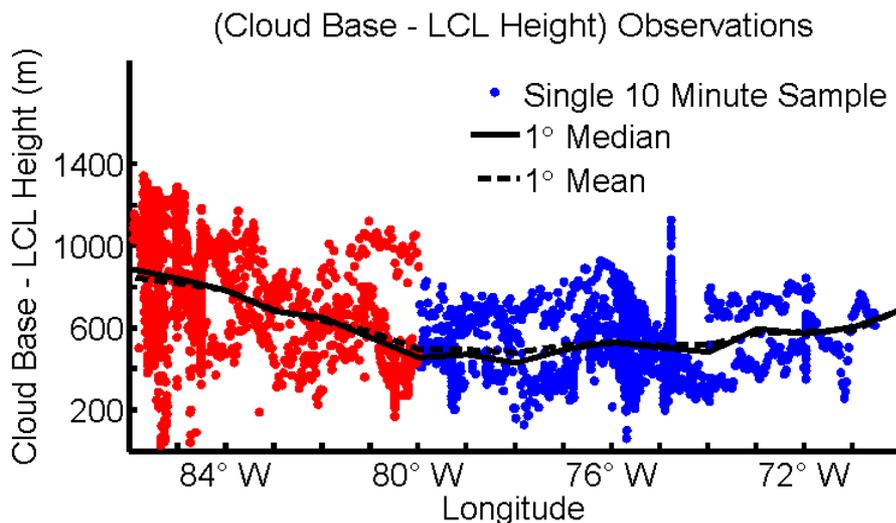
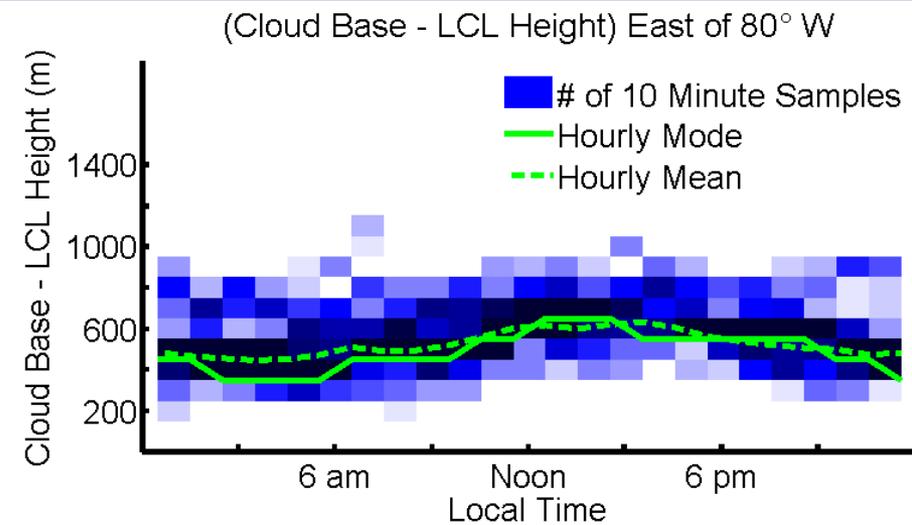
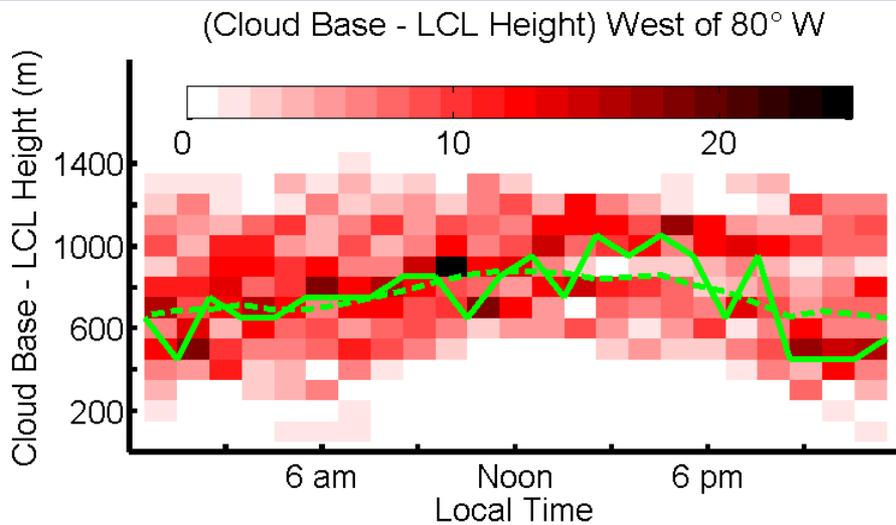
[Radar Top - Ceilometer Base] Cloud Depth



[Sounding Top - Ceilometer Base] Cloud Depth

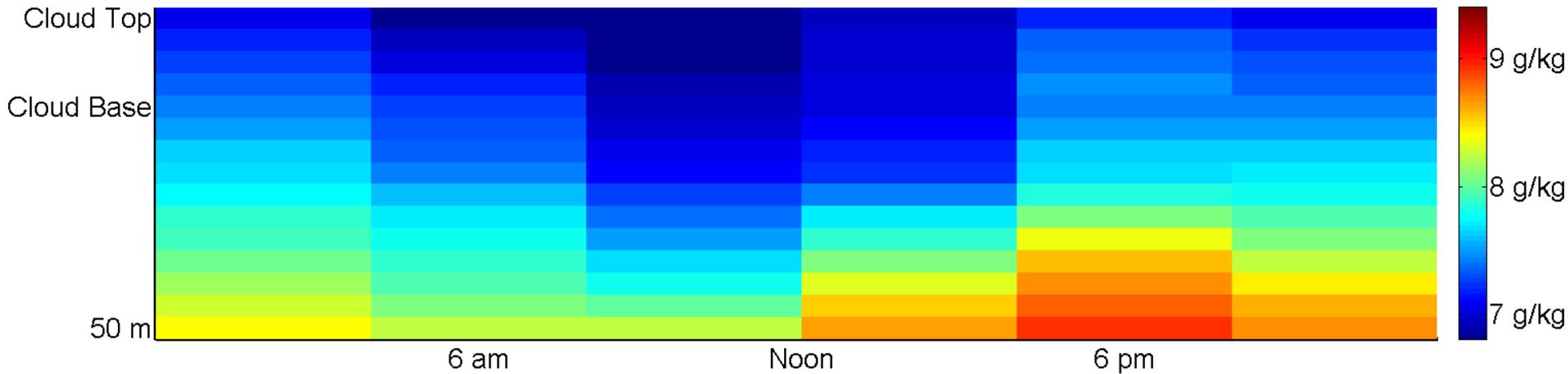


Boundary Layer Coupling

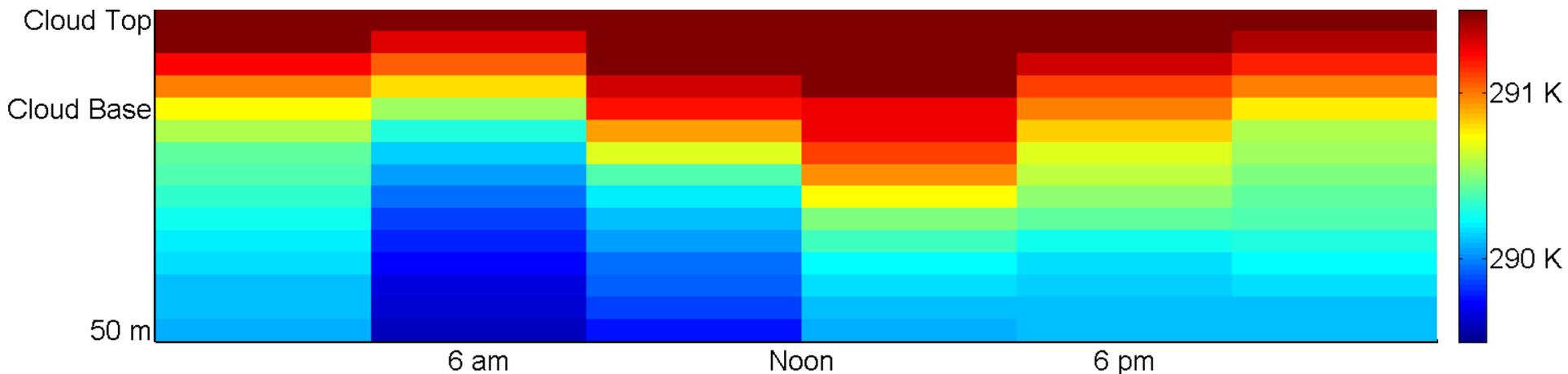


Boundary Layer Mixing

Water Vapor Mixing Ratio

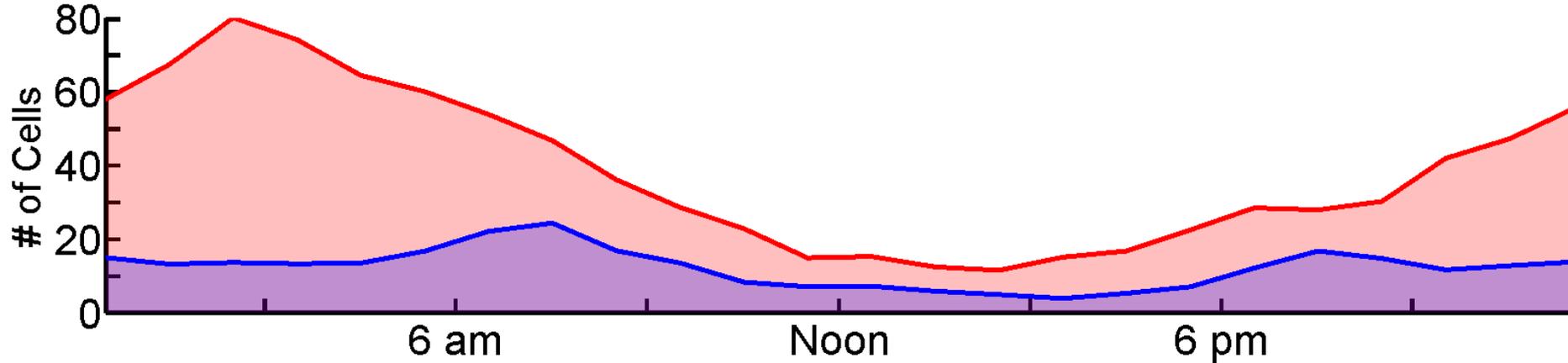


Potential Temperature

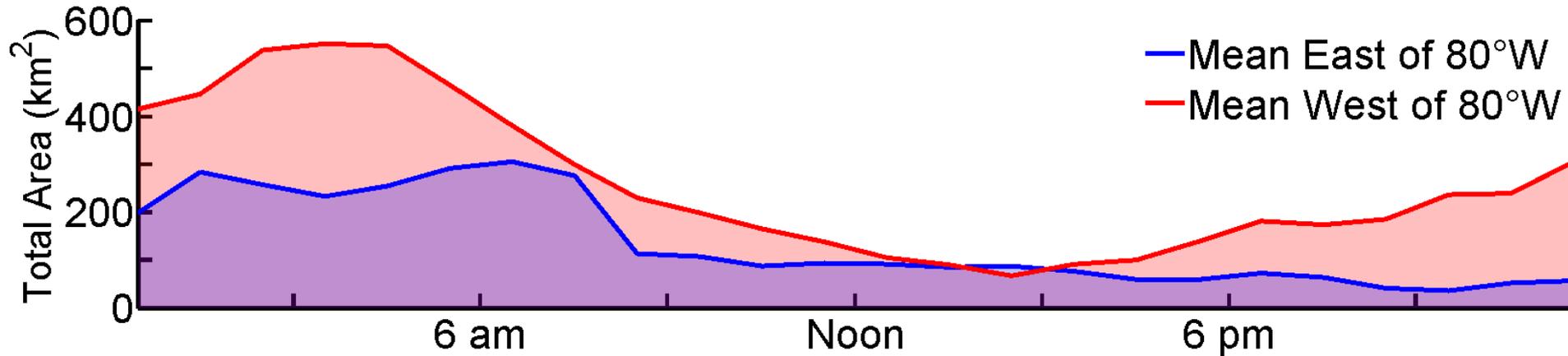


C-Band Radar Observed Drizzle

Drizzle Cells Per Scan

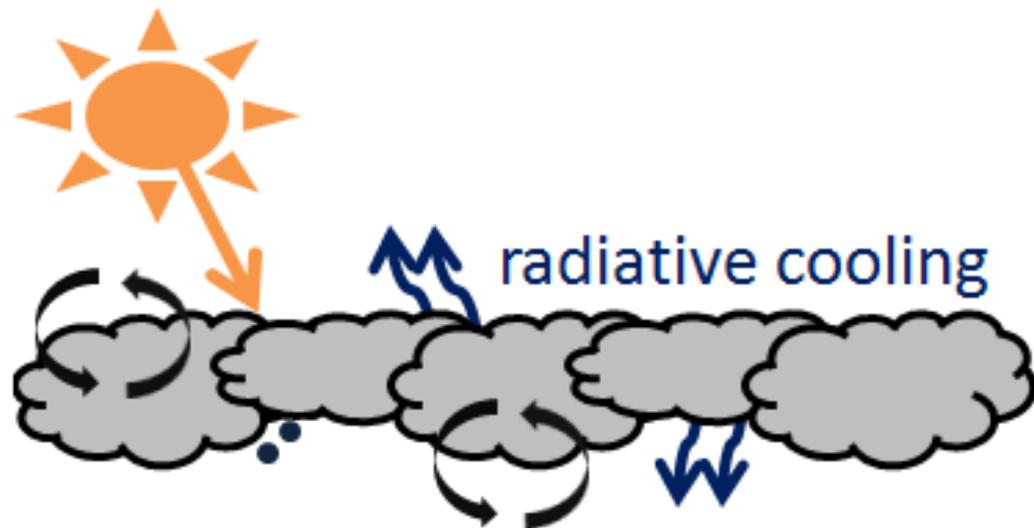


Drizzle Area Per Scan

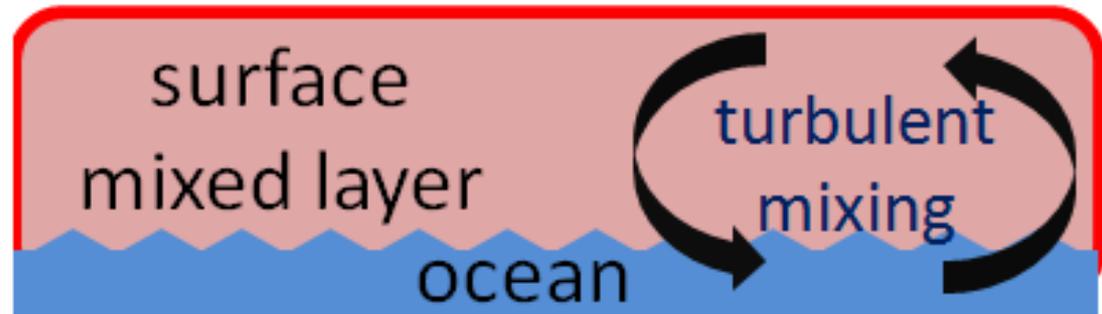


Noon

$h_c =$ cloud depth



$h_{ML} =$ mixed layer depth



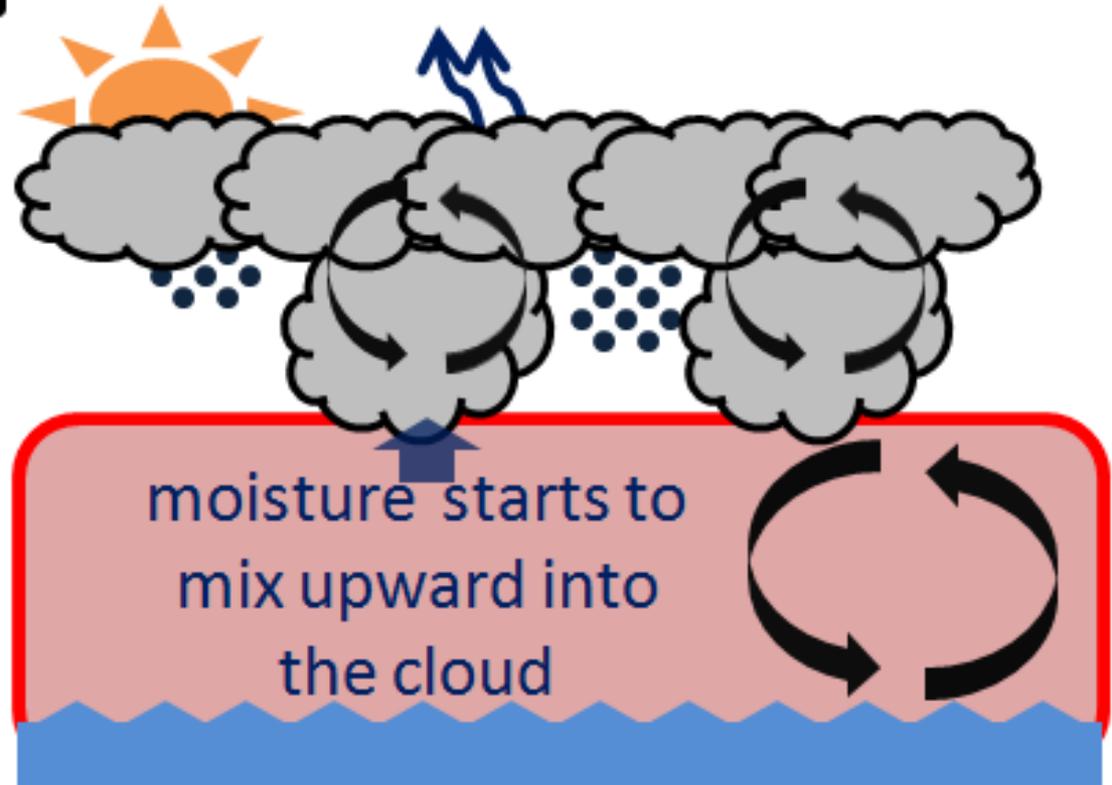
The surface mixed layer is decoupled from the cloud deck.

Sunset

$$\frac{d(h_c)}{dt} > 0$$

$$\frac{d(\text{precip})}{dt} > 0$$

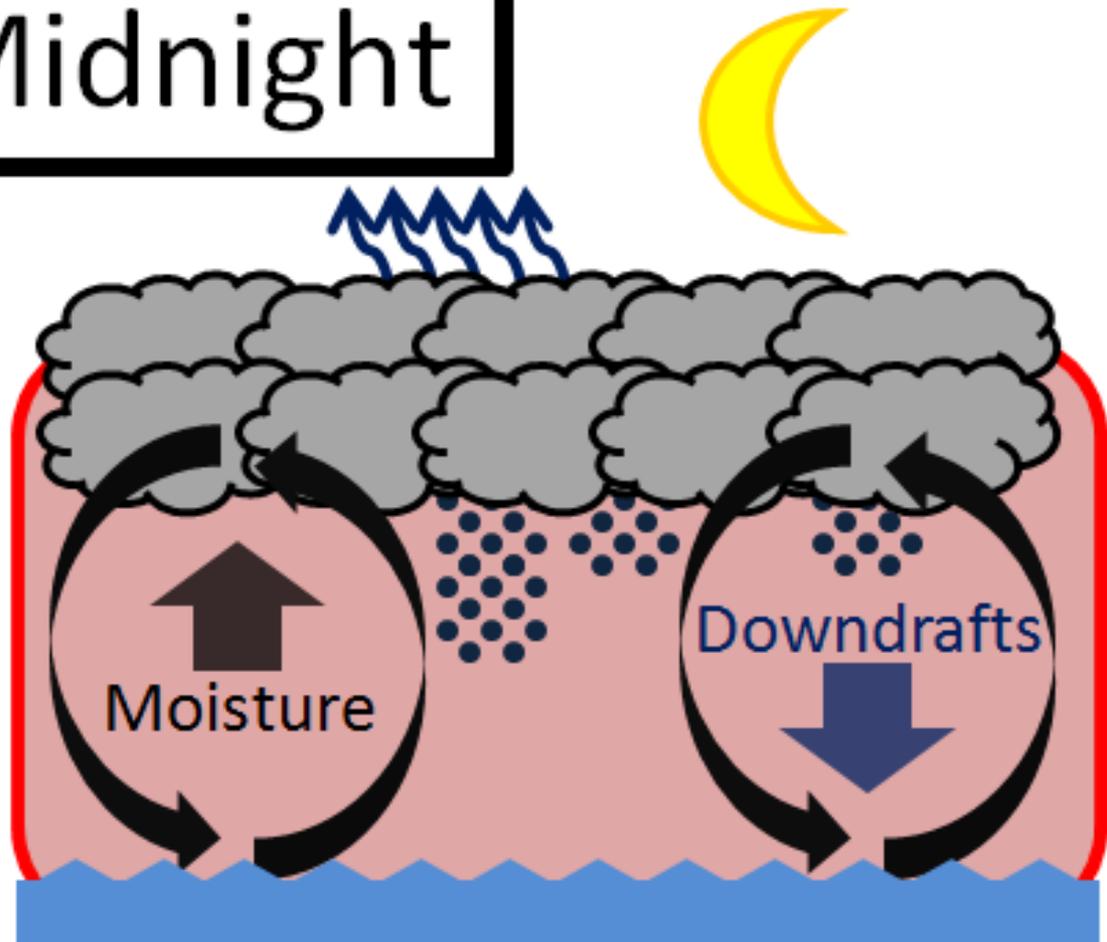
$$\frac{d(h_{ML})}{dt} > 0$$



The mixed layer is deeper and may be coupled with the cloud deck via cumulus clouds.

9 pm -Midnight

$$\frac{d(h_c)}{dt} > 0$$
$$\frac{d(\text{precip})}{dt} > 0$$
$$\frac{d(h_{ML})}{dt} > 0$$



The boundary layer is coupled through cloud top radiative cooling and cellular convection.

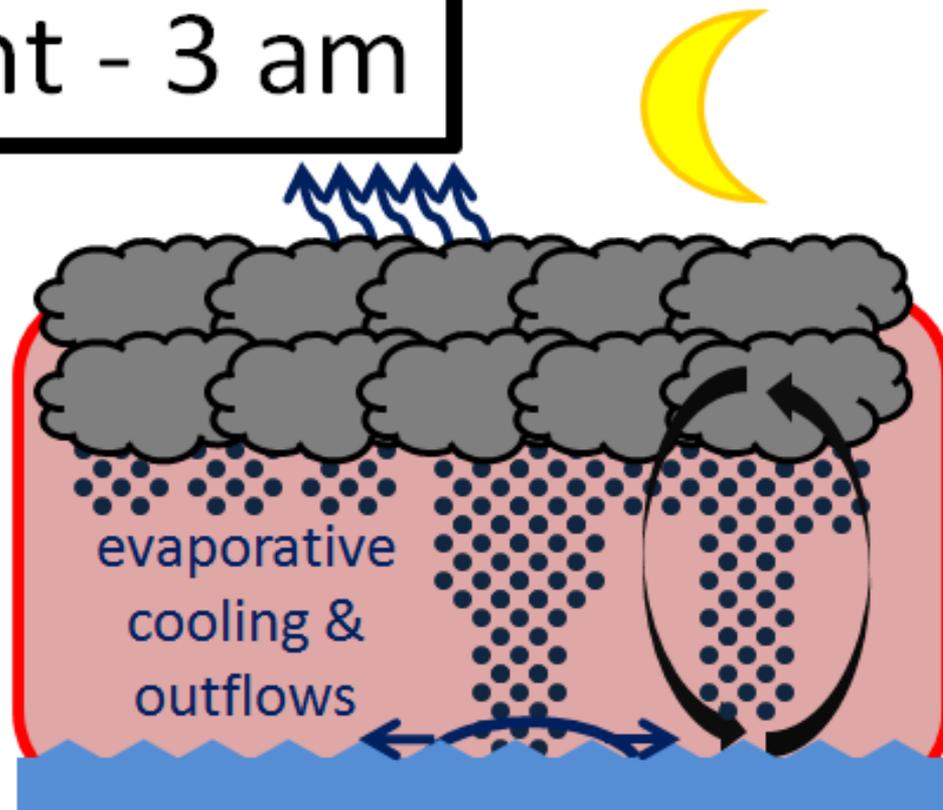
Heavy Drizzle (West)

Midnight - 3 am

$$\frac{d(h_c)}{dt} = 0$$

precip = max

$$\frac{d(h_{ML})}{dt} = 0$$

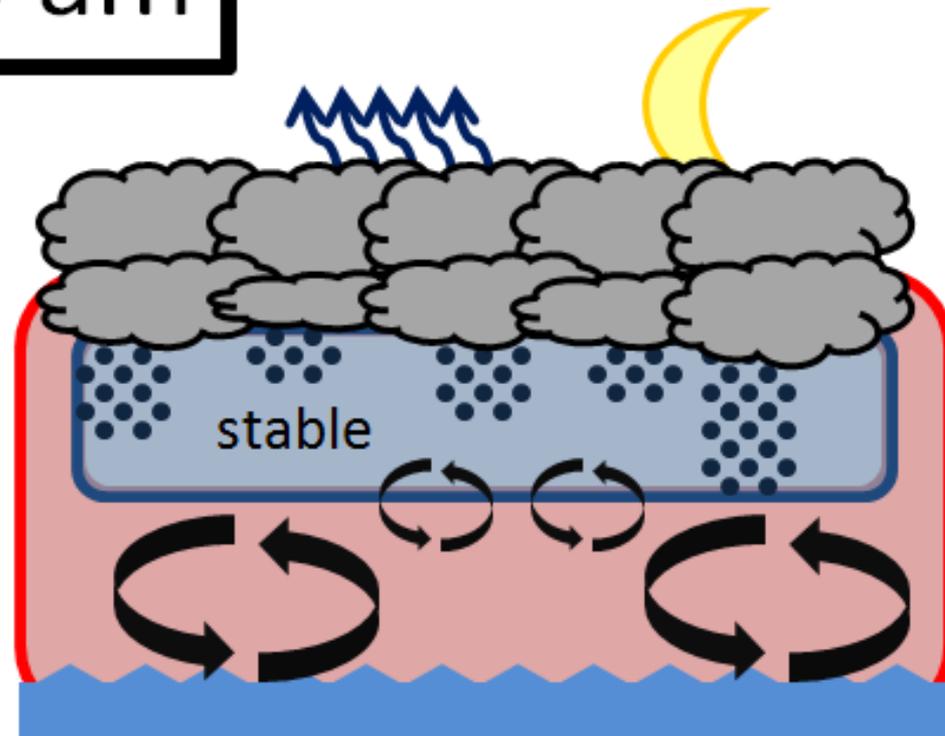


Diurnal maximum in precipitation and degree of coupling. Some drizzle evaporates.

Heavy Drizzle (West)

3 am - 6 am

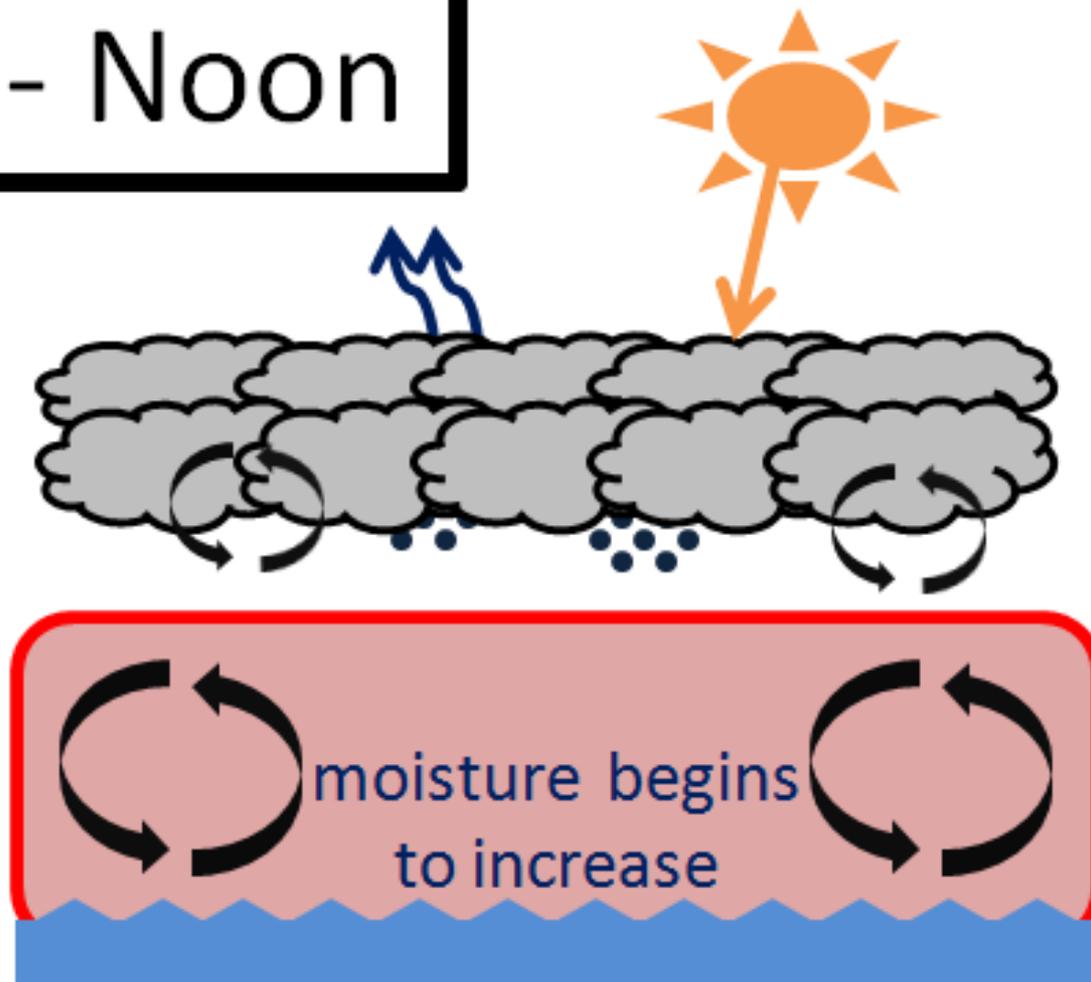
$$\frac{d(h_C)}{dt} \sim 0$$
$$\frac{d(\text{precip})}{dt} < 0$$
$$\frac{d(h_{ML})}{dt} \sim 0$$



An evaporatively cooled stable layer develops and moisture flux into cloud is reduced.

Sunrise - Noon

$$\frac{d(h_c)}{dt} \ll 0$$
$$\frac{d(\text{precip})}{dt} \ll 0$$
$$\frac{d(h_{ML})}{dt} < 0$$



Solar heating offsets radiative cooling and the boundary layer becomes less coupled.

Conclusions

- 1) The diurnal cycle is clearly evident in every variable examined. Diurnal fluctuations in cloud depth, precipitation, and boundary layer mixing should be key features in simulations of STBLs.
- 2) Overnight the cloud depth is similar over all longitudes. The cloud deck thins rapidly after sunrise.
- 3) Heavy drizzle that occurs in the west is self-limiting, maximizing near 3 am. Lighter drizzle in the east only diminishes after the sun comes up.

Noon

$h_c = \text{cloud depth}$
 $h_{ML} = \text{mixed layer depth}$

radiative cooling

surface mixed layer

turbulent mixing

ocean

The surface mixed layer is decoupled from the cloud deck.

Sunset

$\frac{d(h_c)}{dt} > 0$
 $\frac{d(\text{precip})}{dt} > 0$
 $\frac{d(h_{ML})}{dt} > 0$

moisture starts to mix upward into the cloud

The mixed layer is deeper and may be coupled with the cloud deck via cumulus clouds.

9 pm - Midnight

$\frac{d(h_c)}{dt} > 0$
 $\frac{d(\text{precip})}{dt} > 0$
 $\frac{d(h_{ML})}{dt} > 0$

Moisture

Downdrafts

The boundary layer is coupled through cloud top radiative cooling and cellular convection.

Heavy Drizzle (West)

Midnight - 3 am

$\frac{d(h_c)}{dt} = 0$
 $\text{precip} = \text{max}$
 $\frac{d(h_{ML})}{dt} = 0$

evaporative cooling & outflows

Diurnal maximum in precipitation and degree of coupling. Some drizzle evaporates.

Heavy Drizzle (West)

3 am - 6 am

$\frac{d(h_c)}{dt} \sim 0$
 $\frac{d(\text{precip})}{dt} < 0$
 $\frac{d(h_{ML})}{dt} \sim 0$

stable

An evaporatively cooled stable layer develops and moisture flux into cloud is reduced.

Sunrise - Noon

$\frac{d(h_c)}{dt} \ll 0$
 $\frac{d(\text{precip})}{dt} \ll 0$
 $\frac{d(h_{ML})}{dt} < 0$

moisture begins to increase

Solar heating offsets radiative cooling and the boundary layer becomes less coupled.

Drizzle Intensity Distributions

Distribution of Area Average Rain Rates

