Direct observations of reactive trace gases over the eastern Pacific Ocean





CU Ship MAX-DOAS target gasses: <u>CHOCHO, IO</u>, HCHO, NO₂ (OIO, I₂, BrO, SO₂)







Inside: spectrometers/clinometers

Outside: telescope

Cruise tracks (Oct08-Dec08)



Objective: is there glyoxal over the ocean? Where could it be coming from?



C: OMI, VC_{CHOCHO}, Annual mean 2006



D: MODIS Chlorophyll-a, 2005 (ocean color)





Glyoxal (CHOCHO)

Global source α -dicarbonyls:

- Glyoxal: 40%
- Methylglyoxal: 55% → > 185 Tg/yr
- Biacetyl: 5% (uncertain by factor 2)

Glyoxal source:

- 30% biogenic VOC (isoprene)
- 20% anthropogenic VOC
- 50% unaccounted land sources
- Secondary sources dominate over primary sources

108 Tg/yr

- biomass burning? (potentially 5-10 Tg/yr)
- ocean sources?
- energy sector, industry?

Atmospheric lifetime: <1.2h

- Photolysis (60%)
- OH-radical reaction (40%)
- aerosol loss timescales are uncertain (could determine 75 – 95% of the CHOCHO lifetime
- cloud processing timescales (few hours)
- Source for H₂, CO, HCHO and HO₂-radicals

Global SOA source: 5-11 Tg/yr

- Potentially comparable to aromatics and terpenes
- 85% deemed to form in clouds (?)
- Lower limit estimate (sources underestimated !)



SCIA GOME2 2008 annual averages



Vertical distribution: CHOCHO in MBL!



Satellite can not measure the vertical distribution (FT profile) Higher VC by SMAX-DOAS can be explained by MBL profile => 2-4 times higher VC over oceans (largest globally!!)

lodine oxide (IO)

Iodine sources:

- Highly uncertain on global scales
- Mostly constrained from studies in coastal environments: 1.5 Mg I yr⁻¹ km⁻¹
- Very few observations over open ocean

Currently suggested mechanisms:

- Production from alkylhalides
- I₂ from macroalgae (low tide)
- Cape Verde: current mechanisms are insufficient to account for observed IO
- I₂ from microalgae (?)
- Sea salt (?)

Atmospheric lifetime: sec

Atmospheric relevance:

- Destruction of trop ozone
- Oxidative capacity
- New particle formation/particle growth
- Mercury oxidation?









Photoenhanced Reaction of Ozone with Chlorophyll at the Seawater Surface (Reeser et al., J. Phys. Chem. C **2009**, 113, 2071–2077):

 $\begin{array}{l} \mathsf{CHL} + \mathsf{hv} \to \mathsf{CHL}^* \to \mathsf{CHL}^+ + \mathsf{e}^- \\ \\ \mathsf{CHL} + + \mathsf{CI}^- \to \mathsf{CHL} + \mathsf{CI} \\ \\ \\ \mathsf{CI} + \mathsf{RH} \to \mathsf{R} + \mathsf{HCI} \end{array}$

Correlations CHOCHO and IO suggest extended scheme: $\begin{array}{c} CHL^{+}+I^{-}\rightarrow CHL+I\\ I+I\rightarrow I_{2}\\ I_{2}+hv\rightarrow 2I\\ I+O_{3}\rightarrow IO+O_{2}\\ OH, CI, Br+RH\rightarrow \rightarrow CHOCHO, HCHO\\ (chemical reason to accumulate I_{2}, but not Br_{2} \text{ or } CI_{2})\end{array}$



Conclusions



- Spectral proof for elevated glyoxal and IO over the open ocean (3000km from land) !
- Glyoxal and IO are located in the marine boundary layer
- Satellites underestimate VC by factor 2-4
- Tropical ocean is a global hotspot for CHOCHO
- Field evidence compatible with a chemical mechanism for halogen and OVOC release from surface photochemistry
- Tropics are chemically very active atmosphere: destroys ozone (IO), forms/grows aerosols (CHOCHO, IO)
- VOCALS: evidence for aerosol formation/growth from sources other than SO₂ (DMS)

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