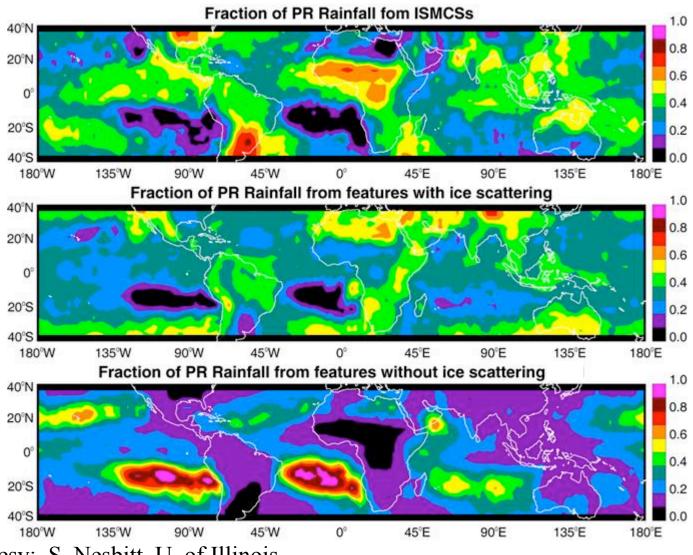
### Fraction of Total Rainfall by PF Type



Courtesy: S. Nesbitt, U. of Illinois

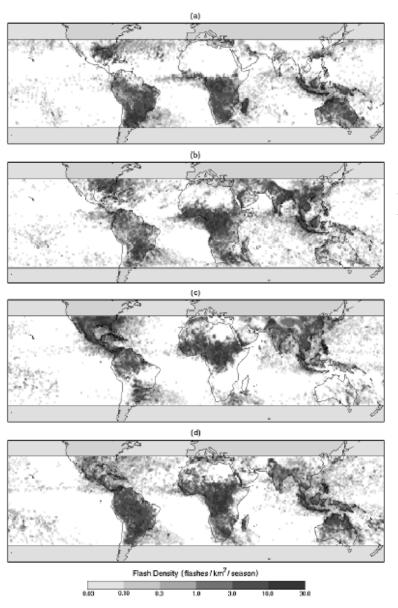


Figure 4. The 1998 seasonal distribution of lightning flashes as observed by LIS for (a) December, January, and February, (b) March, April, May, (c) June, July, and August, and (d) September, October, and November.

#### Dec-Feb

Christian et al. (2000)

Mar-May
Seasonal lightning flash densities from TRMM/LIS for 1998.

June-Aug

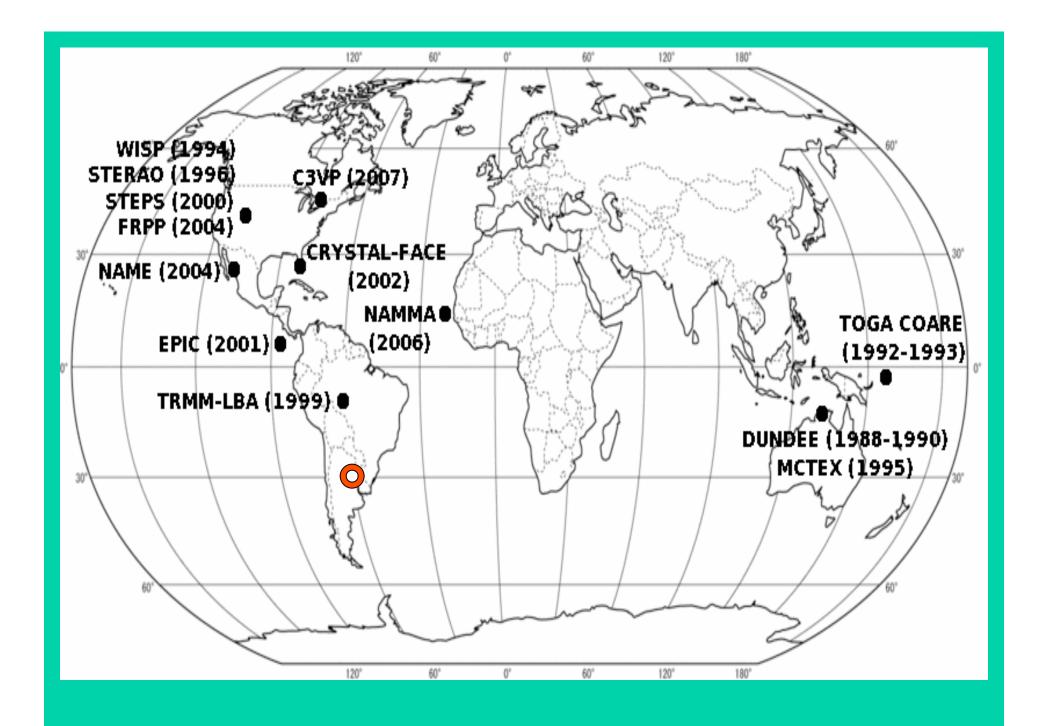
Sept-Nov

## Global Lightning Observed from the NASA-OTD 10-1000 times more lightning over land than Ocean

Global Flash Rate about 40 flashes per second 90 <10 >10 >25 >50 >100 >150 >200 >300 >400 >500 >750 >1000 17371 Orbits 808854 Areas Flashes 4345364 Flash scale 20727622 Groups 42265608 Events April 12, 1995 - December 31, 1999 NASA / MSFC (Created: 02/15/100)

# CSU Radar Meteorology Group Interests in LPB

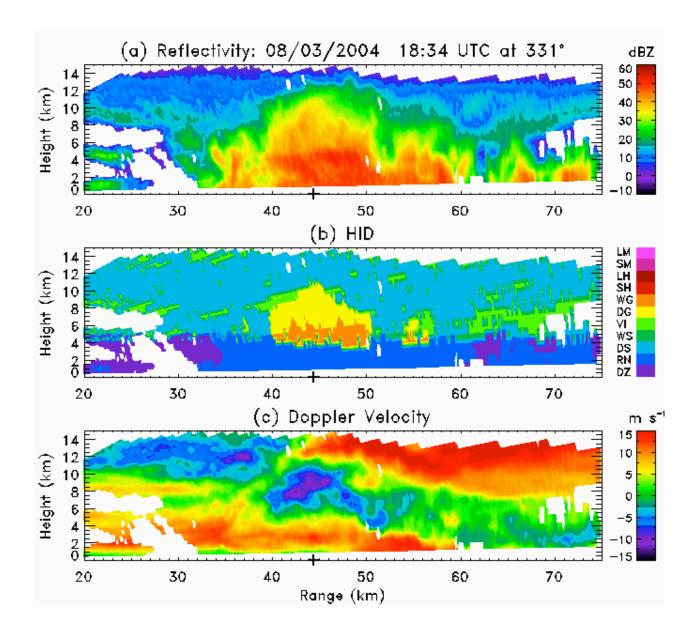
- Connect mostly to MCS related scientific objectives
- Convective scale dynamics and microphysics
- Radar-based rainfall estimation (couple to hydrology goals)
- Precipitation and aerosol influences (biomass burning, pollution)
- Electrification and lightning studies
  - Flash rates, polarity in convection and MCSs
  - Sprites

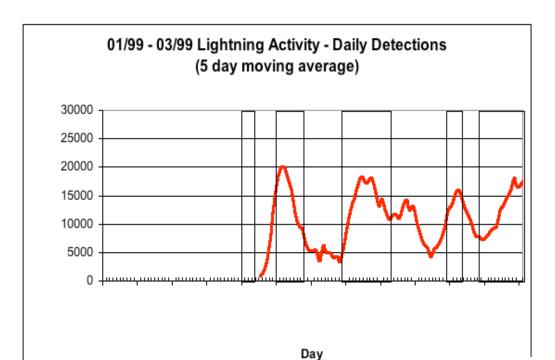


#### **TRMM-LBA**



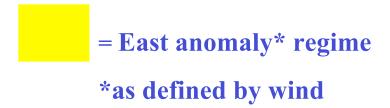




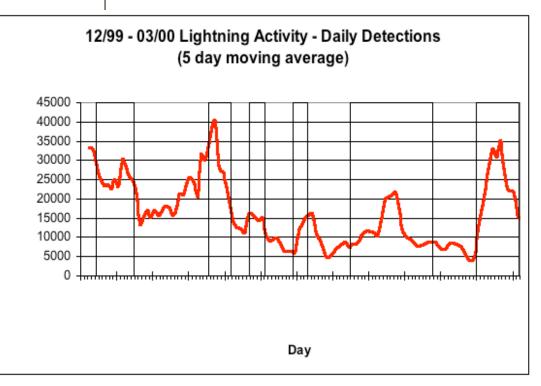


# Brazilian Lightning Detection Network (BLDN):

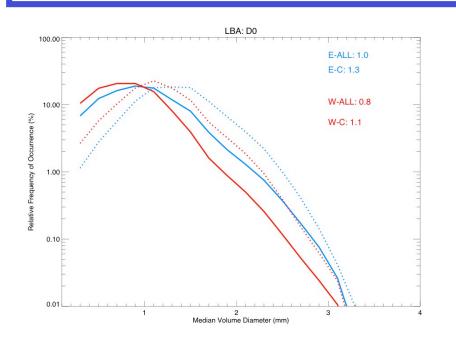
- Oscillations apparent
- East (west) anoms=more (less) lightning.

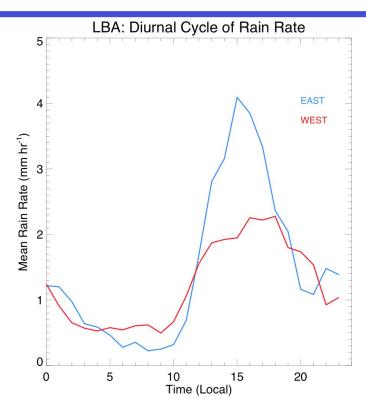


Petersen et al., 2001



Microphysical differences between east and west regimes. Larger  $D_0$  in east regime linked to stronger CAPE and stronger ice-based precipitation. Strong diurnal cycle in east regime.



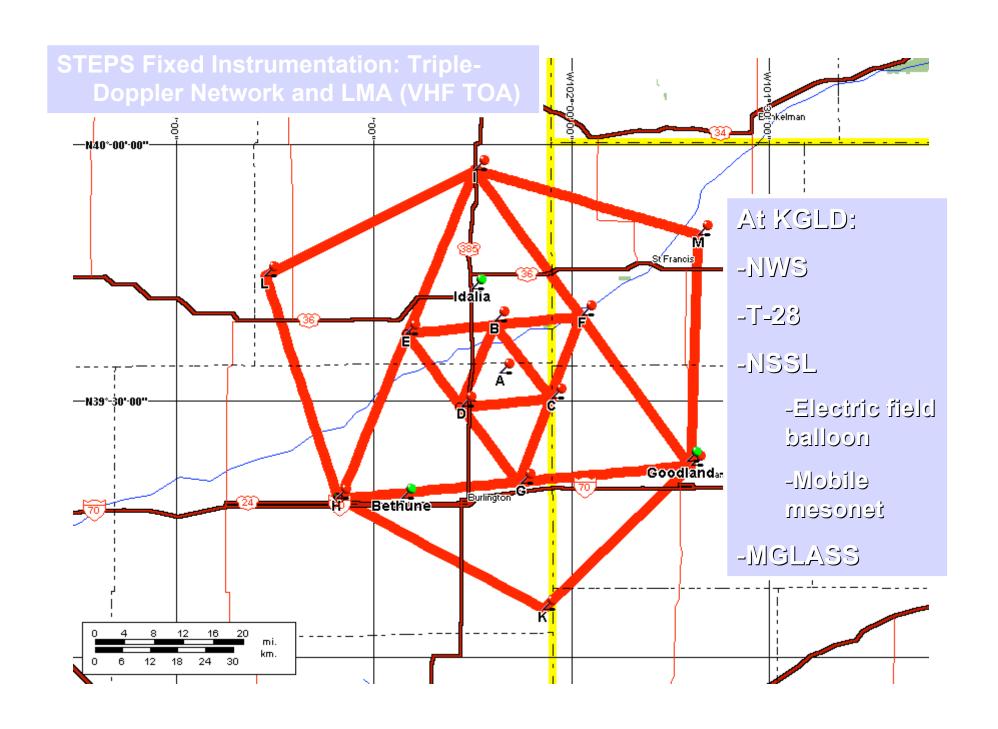


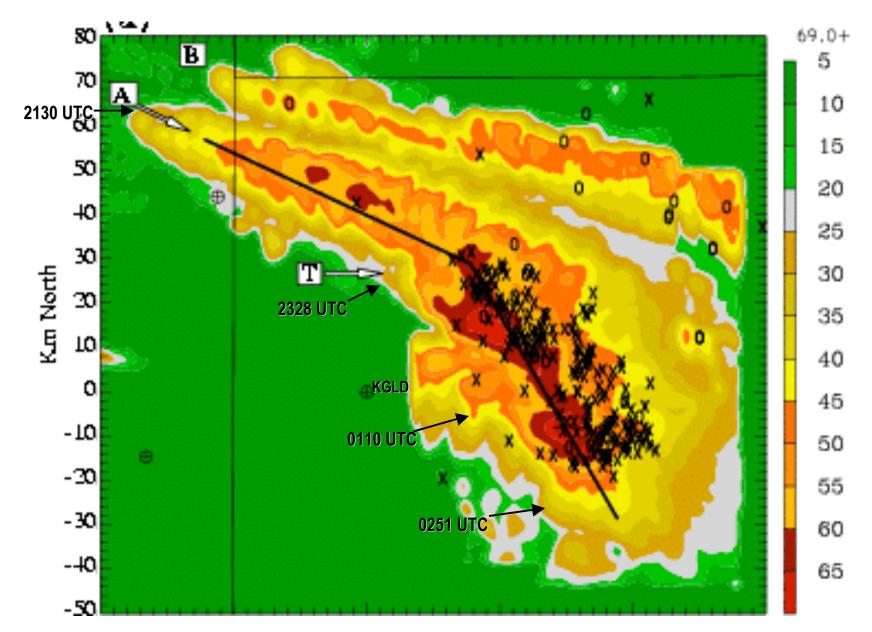
Aerosol influences may be detected through changes in  $D_0$ 





STEPS
Severe Thunderstorm Electrification and Precipitation Study
May-July 2000



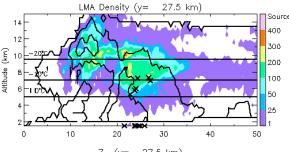


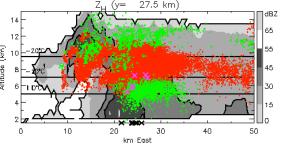
Tessendorf et al., **JAS**, 2005 Storm swath of base reflectivities (2100-0251 UTC) with NLDN lightning data overlaid.

#### 29 June Supercell

- Inverted tripole
- $\bullet$  +CGs

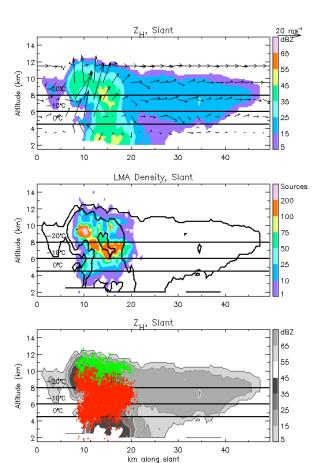
# $Z_{H}$ (y= 27.5 km) 20 mg d62 65 65 55





#### 3 June storm

- Inverted dipole
- No CGs

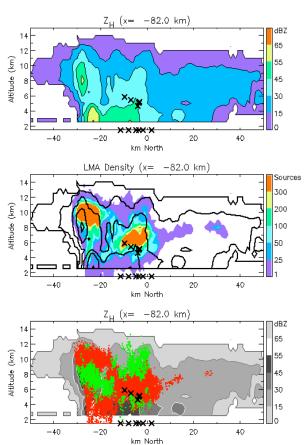


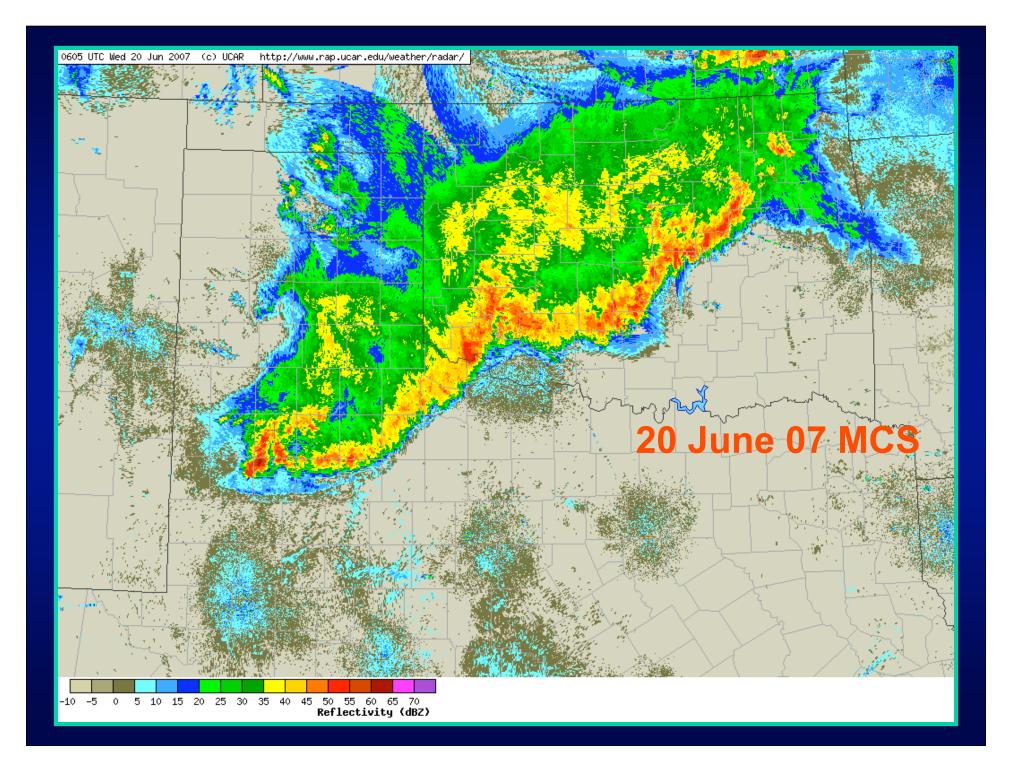
## 23 June storm Early:

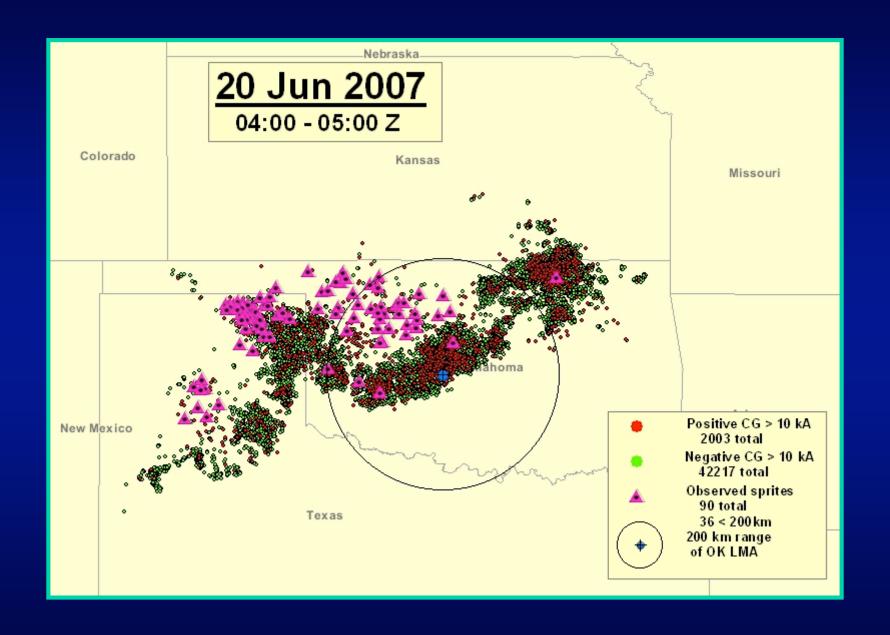
normal tripole, -CGs

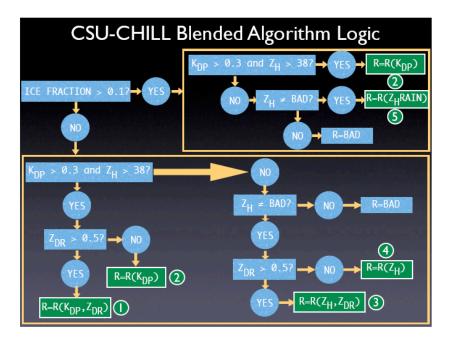
#### Later, collapse:

inverted tripole, +CGs

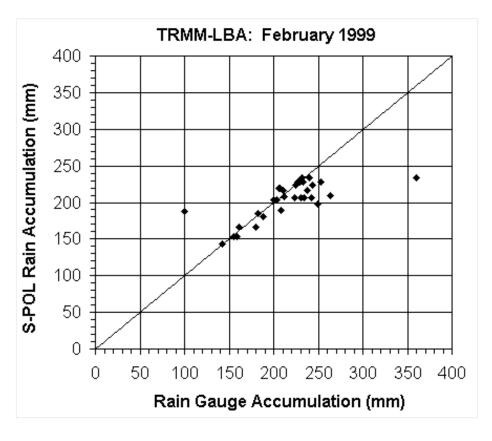






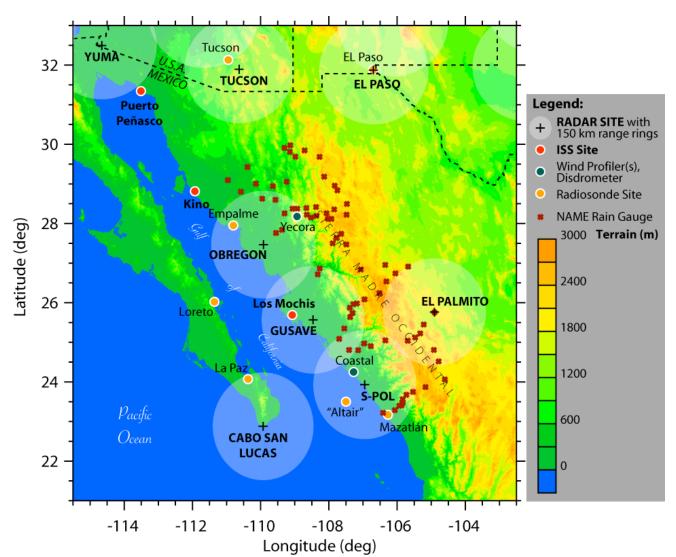


Courtesy R. Cifelli



For non polarimetric radars, use a "pole-tuned" Z-R relationship; did this for NAME and other projects where S-pol was embedded within an operational network

#### NAME Radar Network



#### **Planned**

- -S-Pol
- 4 SMN Radars
- •SMN radars run in full-volume 360s
- •15-min resolution

#### **Actual**

- •S-Pol (7/8-8/21)
- •Cabo (7/15-Fall)
- •Guasave (6/10-Fall)
- •SMN radars single low-level sweep (high temporal resolution)

Courtesy T. Lang

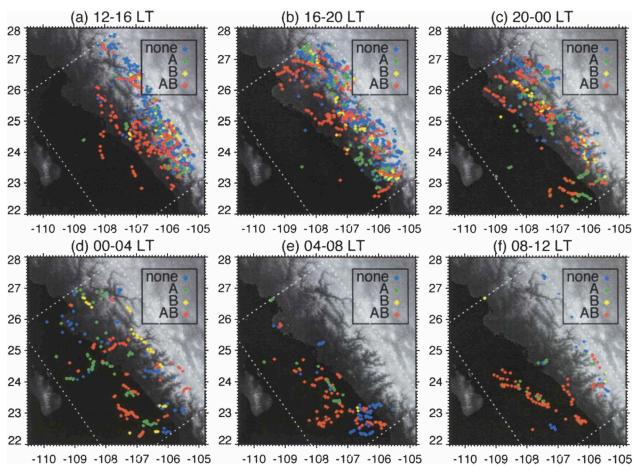


Fig. 11. Centroid locations of organized features as a function of regime (see legends) during the following 4-h periods: (a) 1200–1600, (b) 1600–2000, (c) 2000–0000, (d) 0000–0400, (e) 0400–0800, and (f) 0800–1200 LT. Topography is shaded in the background (grayscale).

### Ground radar based climatology from NAME radar network; Lang et al. (2007)

### Observations

- Dual-Doppler
- Polarimetric radar, ideally not part of the dual-Doppler network
- LMA network
- CG lightning detection network
- Dual-Doppler network should be centered within sounding array for mutual benefit
- Sprite detection via LCD cameras
- Aerosol measurements present a huge challenge
- What are existing quantitative radar and lightning network infrastructures?