

MS-GWaves

Multi-Scale Dynamics of Gravity Waves

U. Achatz

Goethe Universität Frankfurt
and many others:

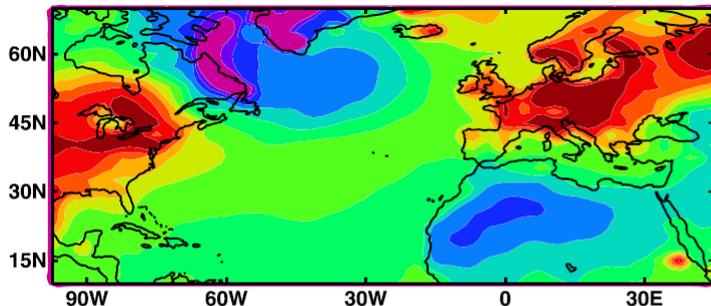
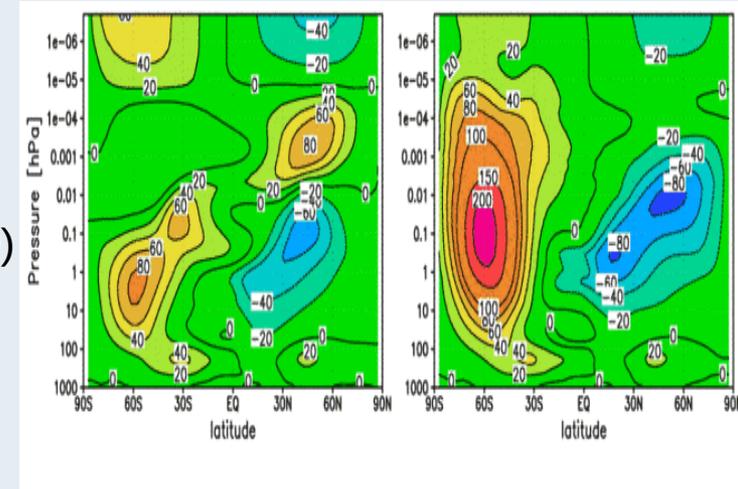
<https://ms-gwaves.iau.uni-frankfurt.de/index.php>



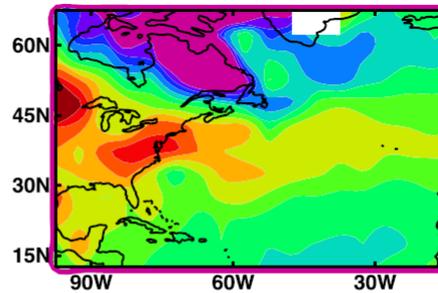
State of the Art: GW Impacts

Gravity-wave effects numerous, e.g.

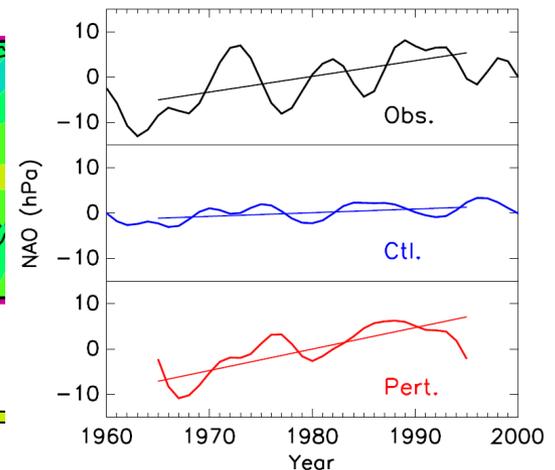
- Clear-air turbulence (e.g. Koch et al 2005)
- **Clouds** (e.g. Zhang et al 2001, 2003, Joos et al 2009)
- **Middle-atmosphere waves** (QBO, solar tides, PWs)
- **residual circulation**
 - GW impact in stratosphere (e.g. Palmer et al 1986)
 - GW control in mesosphere (e.g. Lindzen 1981)
- **Indirectly: Impact middle atmosphere on troposphere (downward control)**



-1.6 -0.8 0 0.8 1.6



-1.6 -0.8 0



Scaife et al (2005)

State of the Art: Parameterization of GW Processes

Sources:

- **Orographic GWs best understood** (Palmer et al 1986, Jiang et al 2002)
- **Convective GWs** (Chun & Baik 1998, Beres et al 2005, Song & Chun 2005, ...)
- **Spontaneous GW emission** (e.g. Plougonven & Zhang 2014)
- **2ndary waves, ...**

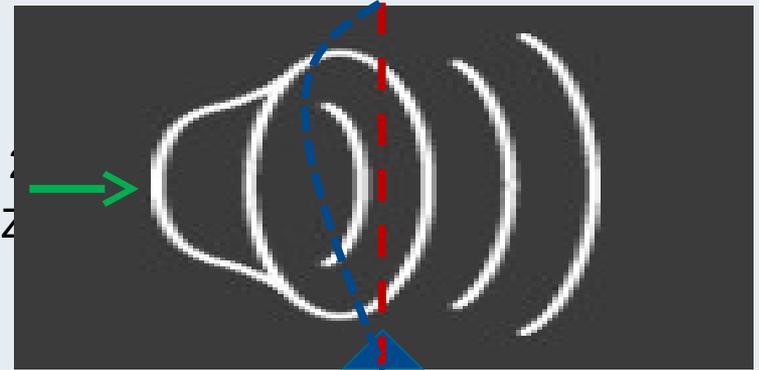
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GW propagation

- Simplifications of WKB Theory (Grimshaw 1975, Achatz et al 2017) for efficiency:
Single-column and steady state limit validity
(e.g. Bühler & McIntyre 2003, Ribstein & Achatz 2016, Bölöni et al 2016)
- Synoptic-scale balanced background assumed
But NWP models resolve some GWs!
- GW propagation through sharp gradients: Tropopause



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GW disipation:

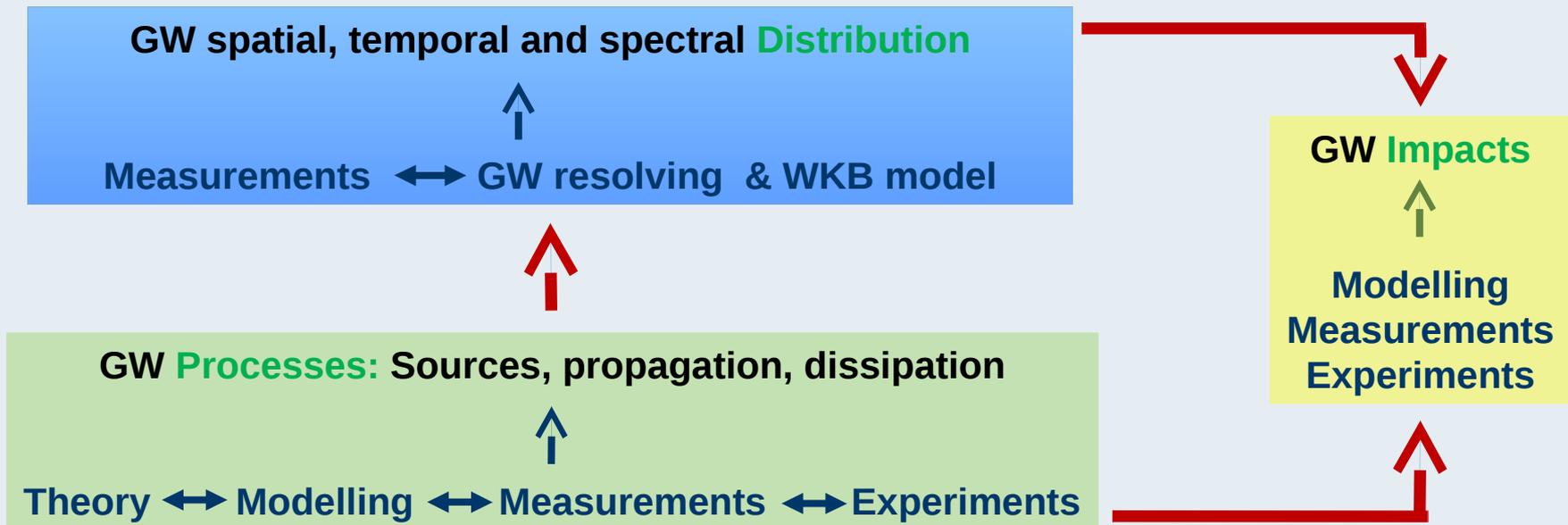
- **Saturation** (Lindzen 1981, ...)
not in agreement with DNS
- **Wave-mean flow interaction** (Dosser & Sutherland 2011, **Bölöni et al 2016**)

Objectives: Central Goals & Key Research Areas

Goals

1. Efficient **parameterization** based on understanding and computational representation of GW processes
2. A **prognostic model** for SGS GWs & implementation into NWP and climate model.

Key Research Areas



D1: Analysis of measurements and weather-service data of the GW distribution

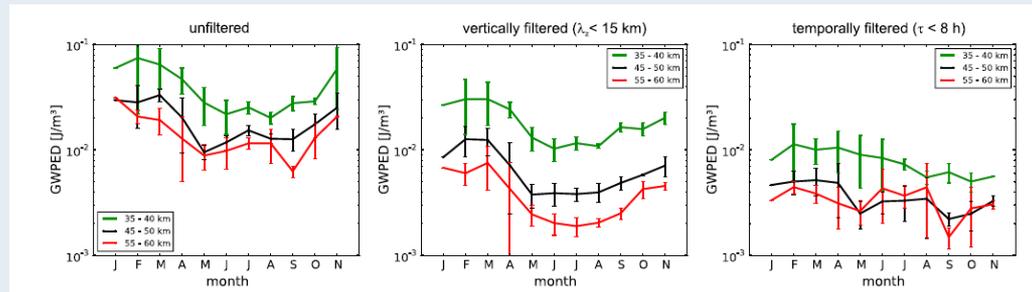
Some examples: Overlap DEEPWAVE & field campaign

- Refraction of GWs into the polar night jet (**Ehard et al 2017**)
- Mountain waves New Zealand (**Portele et al 2017, subm.**)
- **Field campaign** northern Scandinavia winter 2015/16

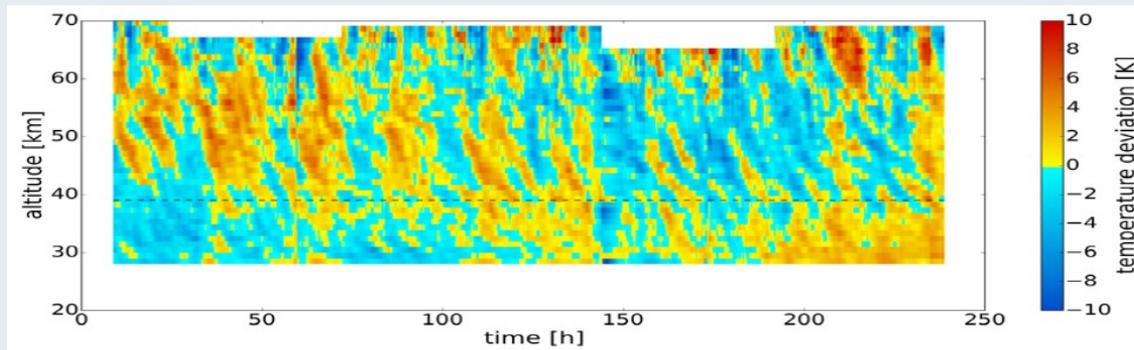
D1: Analysis of measurements and weather-service data of the GW distribution

Some examples: RMR lidar Kühlungsborn

- Climatology T variances (K. Baumgarten et al 2017)



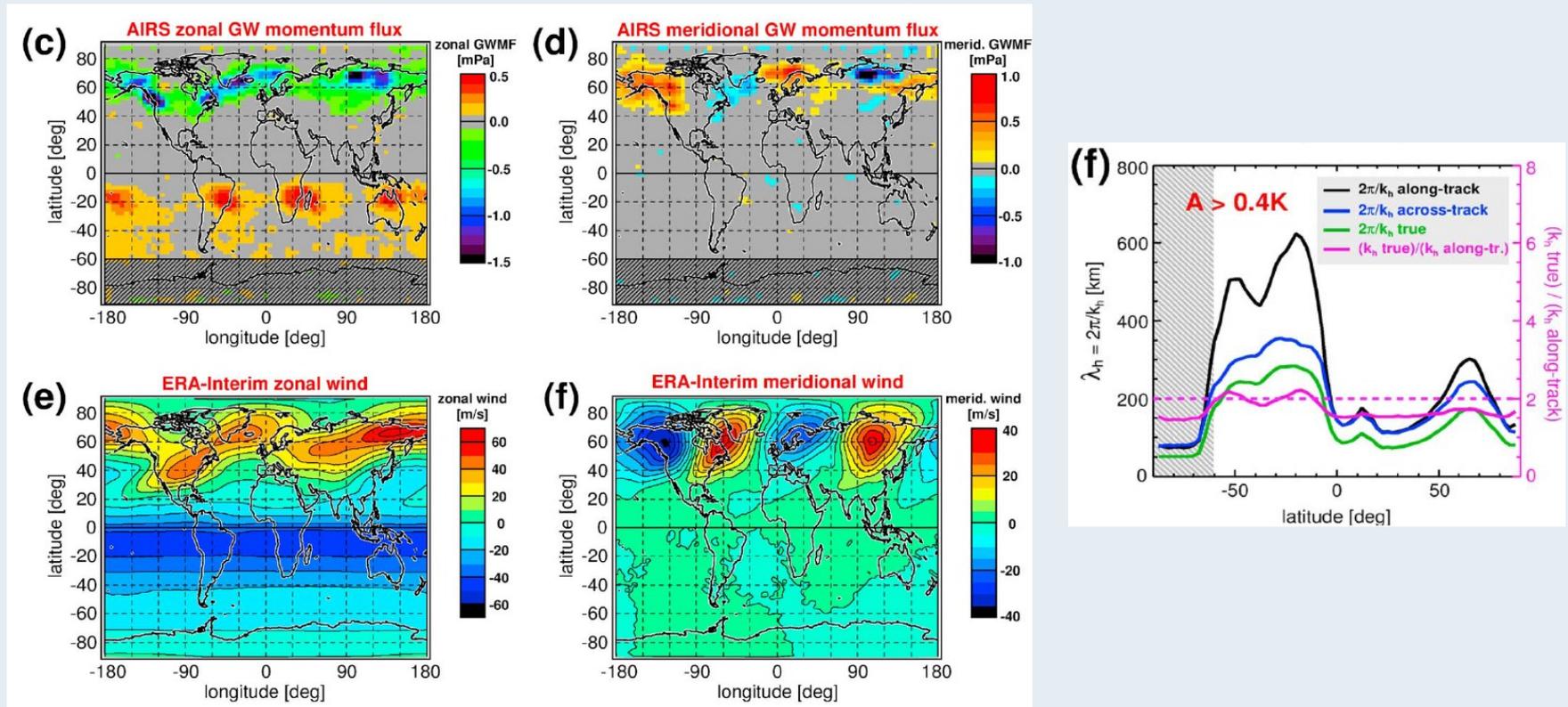
- Unprecedented long data set (4-13 May 2016) (K. Baumgarten et al 2017, subm.)



D1: Analysis of measurements and weather-service data of the GW distribution

Some examples: Satellite data

- Global GW momentum fluxes from AIRS data (Ern et al 2016)



Progress & Results: GW Distribution

Distribution

Impacts

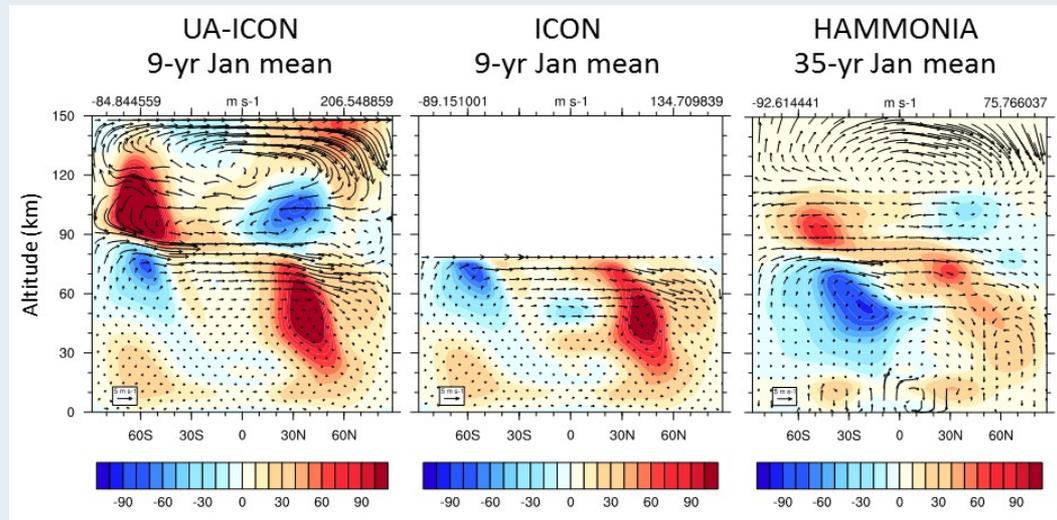
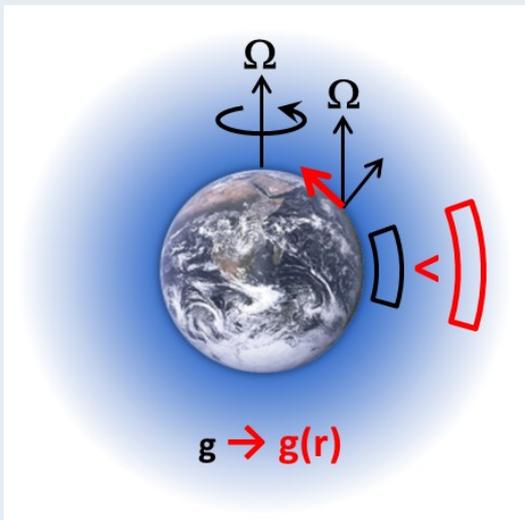


Processes

D2: Non-hydrostatic GW permitting/resolving global model (UA-ICON with MS-GWaM)

Upper-Atmosphere-ICON with standard GW parameterizations (Borchert et al 2017, in prep.)

- height dependence of g
- Coriolis acceleration for all spatial directions
- sphericity changes grid volumes with height
- Development completed (test case and NWP-scores show good results)



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Multi-Scale-Gravity-Wave Model (WKB model)

- implemented (1D, interactive)
- validation in planning (Bölöni et al 2017, in prep.)

Progress & Results: GW Distribution

Distribution

Impacts

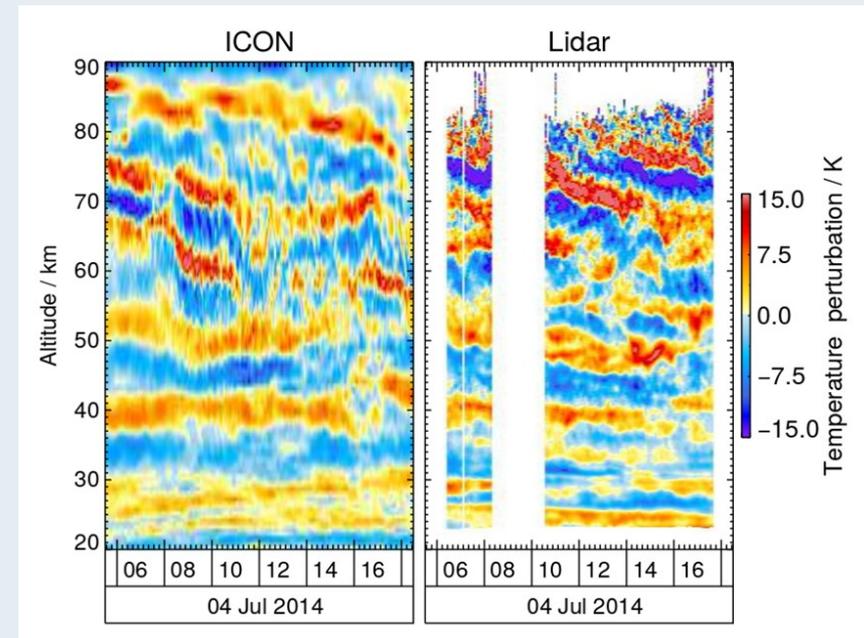
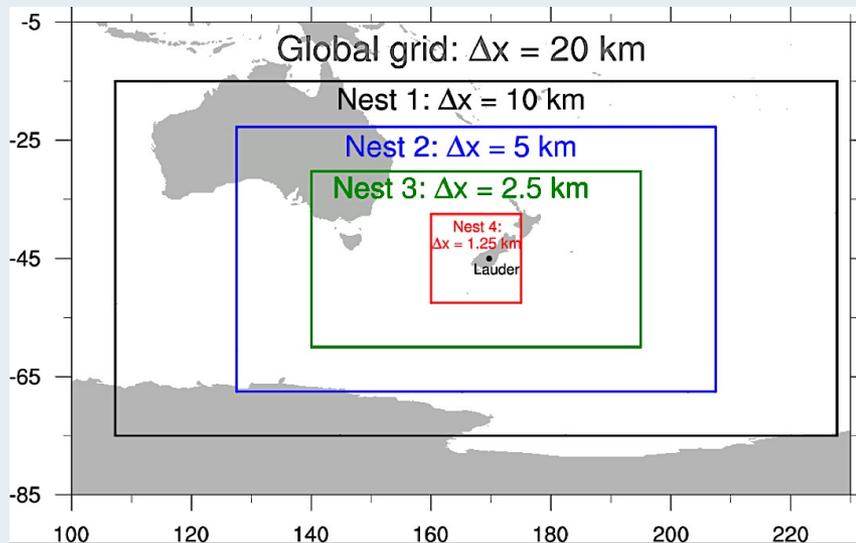


Processes

D3: Validation of the GWs simulated by UA-ICON

Activities so far:

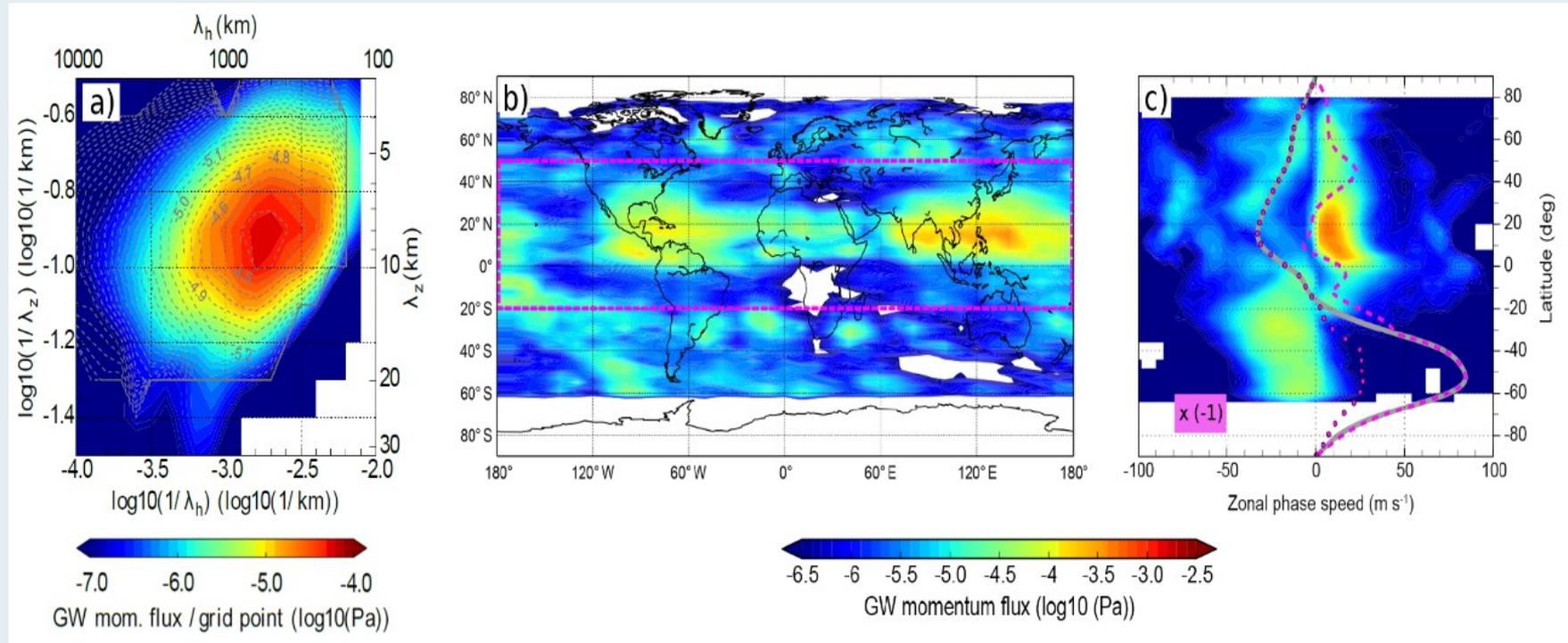
- Implementation of **observational filter** for comparisons against satellite data
- ICON simulations of **campaign episodes** (DEEPWAVE, northern Scandinavia Jan 2016)



P1: GW source processes and their efficient parameterization

Results so far:

- Tuning of convective GW-source parameterization (**Thrinh et al 2016**)



Progress & Results: GW Processes

Distribution

Processes

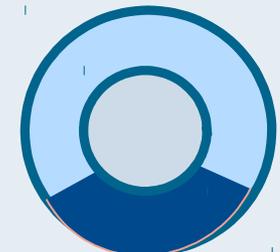
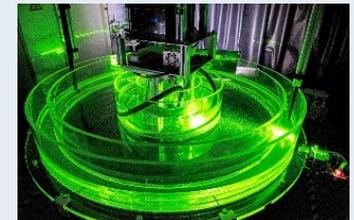
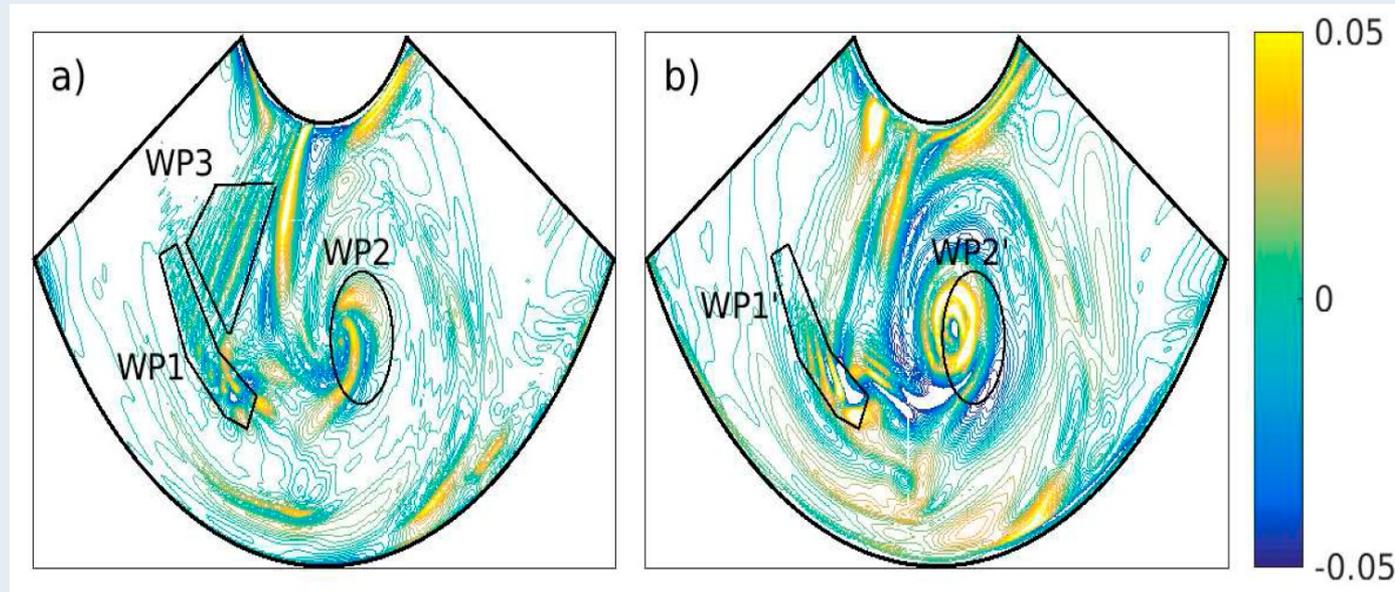
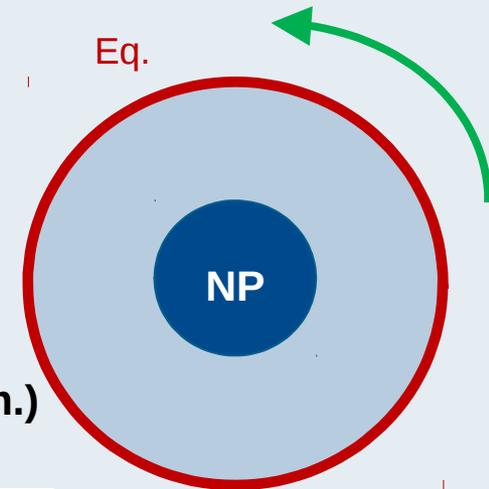
Impacts



P1: GW source processes and their efficient parameterization

Results so far:

- Demonstration/analysis of spontaneous imb. in differentially heated rotating annulus (**Hien et al 2017, revised**)
- ... uses wave analysis tool UWADI (**Schoon & Zülicke 2017, subm.**)



Progress & Results: GW Processes

Distribution

Impacts

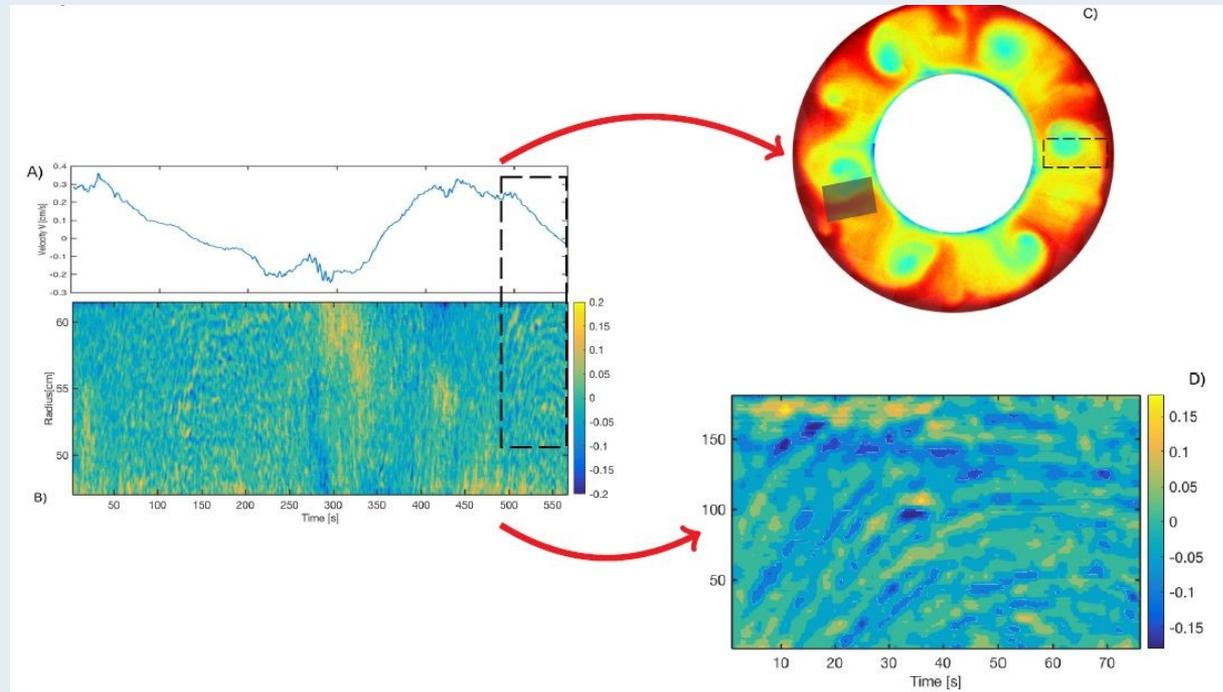
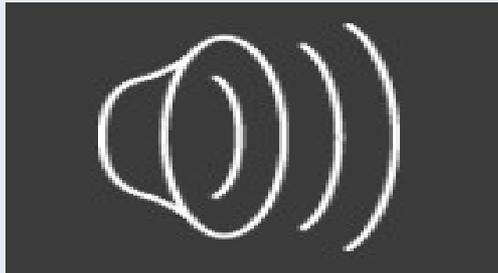
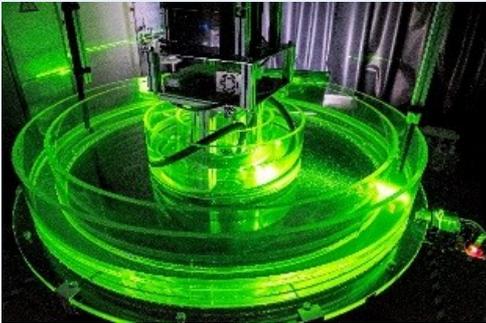


Processes

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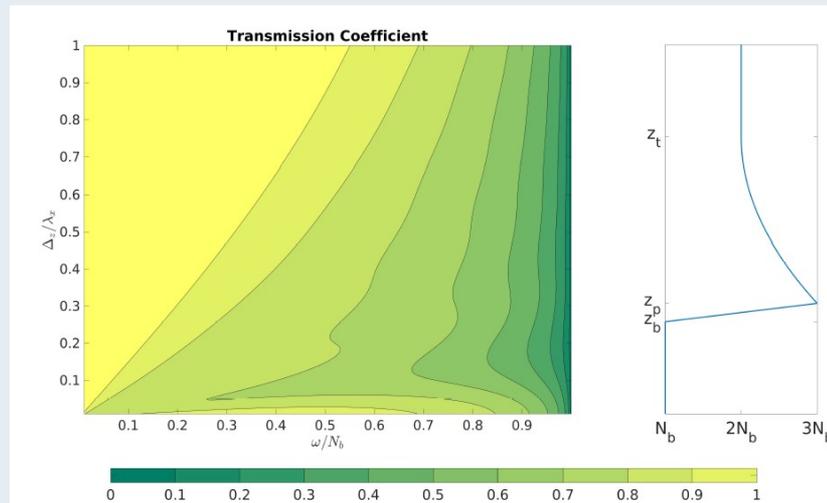
- GWs in the differentially heated annulus (**Rodda et al 2017, subm.**)



P2: GW-mean-flow interactions & Multi-Scale Gravity-Wave Model (MS-GWaM)

Some Results:

- Generalized theory: all stratifications, nonlinear, GMs (Achatz et al 2017)
- Comparison role direct GW-mean-flow interaction with turbulence (Bölöni et al 2016)
- Impact lateral propagation on tides (Ribstein et al 2015, Ribstein & Achatz 2016)
- Interaction sub-mesoscale waves with mesoscale flow (Wilhelm et al 2017, in prep.)
- GW-tropopause interactions (Gisinger et al 2017, **subm.**, Pütz et al 2017, **subm.**)



Progress & Results: GW Processes

Distribution

Impacts

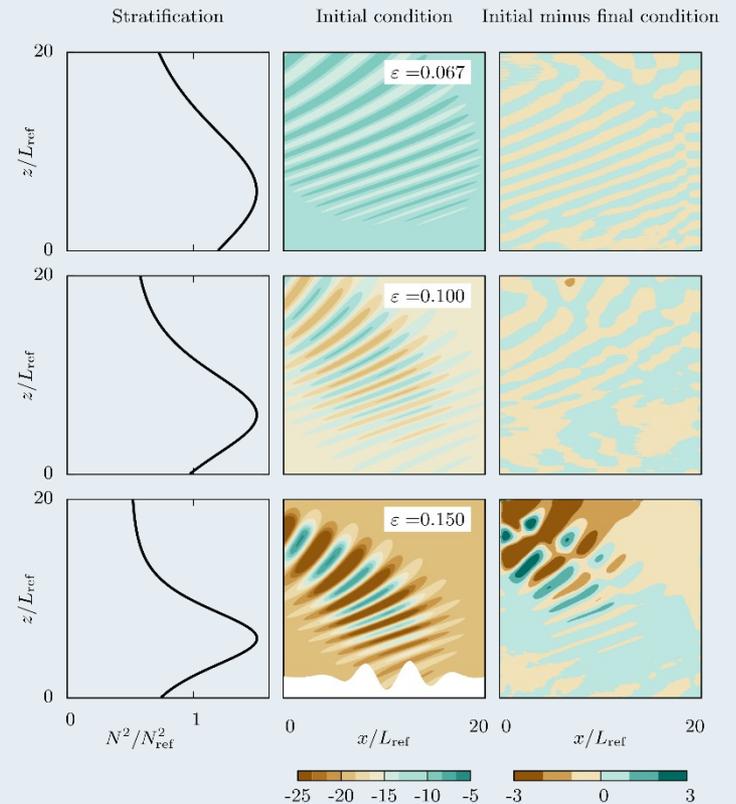
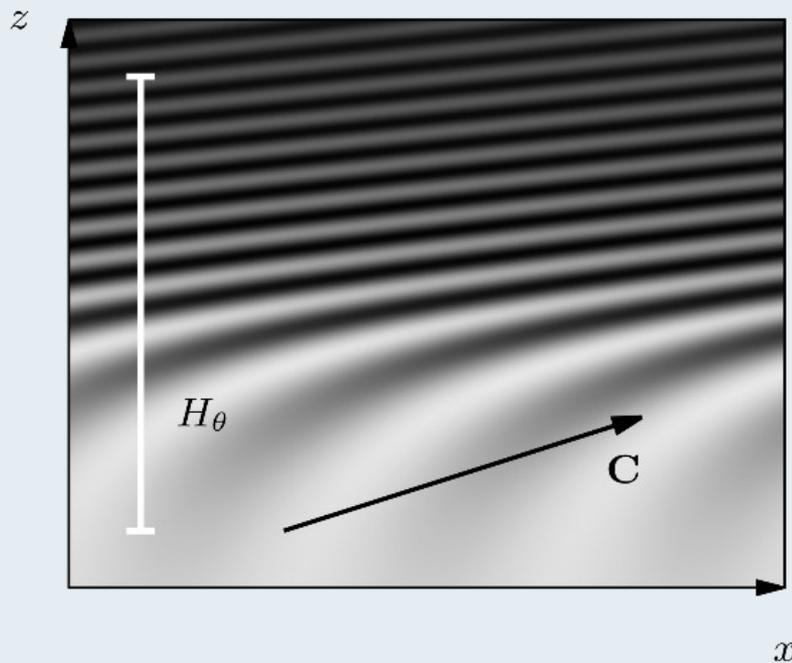


Processes

P3: GW Dissipation

Results:

- Travelling-wave solutions to modulational equations (**Schlutow et al 2017**)
- Stability analysis



Focus:

**Wave-mean-flow Interaction
beyond
traditional parameterization
approaches**

Ray tracing with caustics: Numerics for fully coupled WKB

Classic WKB (Grimshaw 1975, ...) for illustration 1D:
Classic WKB (Grimshaw 1975, ...) for illustration 1D.

Locally monochromatic fields of the form

Locally monochromatic fields of the form $b'(x, t) = \Re B(z, t) e^{i\phi(x, t)}$

local wavenumber and frequency:

local wavenumber and frequency: $k(z, t) = k e_x + m e_z = \nabla\phi, \quad \omega(z, t) = -\partial\phi/\partial t$

wave-action density so that (e.g.)

wave-action density $A(z, t)$ so that (e.g.)

$$E_{GW}(z, t) = A(z, t) \hat{\omega}(m)$$

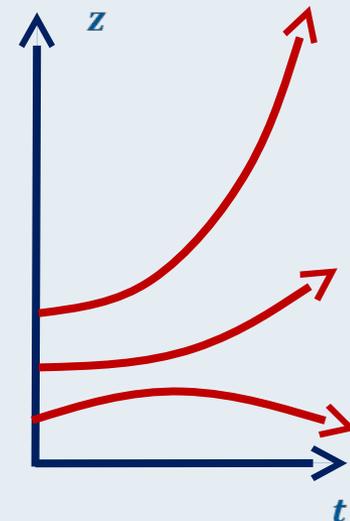
Along rays, defined by

Along rays, defined by

$$dz/dt = c_g$$

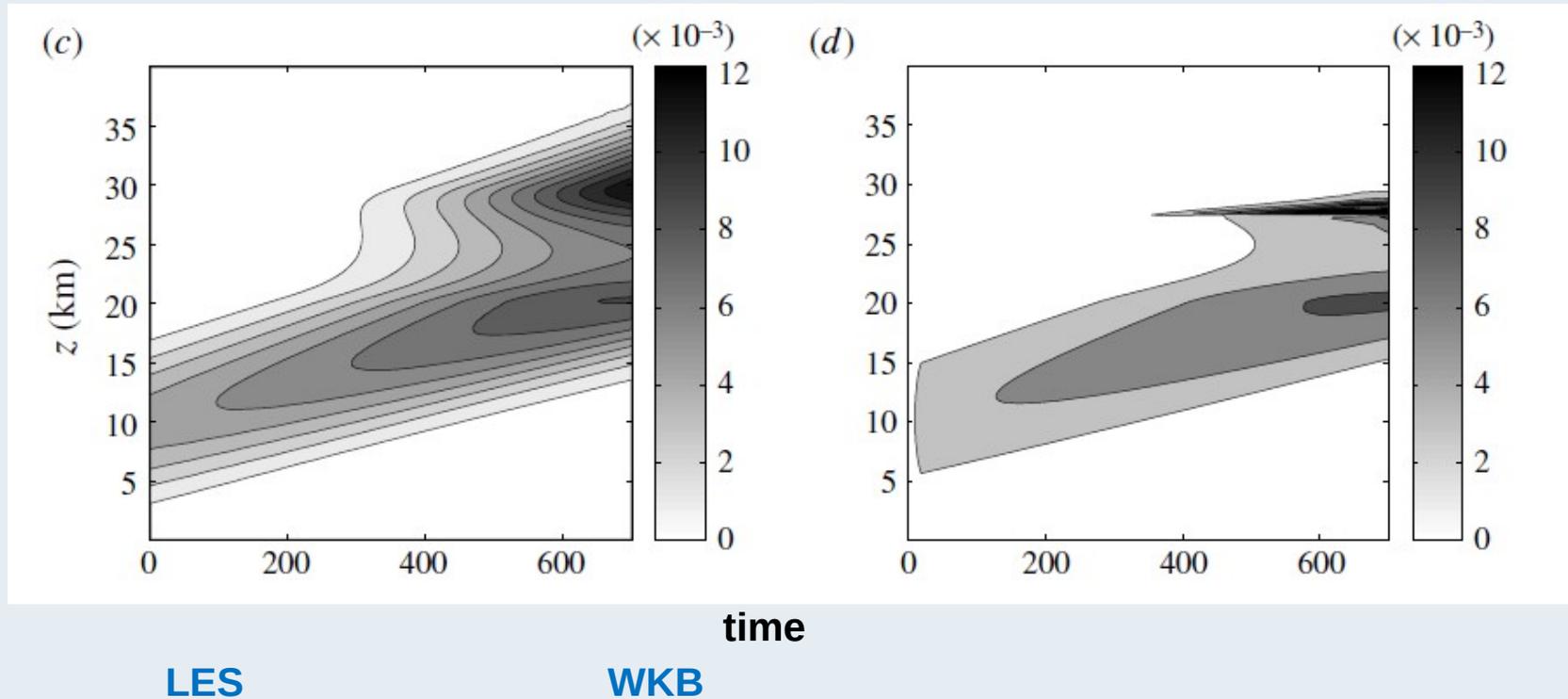
$$\frac{dm}{dt} = -k \frac{\partial U}{\partial z}, \quad \frac{dA}{dt} = -A \frac{\partial c_g}{\partial z}$$

Mean flow:
$$\frac{\partial U}{\partial t} = -\frac{1}{\bar{\rho}} \frac{\partial}{\partial z} (\bar{\rho} \overline{u'w'}) = -\frac{1}{\bar{\rho}} \frac{\partial}{\partial z} (c_g k A)$$



Ray tracing with caustics: Stability Problem

GW packet refracted by a jet



Rieper et al (2013)

Ray tracing with caustics: Uniqueness Problem

Locally monochromatic fields

wave-action density so that (e.g.)
 wave-action density $A(z, t)$ so that (e.g.)

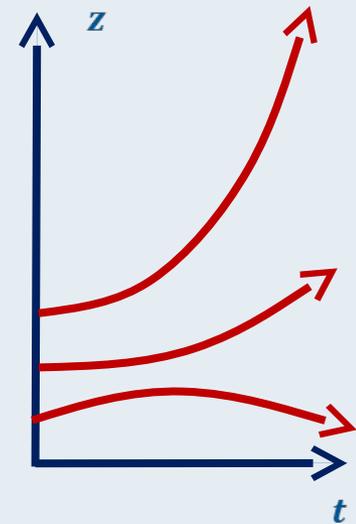
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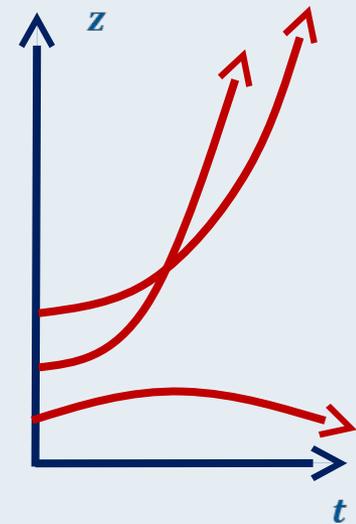
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Crossing rays (caustics): uniqueness problem for A and m !
 Mean flow: $\frac{\partial U}{\partial t} = -\frac{1}{\bar{\rho}} \frac{\partial}{\partial z} (\bar{\rho} u w) = -\frac{1}{\bar{\rho}} \frac{\partial}{\partial z} (c_g k A)$

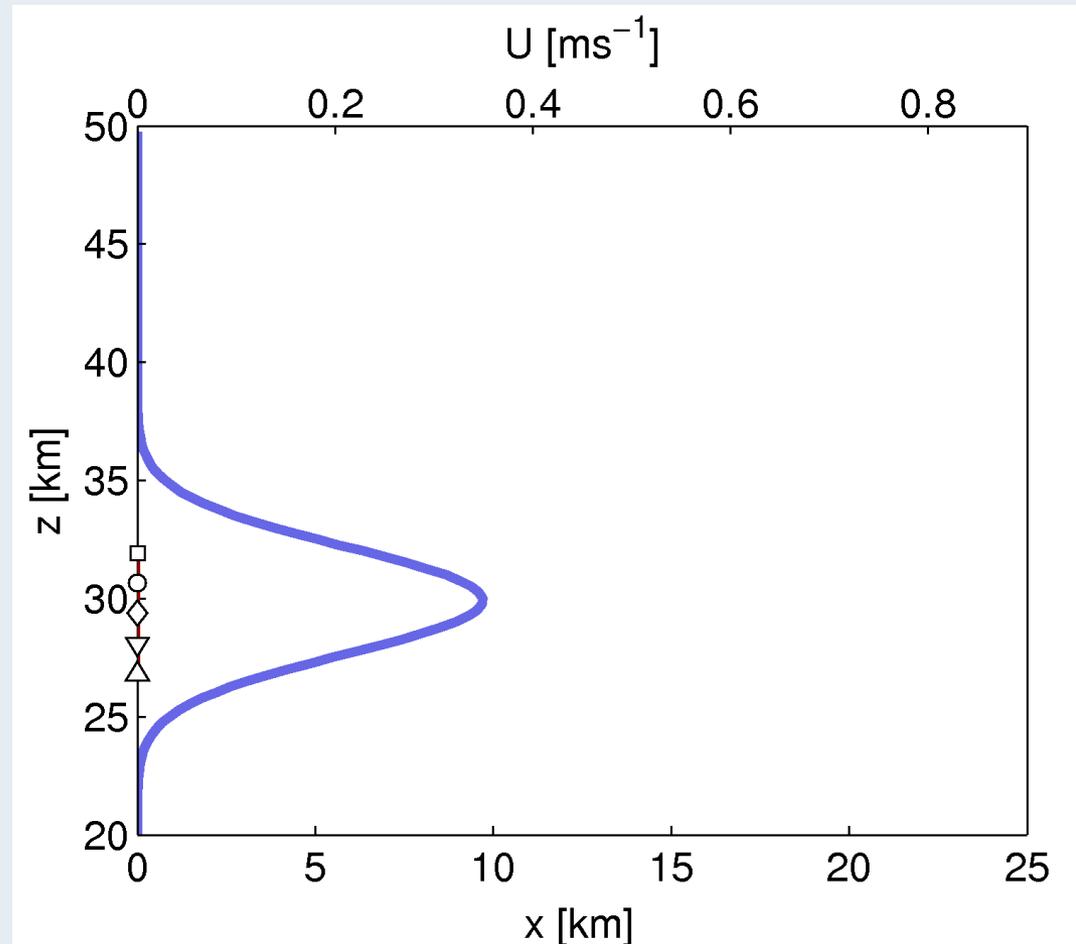
Crossing rays (caustics): uniqueness problem for A and m !



Ray tracing with caustics: examples for caustic situations

Nonuniqueness of wave number and wave-action density arises easily:

e.g. by
wave-induced mean flow



Ray tracing with caustics: spectral approach

linear limit: wave field can be decomposed into fields with single valued wavenumbers

spectral description in phase space (Dewar 1970, Dubrulle & Nazarenko 1997, Bühler & McIntyre 1999, Hertzog et al 2000, Muraschko et al 2015) does this automatically

phase space wave action density

$$\mathcal{N}(m, z, t) = \int d\alpha A_\alpha(z, t) \delta[m - m_\alpha(z, t)] \quad \Leftrightarrow \quad A(z, t) = \int dm \mathcal{N}(m, z, t)$$

satisfies conservation equation

$$\frac{\partial \mathcal{N}}{\partial t} + \frac{\partial}{\partial z} (c_g \mathcal{N}) + \frac{\partial}{\partial m} (\dot{m} \mathcal{N}) = 0 \quad \dot{m} = -k \frac{\partial U}{\partial z}$$

Mean flow:
$$\frac{\partial U}{\partial t} = -\frac{1}{\bar{\rho}} \frac{\partial}{\partial z} (\bar{\rho} \overline{u'w'}) = -\frac{1}{\bar{\rho}} \frac{\partial}{\partial z} \left(\int dm c_g k \mathcal{N} \right)$$

generalization to 3D straightforward

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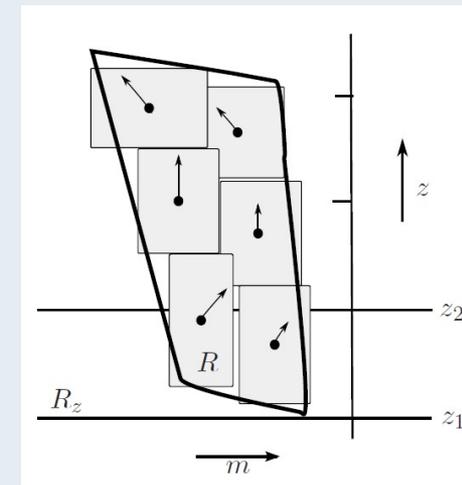
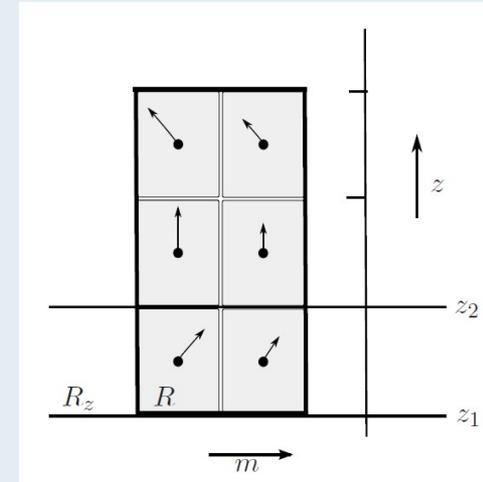
Ray tracing with caustics: efficient numerics (Muraschko et al 2015)

phase-space velocity is non-divergent

hence
$$\frac{\partial c_g}{\partial z} + \frac{\partial \dot{m}}{\partial m} = \frac{\partial}{\partial z} \frac{\partial \Omega}{\partial m} + \frac{\partial}{\partial m} \left(-\frac{\partial \Omega}{\partial z} \right) = 0$$

- flow is volume preserving
- rays cannot cross
- Wave-action density conserved on rays
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- region of nonzero \mathcal{N} approximated by rectangular ray volumes
- ray volumes move with central ray
- ray volumes change height (Δz) and width (Δm) in area-preserving manner

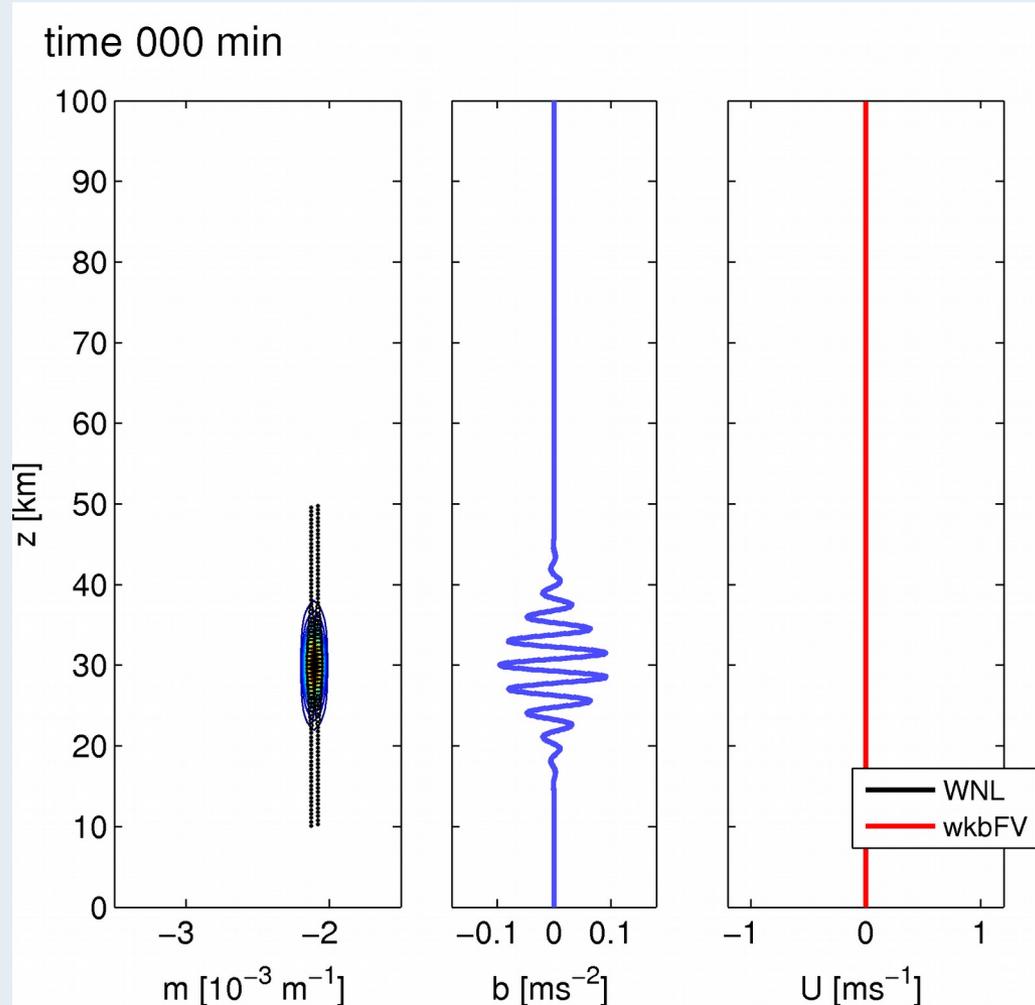


Ray tracing with caustics: efficient numerics (Muraschko et al 2015)

hydrostatic wave packet
(Boussinesq)

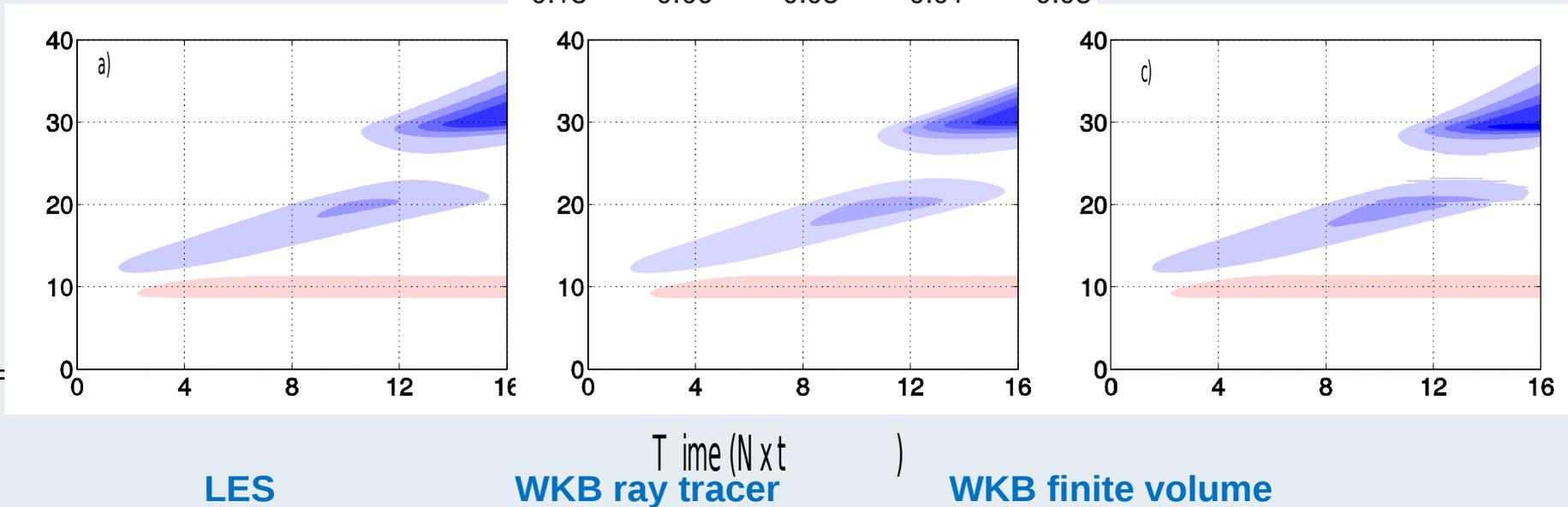
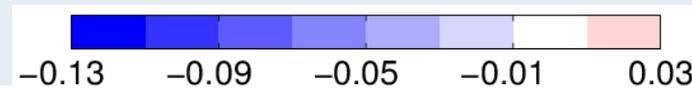
Rays are no wavepackets

No turbulence taken into
account!



Ray tracing with caustics: no numerical instabilities (Bölöni et al 2016)

GW packet refracted by a jet



Direct wave-mean-flow interaction: comparison with role of wave breaking

- **transient GWs** can interact with the mean flow without the onset of turbulence (eg Dosser & Sutherland 2011)
- GW parameterizations (steady-state approximation) only rely on **wave breaking**

comparative role of wave transience (direct interaction) vs wave breaking?

direct wave-mean-flow interaction vs wave breaking (Bölöni et al 2016)

horizontally infinite GW packets in interaction with mean flow

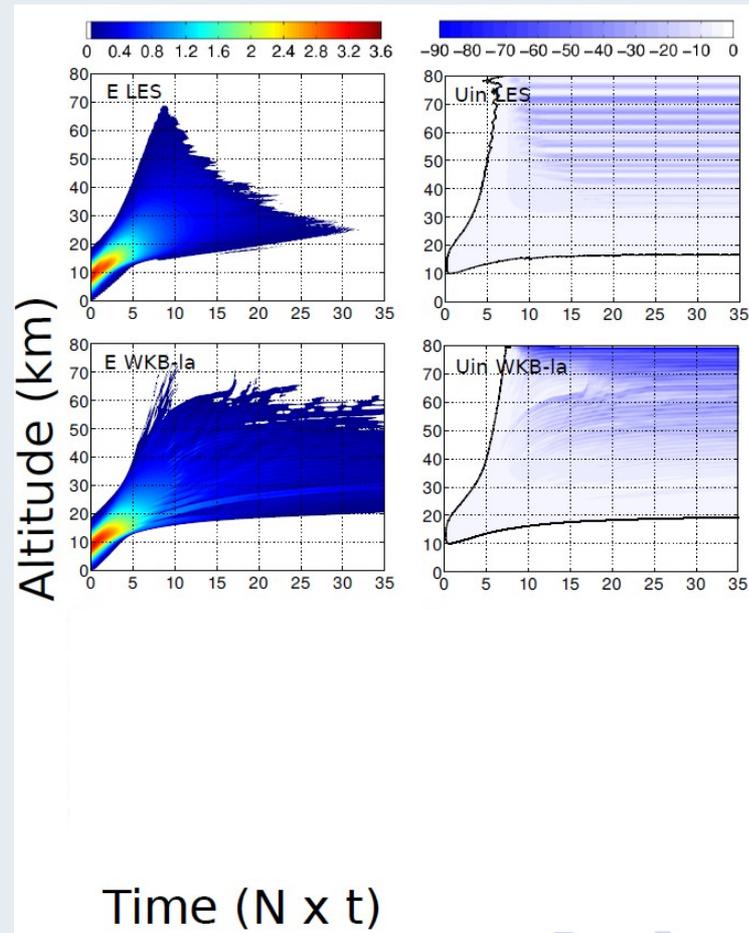
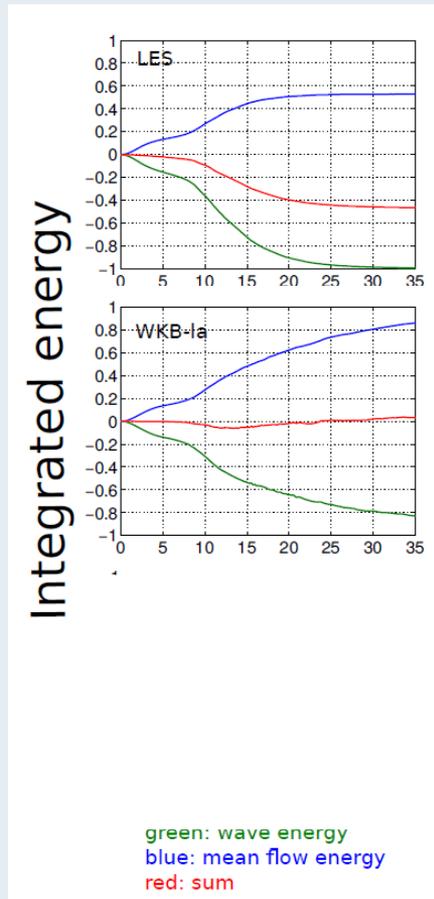
- $1D$: $U(z, t), A(z, t), m(z, t)$
- direct GW-mean-flow interaction always active
- WKB: $E_{mean} + E_{wave} = const.$

tools:

- wave resolving **LES** (reference data)
- fully coupled **WKB** (reference data)
- turbulence onset
- turbulence onset stability threshold can be surpassed
 - once static instability threshold can be surpassed
 - parameter accounting for phase cancellations between spectral components
 - (scale selective) $\frac{eddy\ viscosity/diffusivity}{\rho \omega^2} N f(m) > \alpha N^2$ reduces wave amplitude to inst. threshold
 - parameter $\alpha \in [1,2]$ accounting for phase cancellations between spectral components
 - (scale selective) **eddy viscosity/diffusivity** reduces wave amplitude to inst. threshold

direct wave-mean-flow interaction vs wave breaking (Bölöni et al 2016)

static instability hydrostatic wave packet

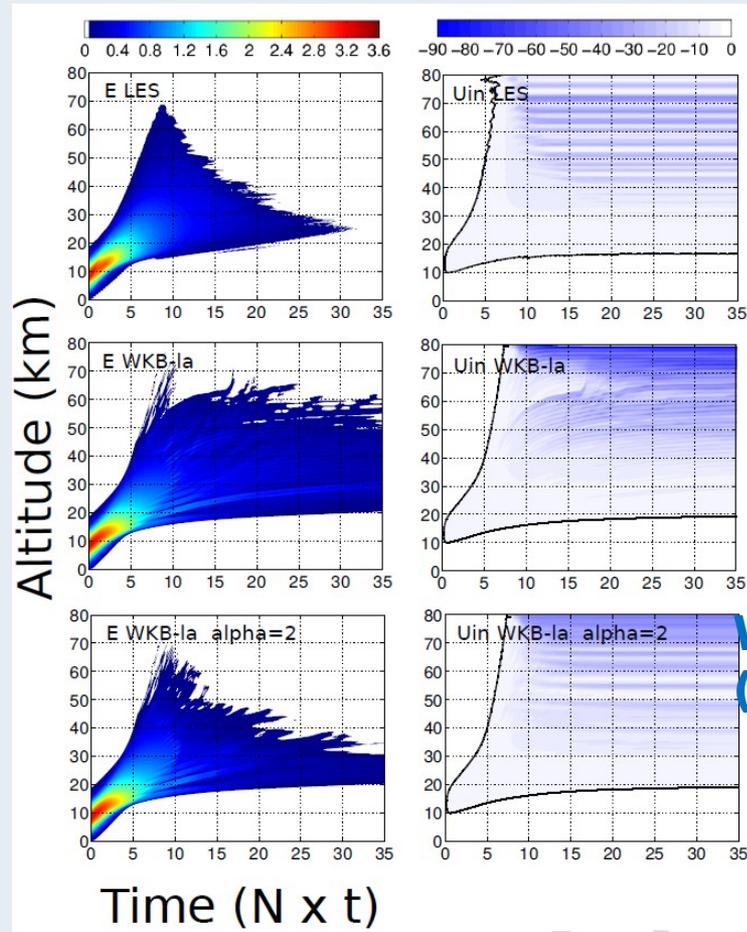
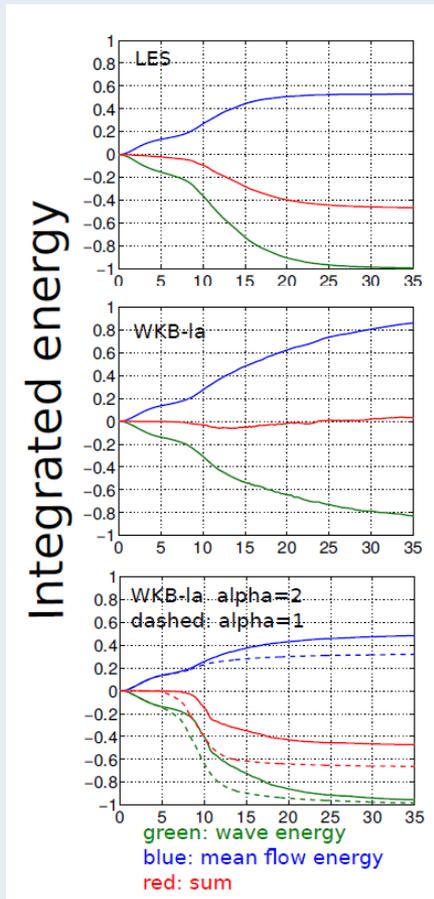


LES
(wave-resolving)

WKB

direct wave-mean-flow interaction vs wave breaking (Bölöni et al 2016)

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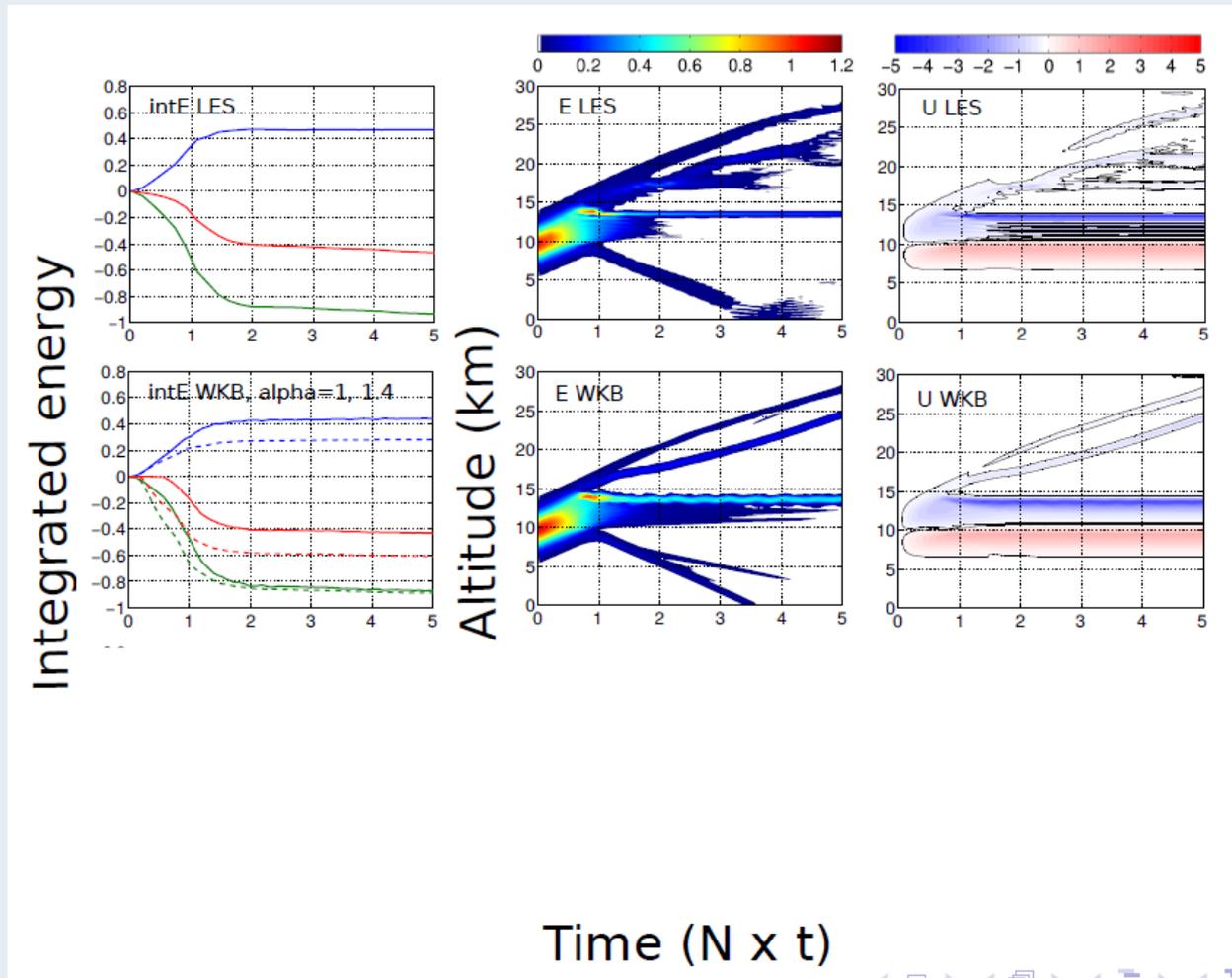
LES
(wave-resolving)

WKB

WKB with saturation
(turbulence param.)

direct wave-mean-flow interaction vs wave breaking (Bölöni et al 2016)

static instability non-hydrostatic wave packet

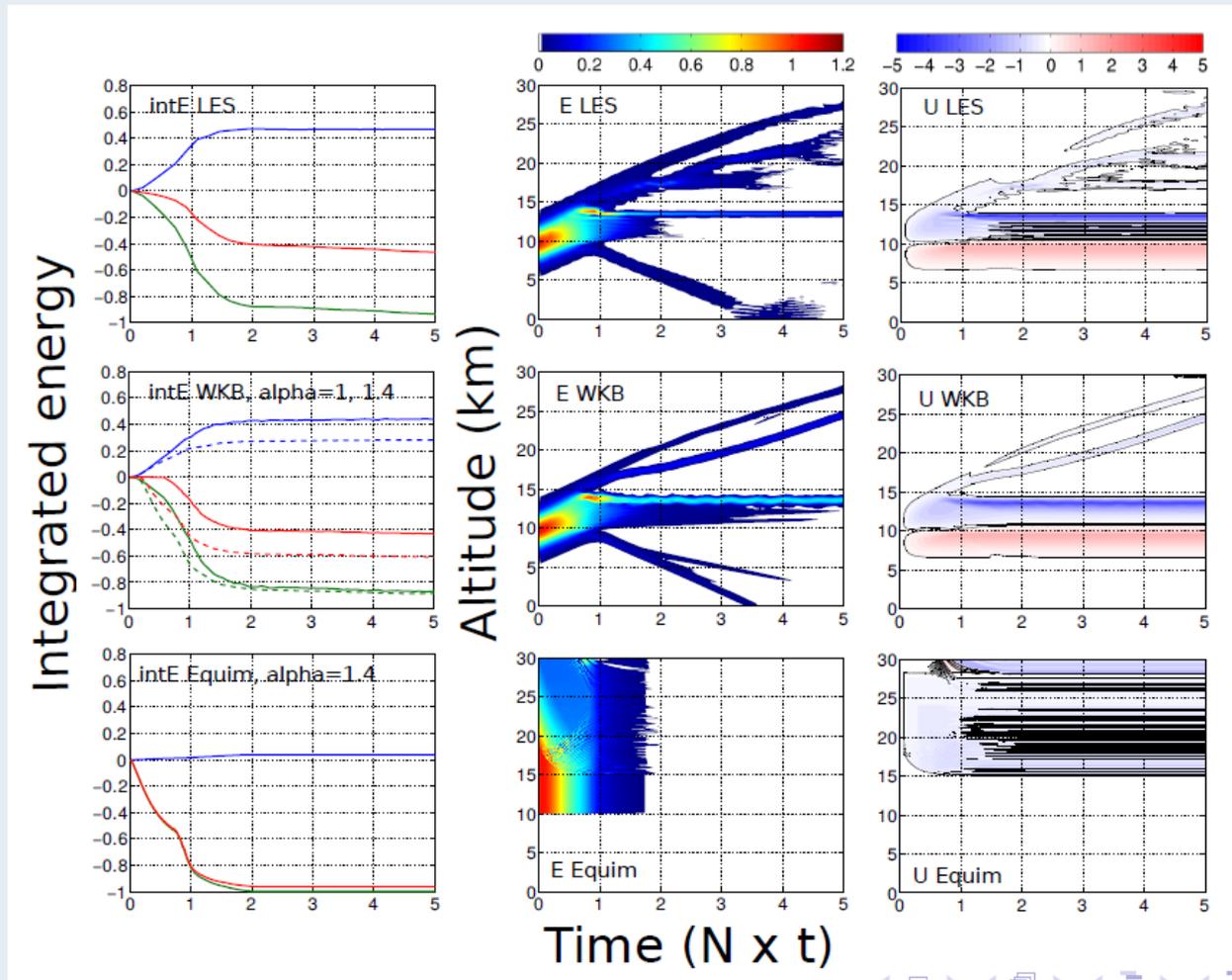


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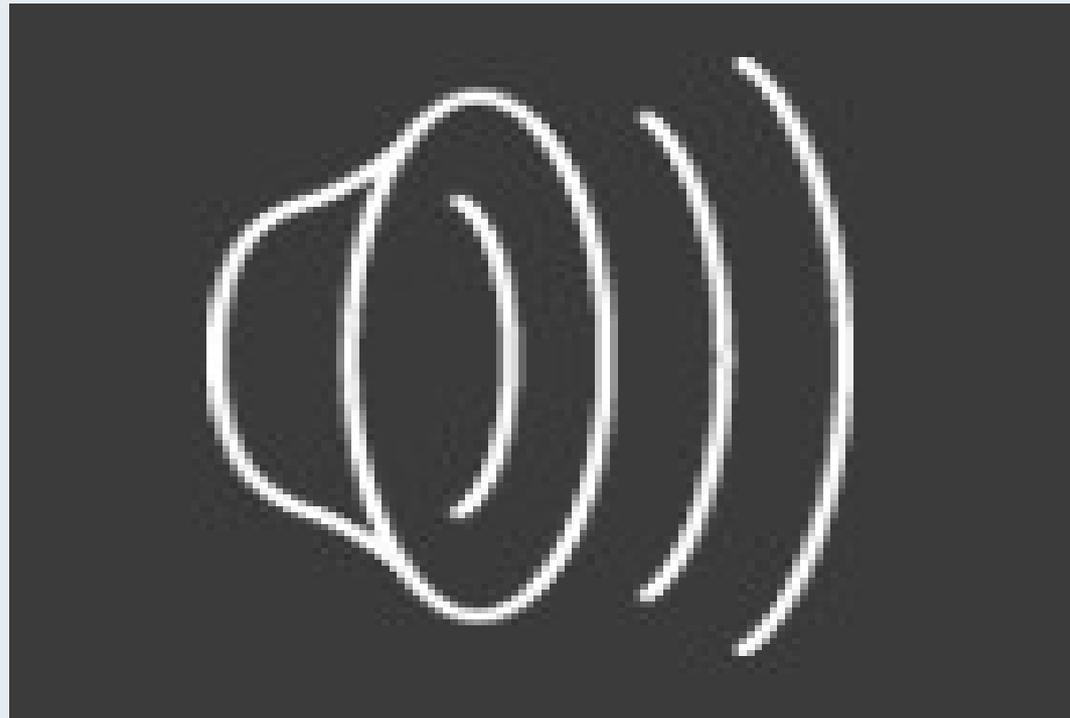


LES
(wave-resolving)

WKB with saturation
(turbulence param.)

steady-state
(GW parameterization)

Large-scale waves forced by the diurnal cycle of solar heating

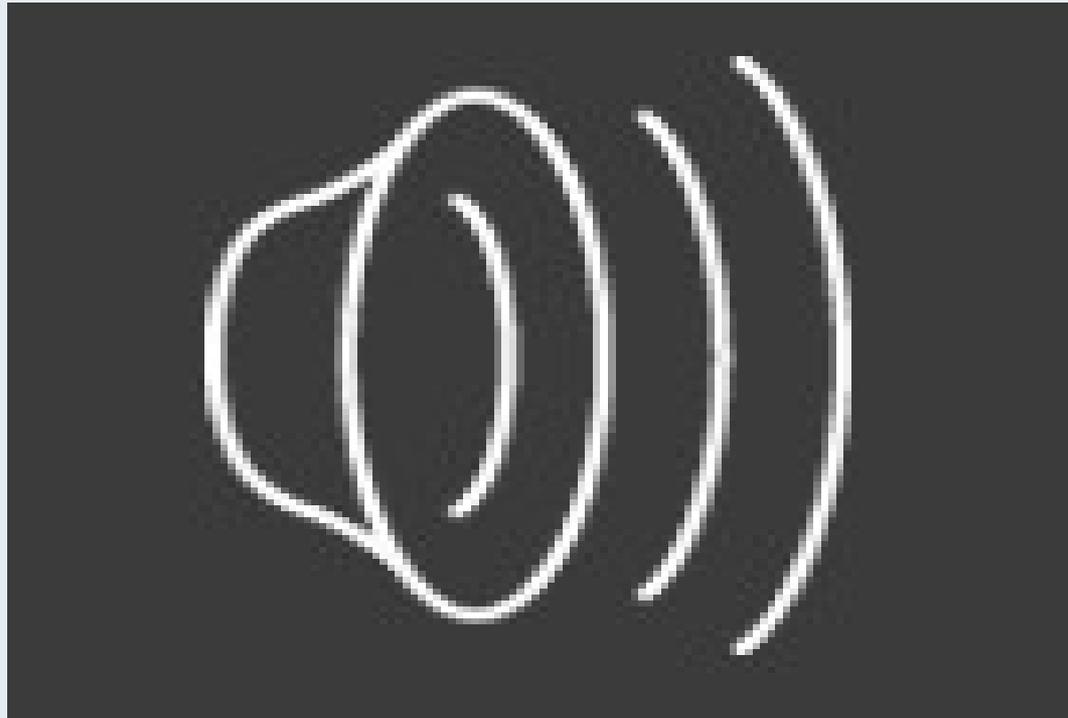


Solar tides

Large-scale waves forced by the diurnal cycle of solar heating

Two components:

- **Migrating tides**
follow solar movement

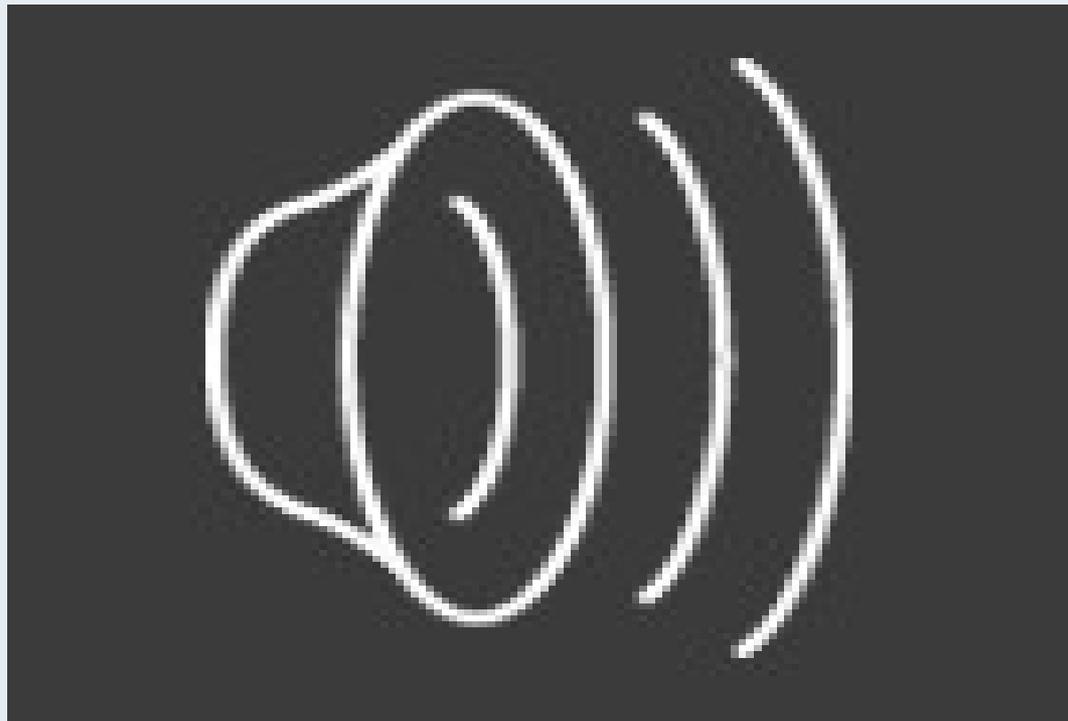


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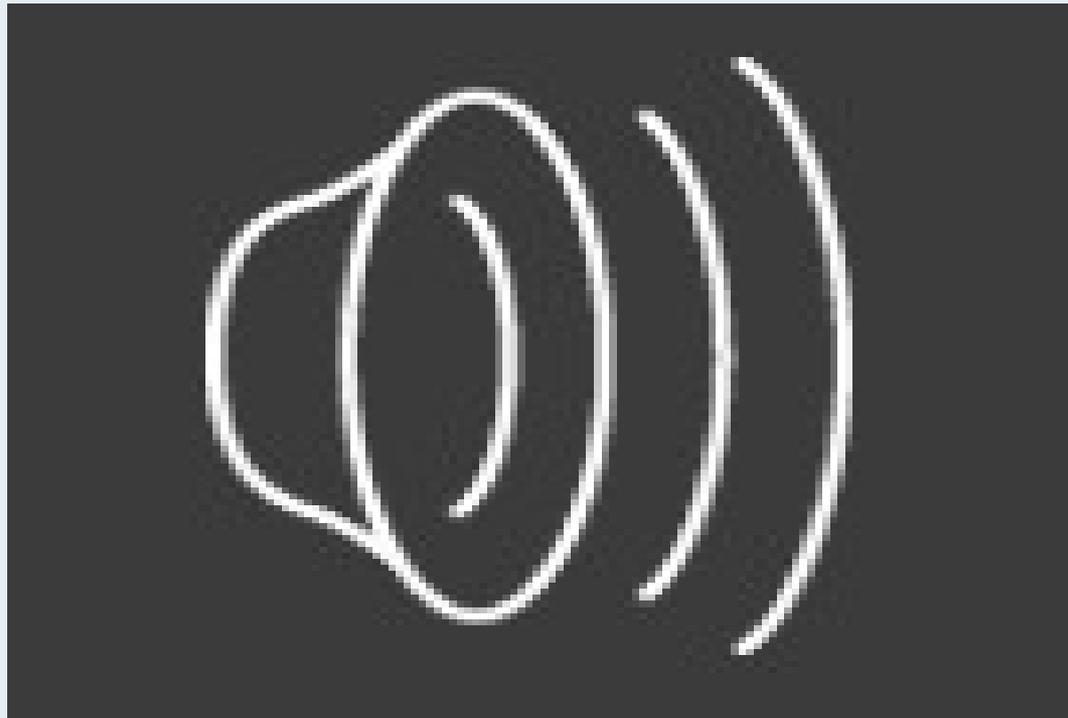


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all the rest



Interaction with GWs:

- STs influence **GW propagation and amplitude development**
- GW impact on STs by **GW momentum and buoyancy deposition**

Tidal model in interaction with GWs (Ribstein et al 2015, Ribstein & Achatz 2016)

From **GCM data** (HAMMONIA, Schmidt et al 2006):

- **Seasonally dependent reference climatology** $\bar{u}(\lambda, \phi, z), \bar{T}(\lambda, \phi, z)$
- **Diurnal heating cycle** $\Re \sum_n Q_n(\lambda, \phi, z) e^{in\Omega t}$

Linear model (Achatz et al 2008, based on KMCM, Becker and Schmitz 2003)

$$\begin{aligned} \mathbf{u} &= \bar{\mathbf{u}} + \mathbf{u}'(\lambda, \phi, z, t) \\ T &= \bar{T} + T'(\lambda, \phi, z, t) \end{aligned}$$

GW fluxes from **4D WKB model with rays propagating on**
First implementation of a fully coupled transient ray tracer into a global model.

$$\left(\frac{\partial}{\partial t} + \bar{\mathbf{u}} \cdot \nabla_h \right) \mathbf{u}' + \dots = -\frac{1}{\bar{\rho}} \nabla \cdot (\bar{\rho} \overline{\mathbf{v}_{GW} \mathbf{u}_{GW}})$$

GW fluxes from **4D WKB model with rays propagating on** $(\bar{\mathbf{u}} + \mathbf{u}', \bar{T} + T')$

First implementation of a fully coupled transient ray tracer into a global model

Tidal model in interaction with GWS (Ribstein et al 2015, Ribstein & Achatz 2016)

3D effects (beyond single column)

- Horizontal GW propagation

- Horizontal gradients in reference climatology and tides

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- Horizontal GW flux convergence, $\frac{dk_h}{dt} = -k \frac{d}{dz} (\bar{u} + u')$, $\frac{dm}{dt} = -k \frac{d}{dz} (\bar{u} + u') - l \frac{d}{dz} (\bar{v} + v')$

- Horizontal GW flux convergence

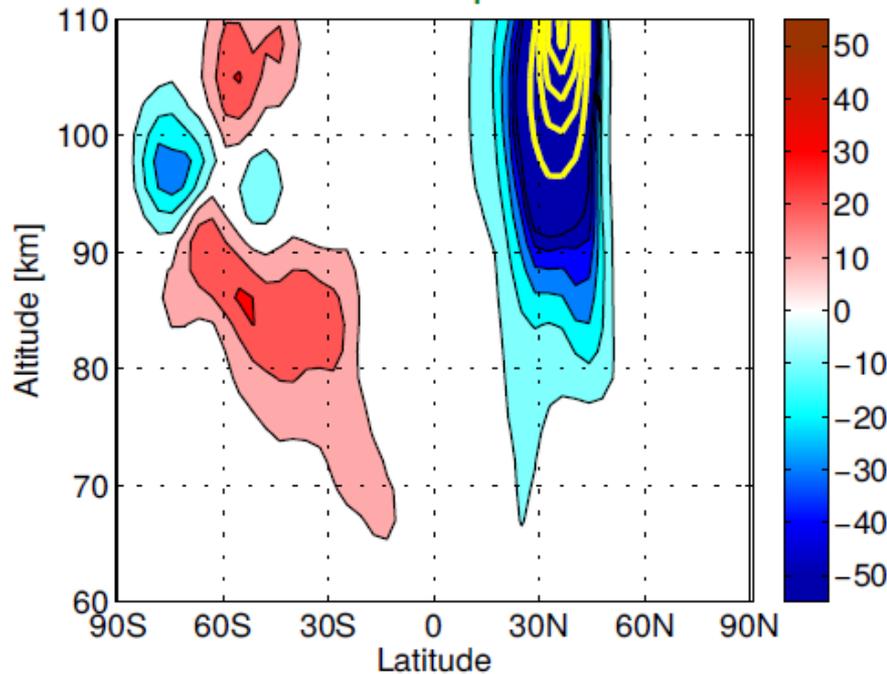
$$\left(\frac{\partial}{\partial t} + \bar{\mathbf{u}} \cdot \nabla_h \right) \mathbf{u}' + \dots = -\frac{1}{\bar{\rho}} \frac{\partial}{\partial z} (\bar{\rho} \overline{\mathbf{w}_{GW} \mathbf{u}_{GW}}) - \frac{1}{\bar{\rho}} \nabla_h \cdot (\bar{\rho} \overline{\mathbf{u}_{GW} \mathbf{u}_{GW}})$$

$$\left(\frac{\partial}{\partial t} + \bar{\mathbf{u}} \cdot \nabla_h \right) T' + \mathbf{v}' \cdot \nabla \bar{T} + \dots = \Re \sum_n Q_n(\lambda, \phi, z) e^{in\Omega t} - \nabla_h \cdot (\overline{\mathbf{u}_{GW} T_{GW}})$$

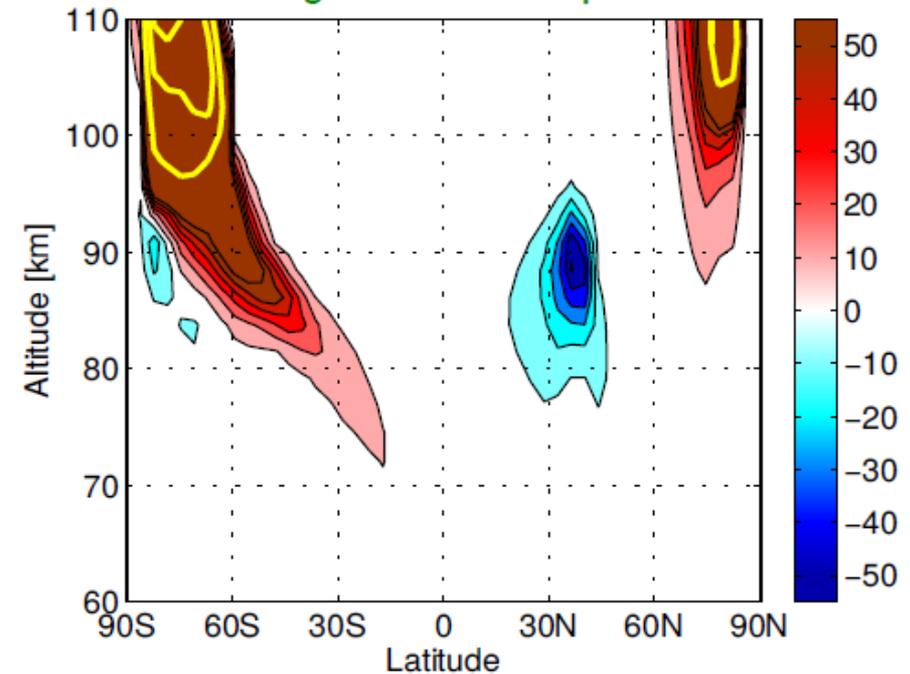
Tidal model in interaction with GWs (Ribstein et al 2015, Ribstein & Achatz 2016)

3D effects (beyond single column)

“Full” experiment



“Single-column” experiment

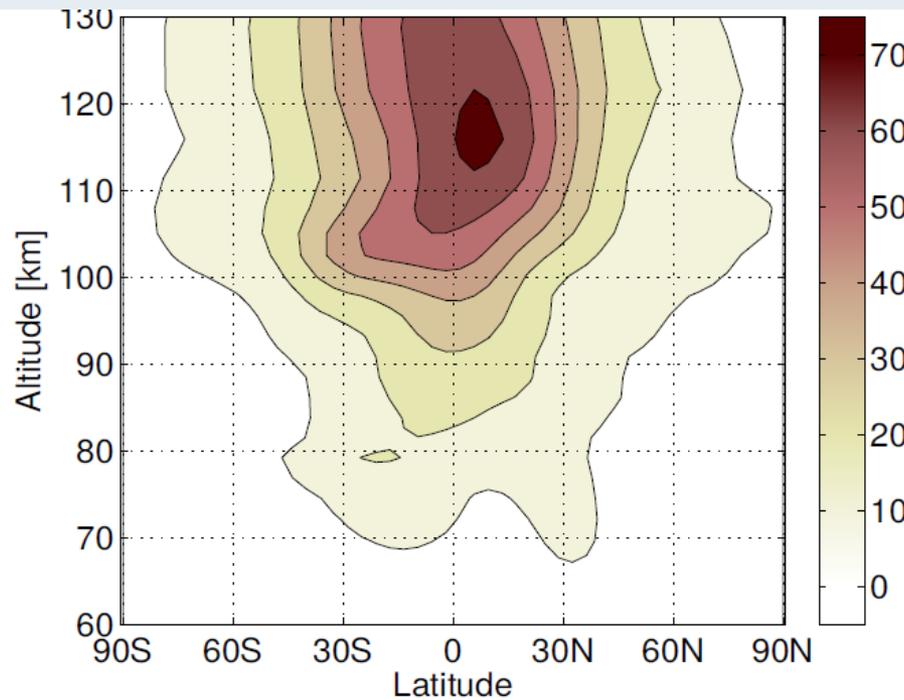


zonal-mean daily-mean GW forcing (December)

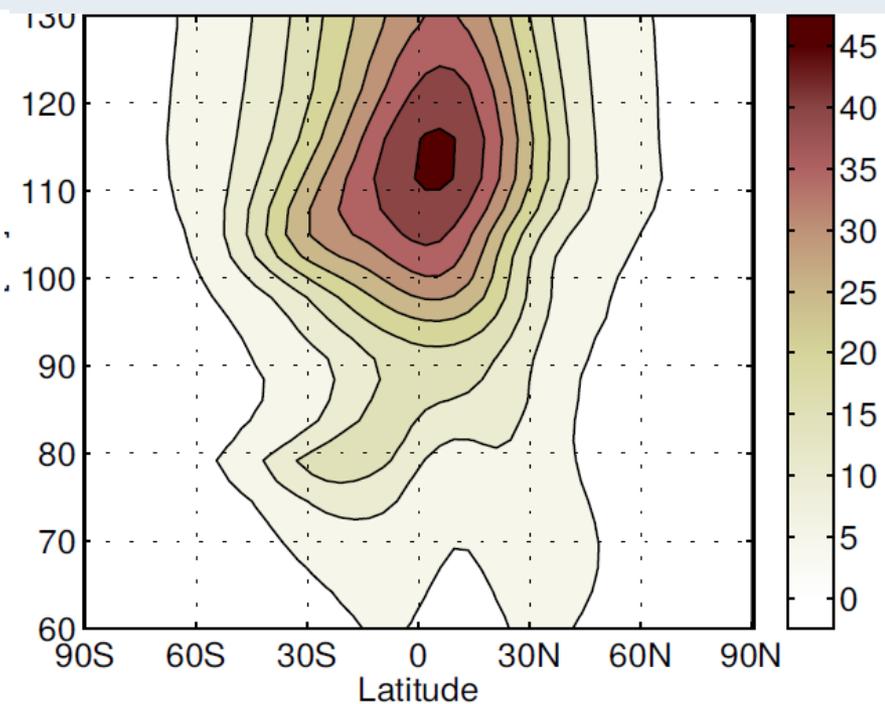
Tidal model in interaction with GWS (Ribstein et al 2015, Ribstein & Achatz 2016)

3D effects (beyond single column)

“Full” experiment



“Single-column” experiment



Tidal model in interaction with GWs (Ribstein et al 2015, Ribstein & Achatz 2016)

3D effects (beyond single column):

• 3D effects (beyond single column):

• Tidal model

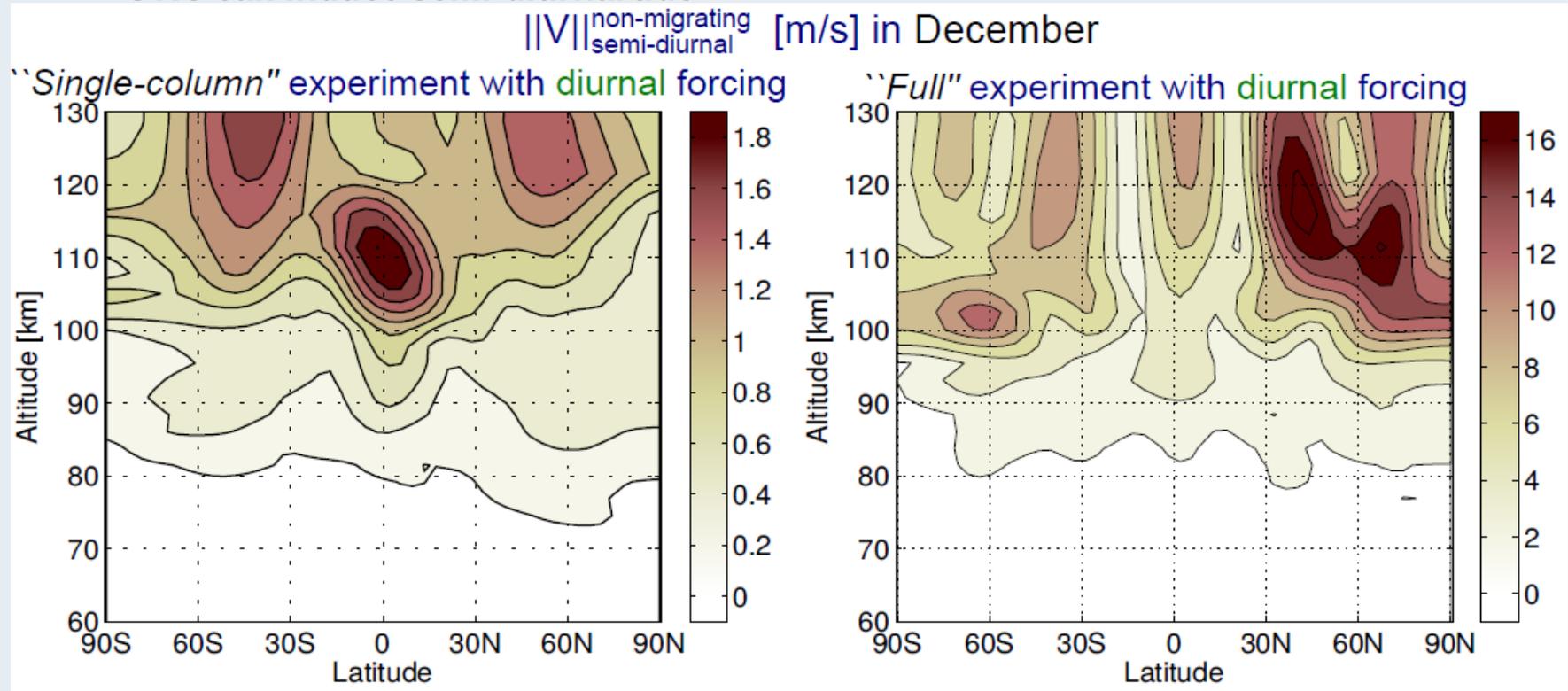
– Tidal model $dY'/dt = L(\bar{u}, \bar{T})Y' + \Re \sum_n Q_n e^{-in\Omega t} + F_{GW}(t)$

– Diurnal forcing only induces diurnal tide

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– **GWs can induce semi-diurnal tide**

– **GWs can induce semi-diurnal tide**



Tidal model in interaction with GWs (Ribstein et al 2015, Ribstein & Achatz 2016)

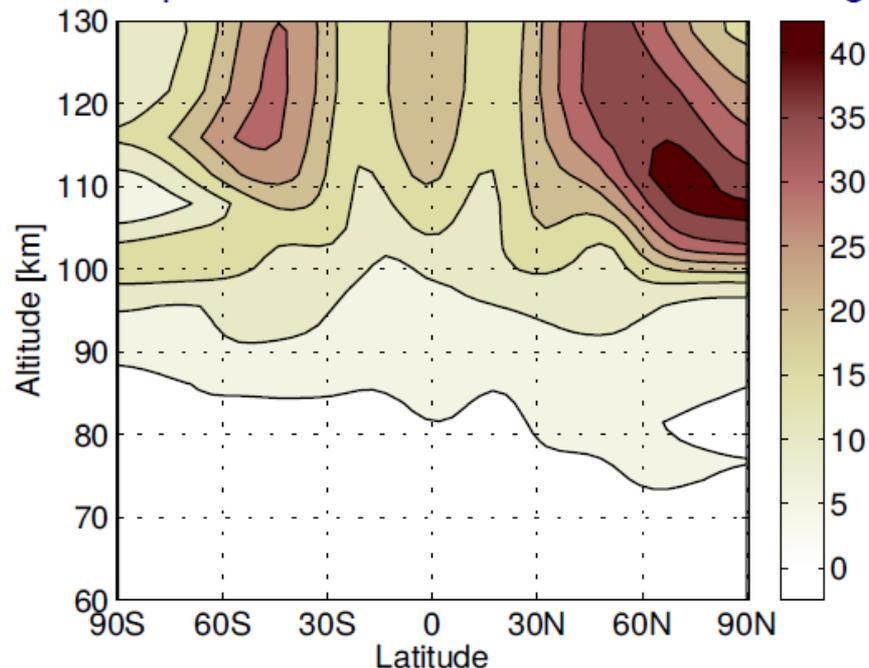
3D effects (beyond single column):

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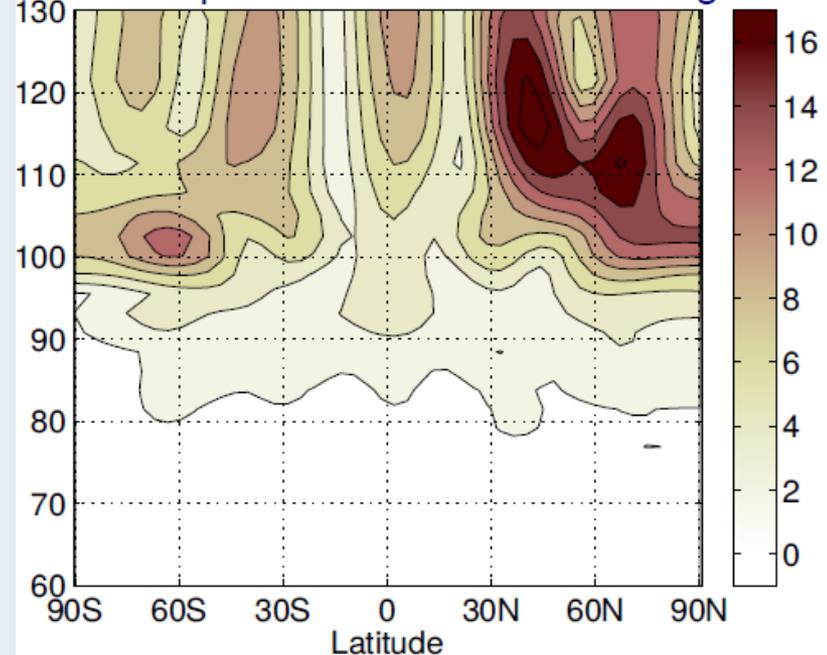
• Tidal model

- Tidal model $dY'/dt = L(\bar{u}, \bar{T})Y' + \Re \sum_n Q_n e^{-in\Omega t} + F_{GW}(t)$
- Diurnal forcing only induces diurnal tide
- Diurnal forcing only induces diurnal tide
- **GWs can induce semi-diurnal tide** (40% effect)
- **GWs can induce semi-diurnal tide** (40% effect)

“Full” experiment with semi and diurnal forcing



“Full” experiment with diurnal forcing



- **Approximations in present-day GW parameterizations critically limit their validity**
 - **Single-column**
 - **Steady state**
- **First implementation of a generalized approach into a global model**
- **Significant impact:**
 - **Zonal-mean forcing**
 - **Solar tides**

Achatz, U., Ribstein, B., Senf, F., and R. Klein 2016: The interaction between synoptic-scale balanced flow and a finite-amplitude mesoscale wave field throughout all atmospheric layers: Weak and moderately strong stratification. *Quart. J. Roy. Met. Soc.*, **143**, 342–361

Ribstein, B., Achatz, U. und F. Senf, 2015: The interaction between gravity waves and solar tides: Results from 4D ray tracing coupled to a linear tidal model, *J. Geophys. Res.*, **120**, doi:10.1002/2015JA021349

Bölöni, G., Ribstein, S., Achatz, U., Muraschko, J., Sgoff, C. und J. Wei, 2016: The interaction between atmospheric gravity waves and large-scale flows: an efficient description beyond the non-acceleration theorem. *J. Atmos. Sci.*, **73**, 4833-4852

Ribstein, B. und U. Achatz, 2016: Gravity wave propagation and impacts on a diurnal middle atmosphere : results from 4D ray tracing directly coupled to a linear tidal model. *J. Geophys. Res.*, doi:10.1002/2016JA022478

- Investigation **multi-scale dynamics of GWs** in 6 projects
- **prognostic WKB GW parameterization** to be developed for NWP and climate model
- To be addressed:
 - Sources
 - Propagation
 - dissipation
- Combined effort:
 - Theory,
 - modelling,
 - measurements,
 - laboratory experiments

