

HYDRAULIC REDISTRIBUTION BY PLANT ROOTS:

A Mechanism of Interaction between the Biosphere Moisture Reservoirs of the Deep Soil, the Near-Surface Soil, and the Lower Atmosphere

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1. INTRODUCTION

Objective:

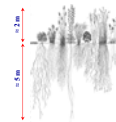
To model hydraulic redistribution and investigate its effect on (a) soil moisture profile and (b) fluxes at land-atmosphere interface in water-limited ecosystems.

Motivation:

- The long-memory nature of deep soil layer
- The availability of plant roots at deeper soil layers

Hypothesis:

"The hydrologically active depth of rooting zone in current climate models is underestimated and that the long-memory deep soil-layer may have significant impact on land-atmosphere interaction in vegetated environments."



3. MODEL

INPUT DATA	SOURCE
Atmos. Forcing	NARR
Soil Properties	Global Soil Data Task
Land Cover	IGBP-DIS
Vegetation Properties	AVHRR

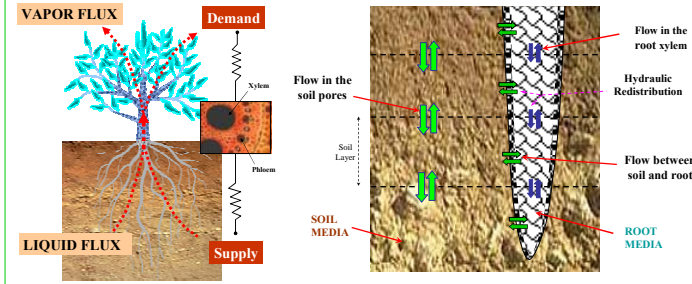


Fig.2: Conceptual representation of the soil-plant-atmosphere continuum (left) and the hydraulic redistribution model (right) as used in this study.

MODULE	PRINCIPAL ALGORITHM
Soil Model	Based on Darcy's Law $\frac{\partial \theta}{\partial t} + \frac{\partial}{\partial z} \left[K_{sat} \left(\frac{\partial \psi_{mat}}{\partial z} + 1 \right) \right] = -\rho_{mat} K_{mat} (\psi_{mat} - \psi_{atm})$
Root Model	Based on Hagen-Poiseuille's Law $\frac{\partial}{\partial z} \left[\rho_{mat} K_{mat} \left(\frac{\partial \psi_{mat}}{\partial z} + 1 \right) \right] = \rho_{mat} K_{mat} (\psi_{mat} - \psi_{atm})$
Stomatal Conductance	Jarvis Algorithm $g_s = g_{s,max} f(Rad) f(T) f(\psi_{mat}) f(D) f(C)$
Canopy Transpiration	Penman-Monteith Algorithm $E = \frac{\Delta \lambda Q_n + \rho_a c_p (e_s - e_a) g_s}{\Delta \lambda g_s + \gamma (g_s + g_a)}$
Carbon Assimilation	Farquhar-Collatz Algorithm $A = \min(A_p, A_m)$

2. STUDY SITE

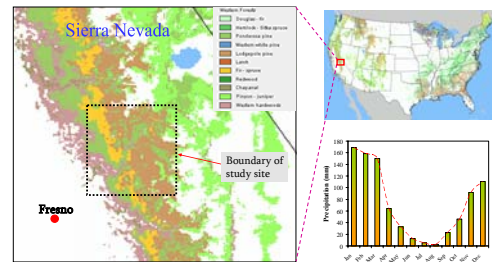


Fig.1: Location of case study site – Sierra Nevada ecosystem

Soil Depth & Root Distribution

The depth to center of i^{th} soil layer:
 $z_i = 0.025 (e^{0.5(i-0.5)} - 1)$

The root fraction in the i^{th} soil layer:
 $F_{root}(z_i) = -\Delta z \frac{c}{z_{50} + z_{50}} \left[1 + \left(\frac{z_i}{z_{50}} \right)^2 \right]^{-c} = \frac{1.27875}{\log \left(\frac{z_{50}}{z_{10}} \right)}$

Where z_{50} and z_{10} are vegetation-dependent constants.

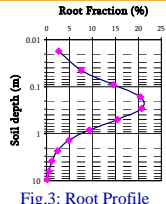


Fig.3: Root Profile

4. RESULTS

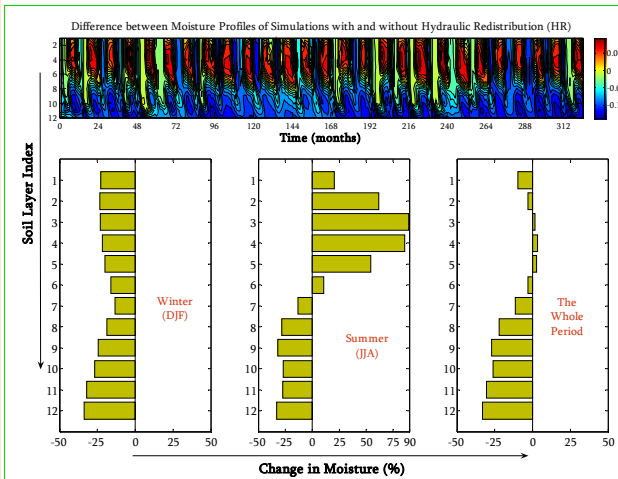


Fig.4: The effect of Hydraulic Redistribution (HR) on soil moisture profile. Positive (negative) values indicate an increase (decrease) in soil moisture due to HR.

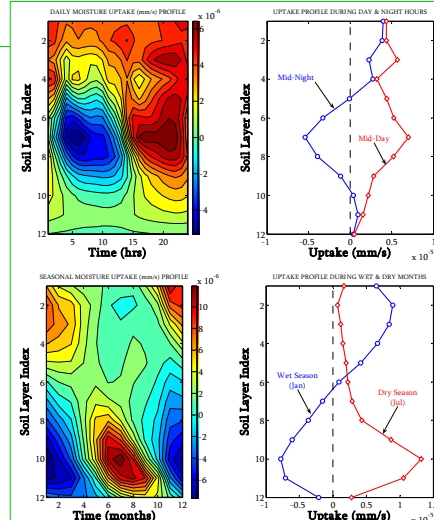


Fig.5: Daily (top) and Seasonal (bottom) profiles of water uptake by vegetation root system.

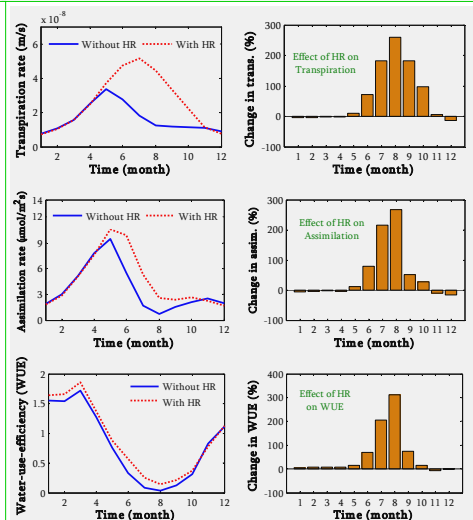


Fig.6: The effect of HR on Transpiration (top), Photosynthesis (middle), & Water-Use-Efficiency (bottom).

5. CONCLUSIONS

- Hydraulic redistribution (HR) – a passive transport of soil water across soil layers via plant roots – could be modeled by assuming the plant root system as a conduit for moisture transport along a pressure gradient, analogous to pipe flow.
- HR coupled with rooting depth could be a main mechanism for facilitating the dynamic interactions between moisture reservoirs of the long-memory deep soil, the medium-memory near-surface soil, and the short-memory atmosphere.
- The incorporation of HR results in a significant (i) alteration in the profiles of soil moisture and moisture uptake, and (ii) increase in dry-season transpiration, carbon assimilation, and water-use-efficiency.
- By enhancing the effectiveness of water uptake by plant roots from long-memory deep soil moisture reservoirs (and hence, controlling the water, energy, and carbon cycles), the dynamic mechanism of HR may have an important implication for understanding processes that govern long-term climate and ecological predictions.

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