Observational and diagnostic studies for AMMA

[land-surface-atmosphere feedbacks]

Alan K. Betts Atmospheric Research, Pittsford, VT <u>akbetts@aol.com</u>

Silver Springs, May 5, 2006

Background references

- Betts, A. K and P. Viterbo, 2005: Land-surface, boundary layer and cloud-field coupling over the south-we stern Amazon in ERA-40. *J. Geophys. Res.*, 110, D14108, doi:10.1029/2004JD005702.
- Betts, A.K., J.H. Ball, A.G. Barr, T.A. Black, J.H. McCaughey and P. Viterbo, 2006: Assessing landsurface-atmosphere coupling in the ERA-40 reanalysis with boreal forest data. *Agric. For. Meteorol.(accepted)*

(*Preprint: <u>ftp://members.aol.com/akbetts/BettsetalFCRN26.pdf</u>)*

• Betts, A. K., 2006: Radiative scaling of the nocturnal boundary layer and the diurnal temperature range, *J. Geophys. Res.*, 111, *D07105*, *doi:10.1029/2005JD006560*

What does AMMA need?

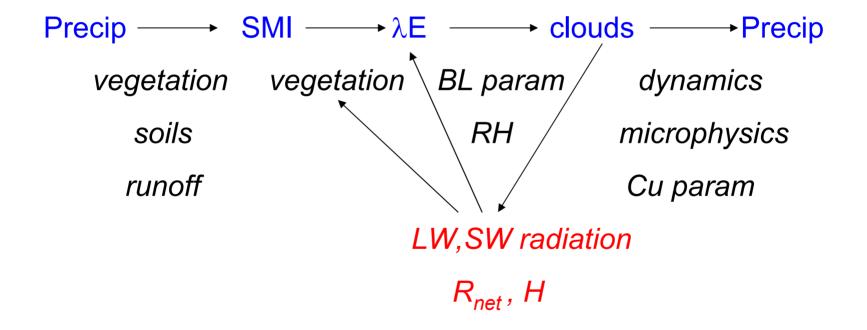
Models that are properly coupled across all scales!

- Point and spatial evaluation of models
- Evaluation of model land-surfaceatmosphere coupling against data
- Conceptual understanding of landsurface-atmosphere coupling!

Clouds are a crucial link in landsurface feedback

- Cloud fields are a tightly coupled component; impacting surface energy budget and evaporation
- Partly linked locally to 'soilwater' which impacts evaporation, and LCL
- Partly linked to larger-scale dynamics

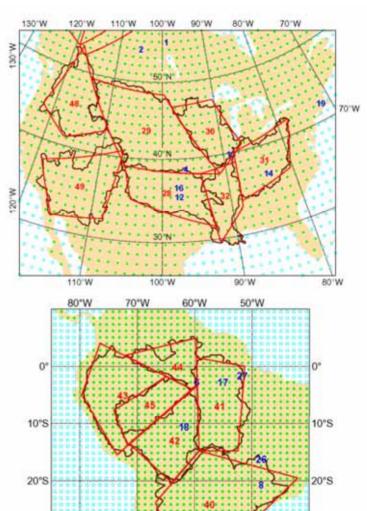
Consider the chain of processes involving water



SMI : soil moisture index [0<SMI<1 as PWP<SM<FC] α_{cloud}: 'cloud albedo' viewed from surface

ERA40 river basin budgets

- Basin averages: hourly archive
- Daily averages
- Madeira : Amazon: 1990-2001
- Mississippi: 1980-2002



30°S

80°W

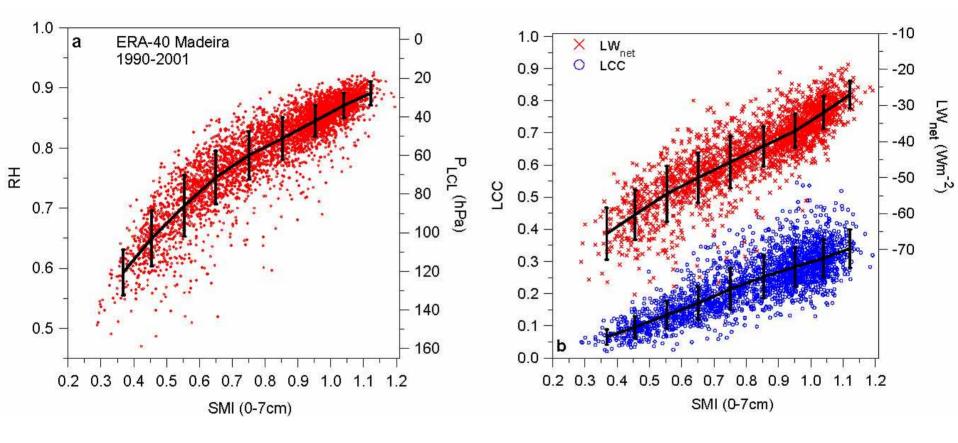
70°W

60°W

50°W

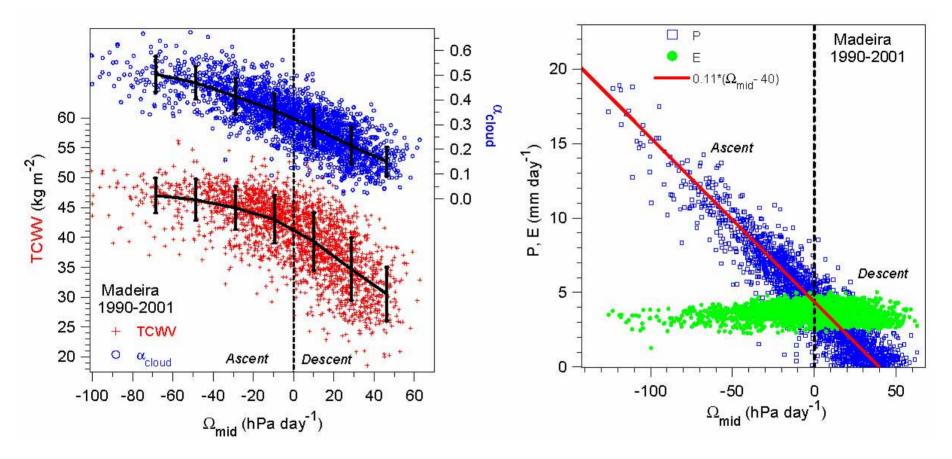
30°S

ERA40: Surface 'control'



- Madeira river, SW Amazon
- Soil water → LCL, LCC and LW_{net}

ERA-40 dynamic link (mid-level omega)



• $\Omega_{mid} \rightarrow Cloud albedo, TCWV and Precipitation$

How well are physical processes represented?

• Basin-scale assessment of ERA40 biases [Betts et al. 2003a, 2003b, 2005]

• Flux tower data can assess both biases and the coupling of physical processes *on the point scale*

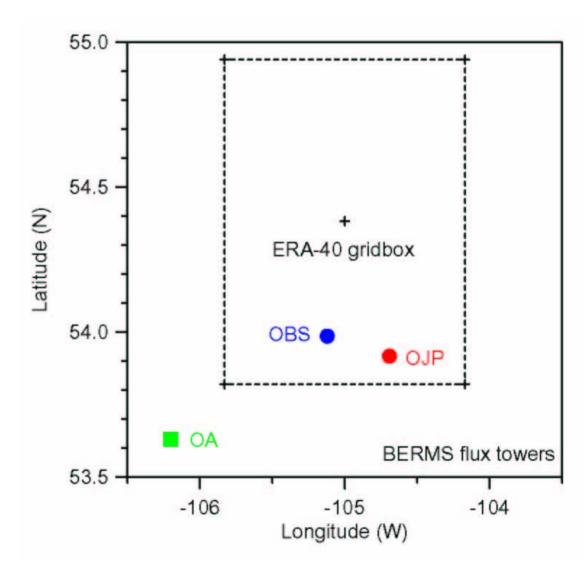
Compare ERA-40 with BERMS flux towers in Saskatchewan

Focus:

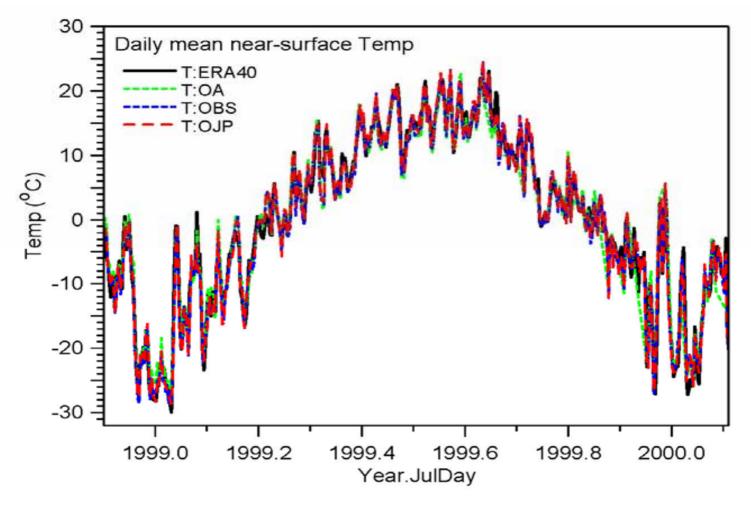
- Coupling of clouds to surface fluxes
- Define a 'cloud albedo' that reduces the shortwave (SW) flux reaching surface
 - Basic 'climate parameter', coupled to surface evaporation [locally/distant]

Compare ERA-40 with BERMS

- ECMWF reanalysis
- ERA-40 hourly time-series from single grid-box
- BERMS 30-min time-series from Old Aspen (OA)
 Old Black Spruce (OBS)
 Old Jack Pine (OJP)
- Daily Average

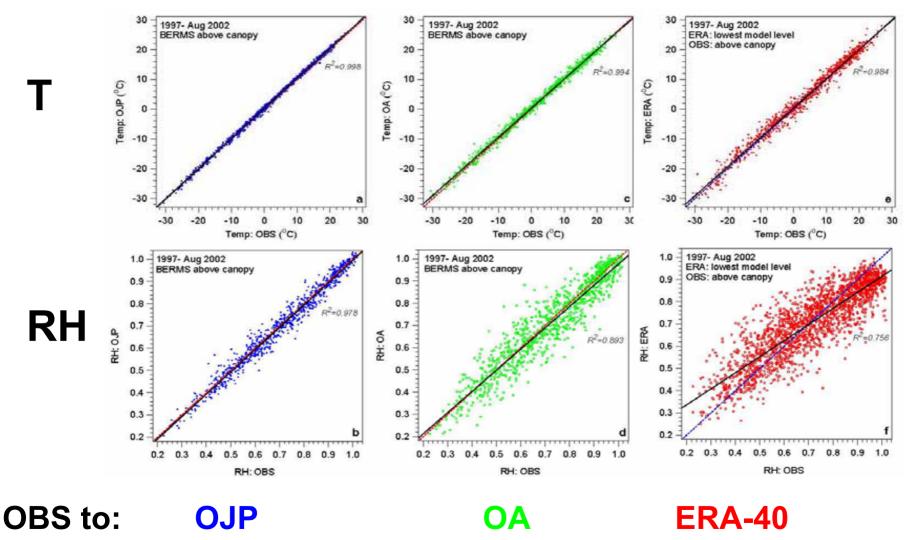


Global model improvements [ERA-40]



- Reanalysis T bias is now small in all seasons [ERA-40 land-surface model developed from BOREAS]
- BERMS inter-site variability of daily mean T is small

Comparison of BERMS and ERA-40



81km

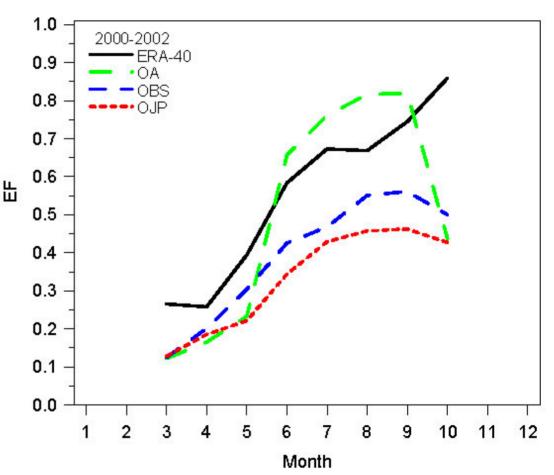
[grid-point]

Spacing:

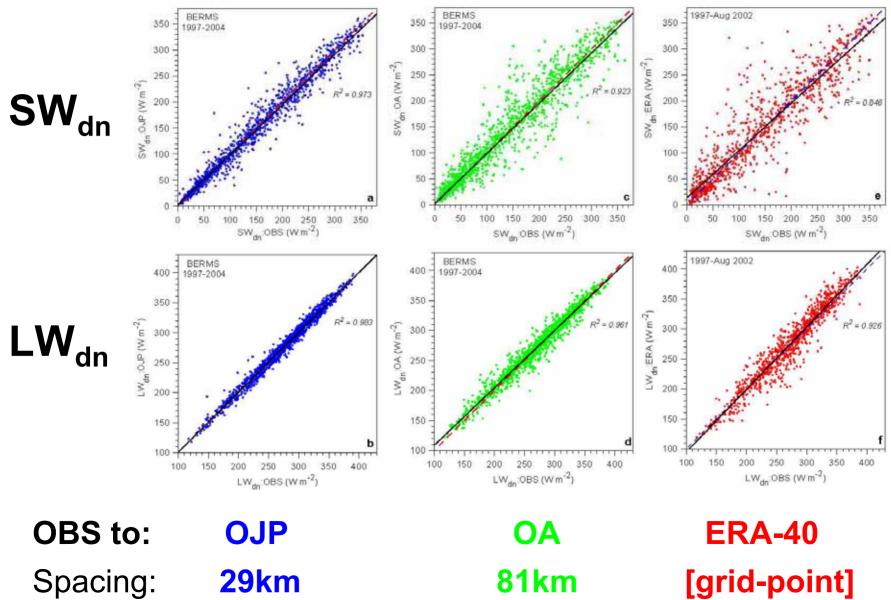
29km

Seasonal Evaporative Fraction

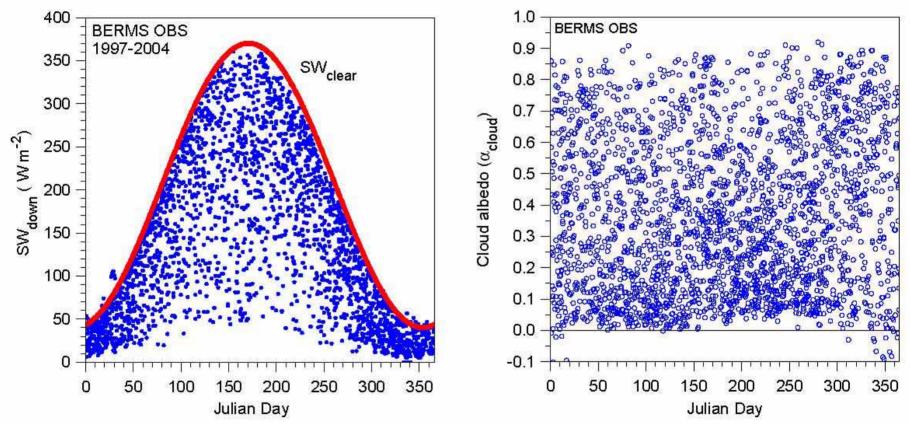
- Data as expected OA>OBS>OJP
- ERA-40 too high in spring and fall
 [Lacks vegetation seasonal cycle]
- ERA a little high in summer?



Comparison of BERMS and ERA-40

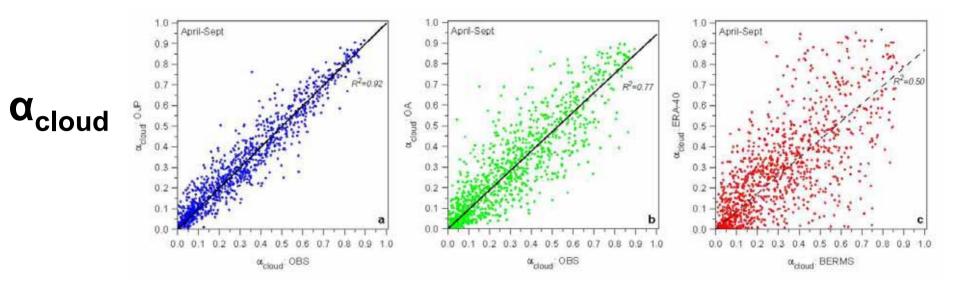


BERMS: Old Black Spruce



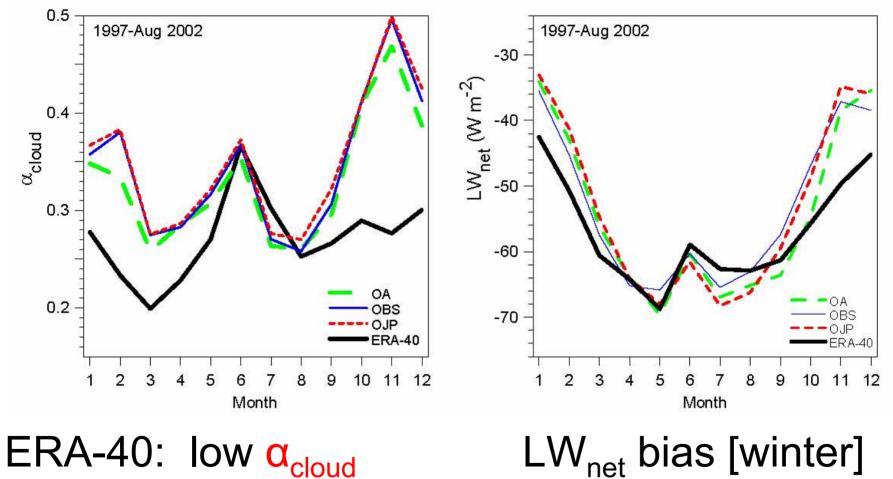
- Cloud 'albedo': $\alpha_{cloud} = 1 SW_{down}/SW_{clear}$
- $SW_{net} = (1 \alpha_{cloud})(1 \alpha_{surface})SW_{clear}$

Cloud albedo comparison (daily)



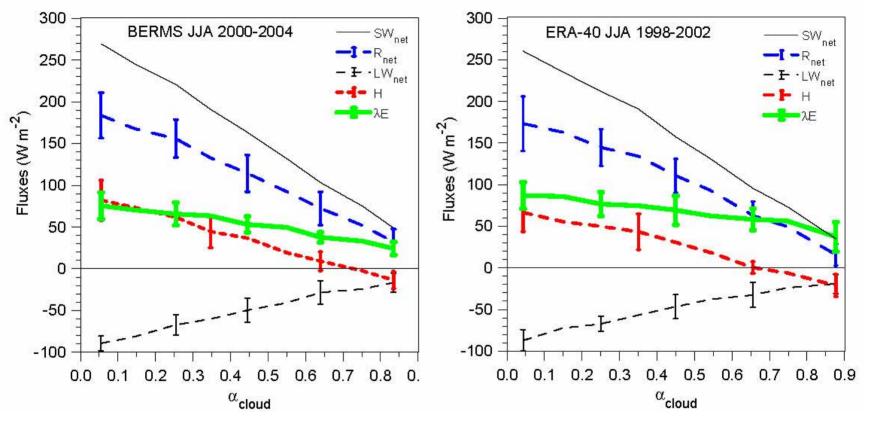
OBS to:	OJP	ΟΑ	ERA-40
Correlation:	Good	Fair	Poor
Spacing:	29km	81km	[grid-point]

Cloud albedo and LW comparison



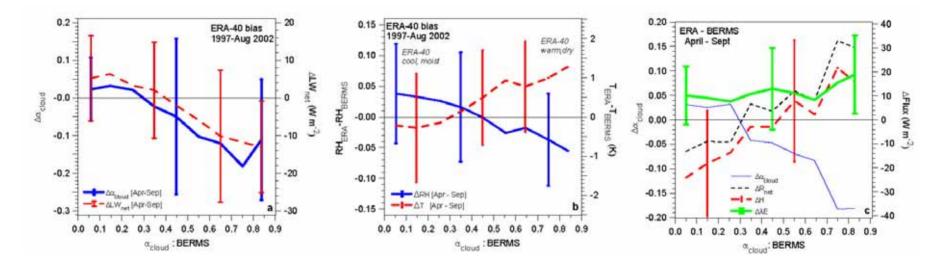
[except summer]

How do fluxes depend on cloud cover?



- Quasi-linear variation
- Evaporation varies less than other fluxes

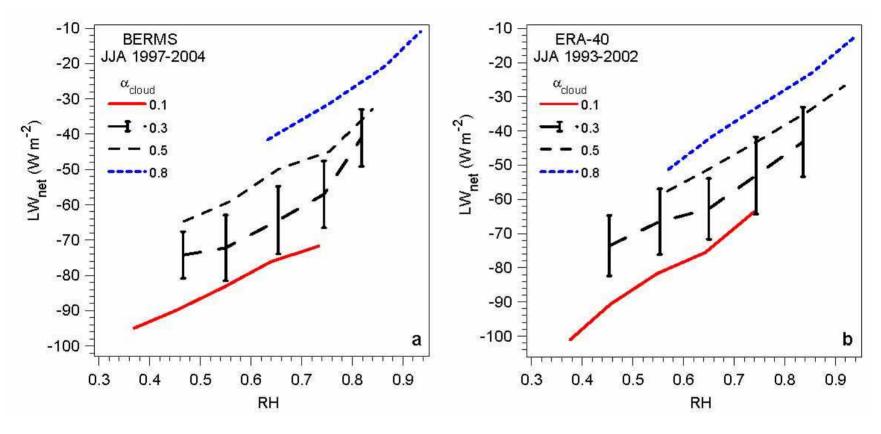
How do model biases depend on cloud cover?



Across range of observed cloud cover shift from

- cloudy-cool-moist bias to warm-dry-less-cloud bias
- H and R_{net} [but not λE] coupled to cloud bias

 LW_{net} on RH and α_{cloud}



- Outgoing LW_{net} falls as RH and cloud cover increase
- Higher RH means lower LCL & depth of ML
- LW coupling same for BERMS and ERA-40
- *LW_{net} is linked to diurnal temperature range [Betts, 2006]*

Coupling of regional water vapor convergence, clouds and landsurface processes

Alan K. Betts Atmospheric Research, Pittsford, VT <u>akbetts@aol.com</u>

A03 Role of Circulation in Regional Hydroclimate Variability AGU, May 23, 2006

Historical perspective

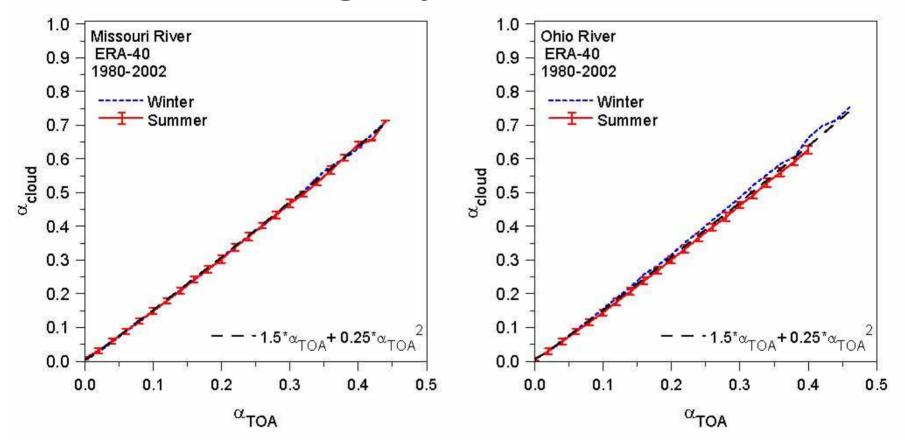
- For about 15 years, 'cloud feedbacks' have been labeled a 'challenge'; a 'major source of uncertainty in climate modeling'
- Why? Seems odd because they are so easily observed!
- A quantitative framework, which links clouds to both surface and large-scale processes has been missing.

Organize data by 'cloud albedo' and the links become transparent and verifiable

Definitions

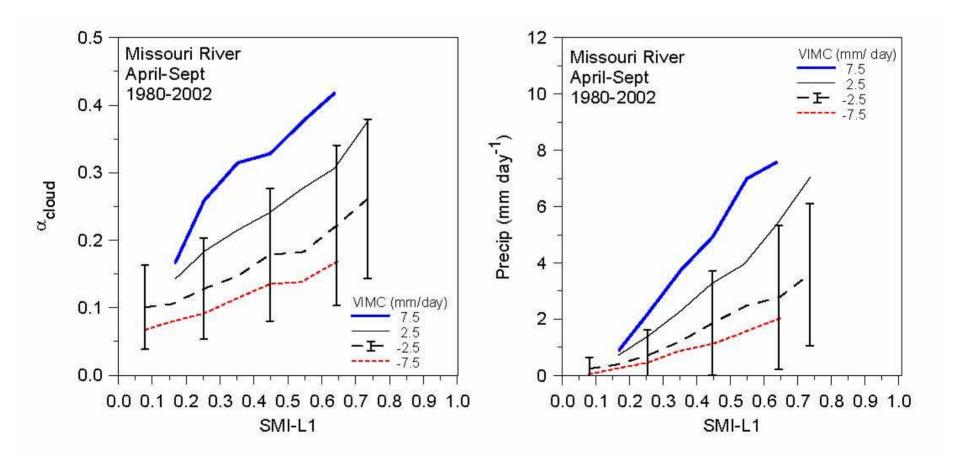
- VIMC: Vertically integrated moisture convergence
- α_{cloud}: 'cloud albedo' viewed from surface – measure of surface SW cloud forcing
- SMI : soil moisture index
 [0<SMI<1 as PWP<SM<FC]

Surface and TOA cloud albedo are tightly related

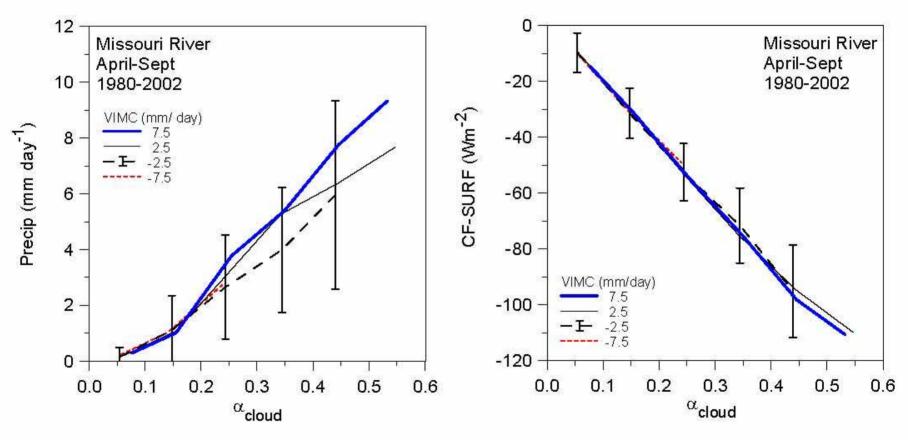


- α_{cloud} = SWCF:SURF/SW_{net}(clear)
- α_{TOA} = SWCF:TOA/SW_{dn}(clear)

α_{cloud}, Precipitation increase with SMI and VIMC



α_{cloud} is the critical link

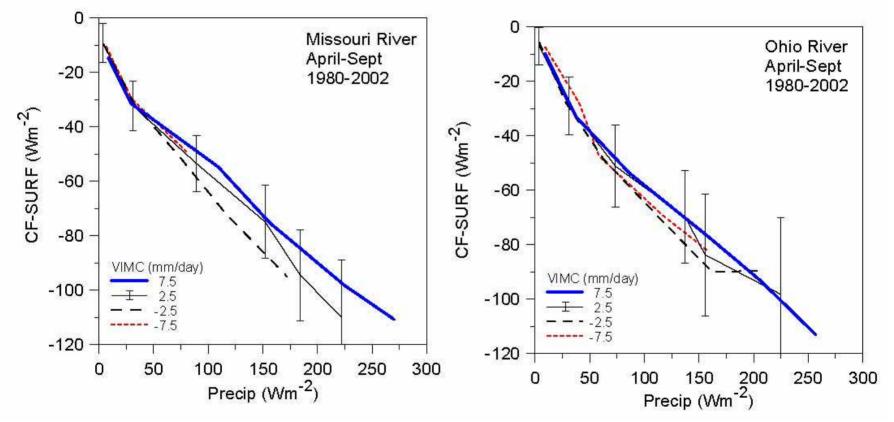




and surface cloud radiative forcing

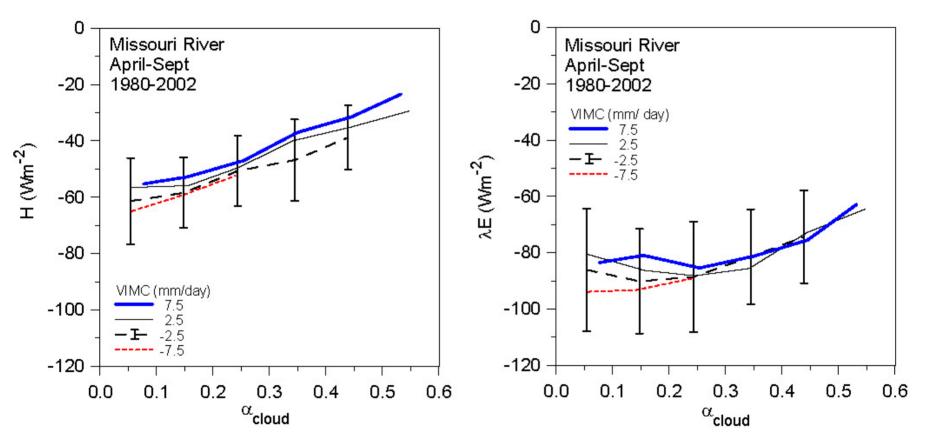
Not dependent on moisture convergence, VIMC

Relation of surface cloud forcing to precipitation forcing [Wm⁻²]



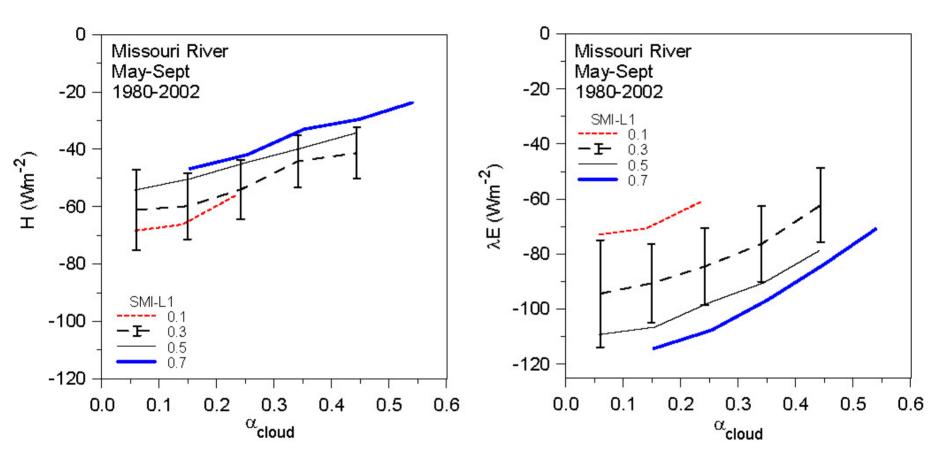
- Same relationship, not dependent on VIMC
- CF-SURF is ≈ 40% of precipitation *in ERA-40*

Surface fluxes vary with α_{cloud} but not VIMC



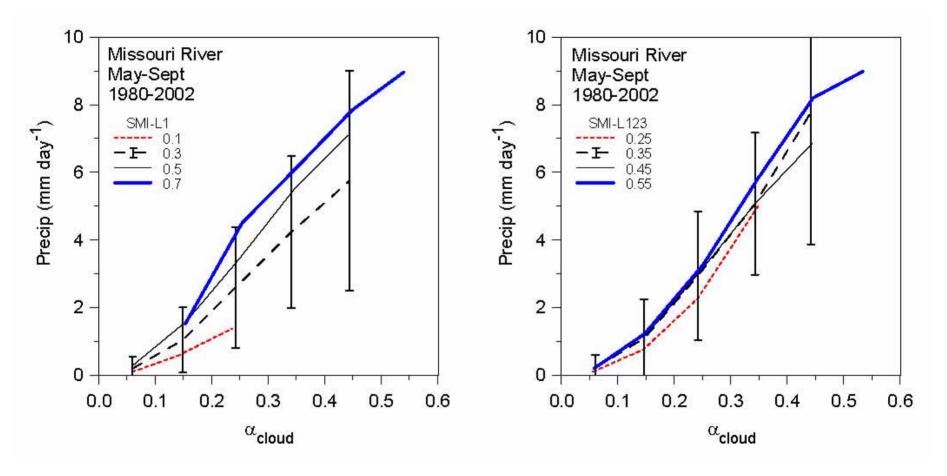
• H varies more with α_{cloud} than λE

Surface fluxes vary with SMI



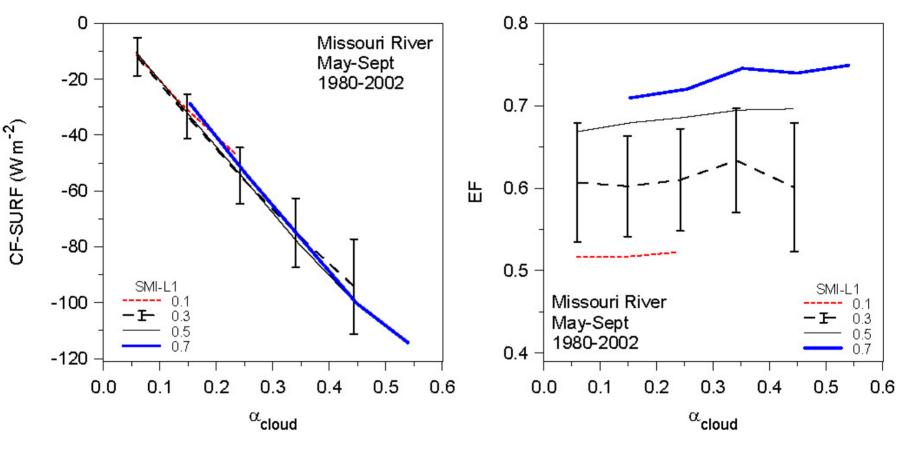
L123 root-zone soil moisture similar

Precipitation varies with α_{cloud} and SMI-L1



Coupling to SMI-L1 stronger than SMI-L123

ERA-40 SEB is partitioned into



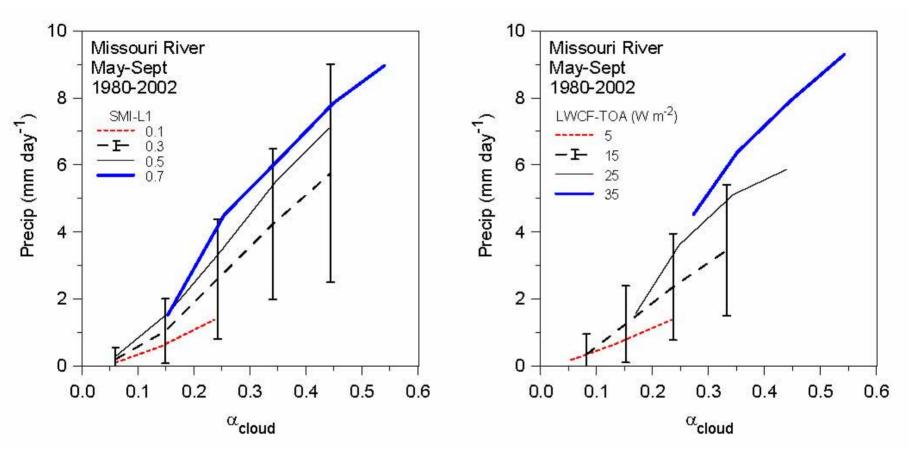
- Energy, a function of cloud albedo
- EF, a function of SMI

Satellite perspective on SEB

SEB energy balance a 'soluble problem' ?

Surface cloud forcing/α_{cloud} [visible]
 EF from surface layer SMI [microwave]
 Vegetation, slower component [NDVI]

Satellite perspective on Precip



Dependence of precipitation on α_{cloud} and SMI-L1 LWCF-TOA

Conclusions

- Organize data by 'cloud albedo'
- Cloud albedo is as important as surface albedo [with higher variability]
- Clouds, BL and surface are a coupled system; so biases coupled
- Coupling of LW_{net} to RH and α_{cloud} important to diurnal temperature range

Critical Issues

- Are observables coupled correctly in a model? Accuracy of model 'daily climate' important to forcing of monsoon
- Key non-local observables:
 - BL quantities: RH, LCL, θ_E linked to soilwater
 - Clouds: reduce SW reaching surface, α_{cloud} and outgoing LW_{net}
 - Aerosol: shift heating from surface to lower atmosphere

For AMMA

- Will we have sufficient/representative fluxtower data to assess models?
- Bamba (17.0N, 1.4W)
- Agoufou (15.33N, 1.48W)
- Wankama (13.63N, 2.63E)
- Banizoumbou (13.52N, 2.62E) [+ Niamey airport- ARM]
- Djougou (9.82N, 1.72E)
 [2 or 3 flux stations at each location each over a different land surface type]
- Can we link 'cloud albedo' derived from satellite data to surface fields and model analyses?
- Stratify by aerosol surface radiative feedback